Study of macropore flow using electrical resistivity tomography?

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Outline

- I. Summary
- II. ERT principle and layout
- III. Inversion and petrophysical relationship
- IV. Errors with regard to quantitative calculation?
- V. The test site Holtensen (loess)
 - \checkmark sprinkling experiment, the process, the problems
 - ✓ ponding experiment, the process, the problems
- VI. Conclusion and outlook



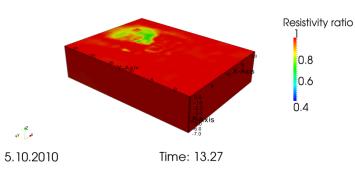


Summary

ERT is a wonderful method to observe flow processes minimal invasively ERT is easy, everybody can use it, and it is relatively cheap Flow paths can be visualised, but mind the resolution constraints Quantitative interpretation in unsaturated condition is still a challenge Try to avoid conductive tracers



Water movement during and after sprinkling



(collaboration with TU Delft, Bogaard et al.)

all inversions done using the programme Bert (Günther et al., 2006, www.resistivity.net)

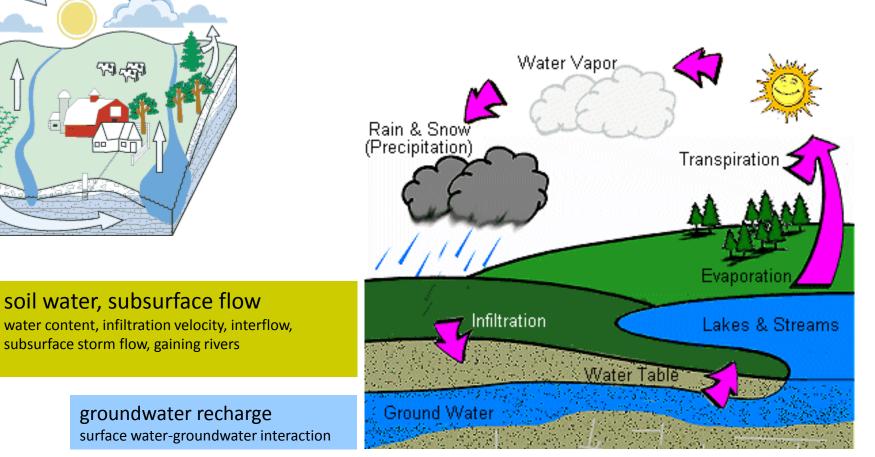




Motivation

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Can we use ERT to directly observe those hidden subsurface processes on a meso scale (1-100m²)?



www.dnr.state.wi.us/.../images/groundwater.gif (21.6.07)



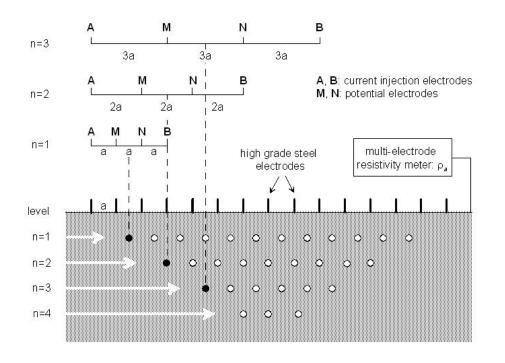


The method (ERT = electrical resistivity tomography)

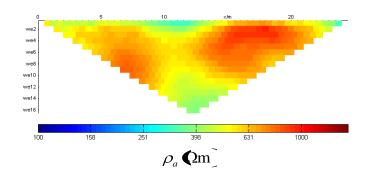
sketch (Wenner-Configuration):

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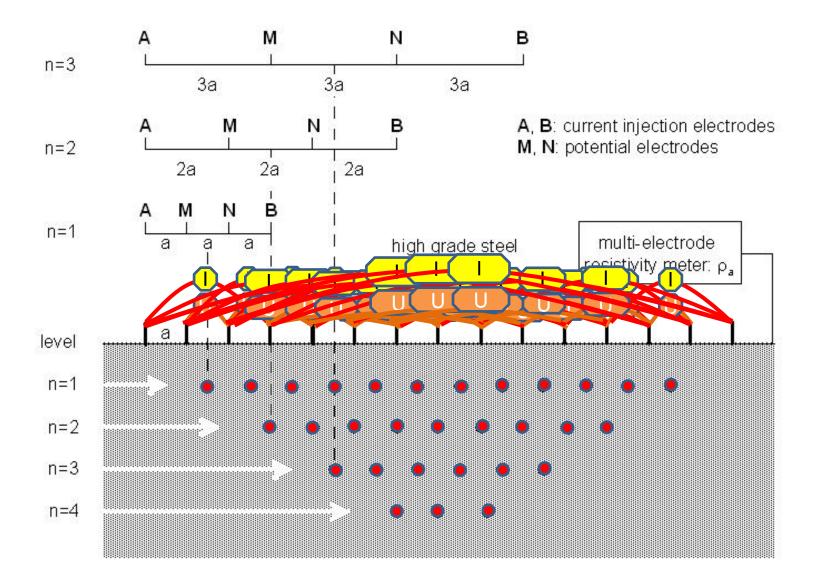
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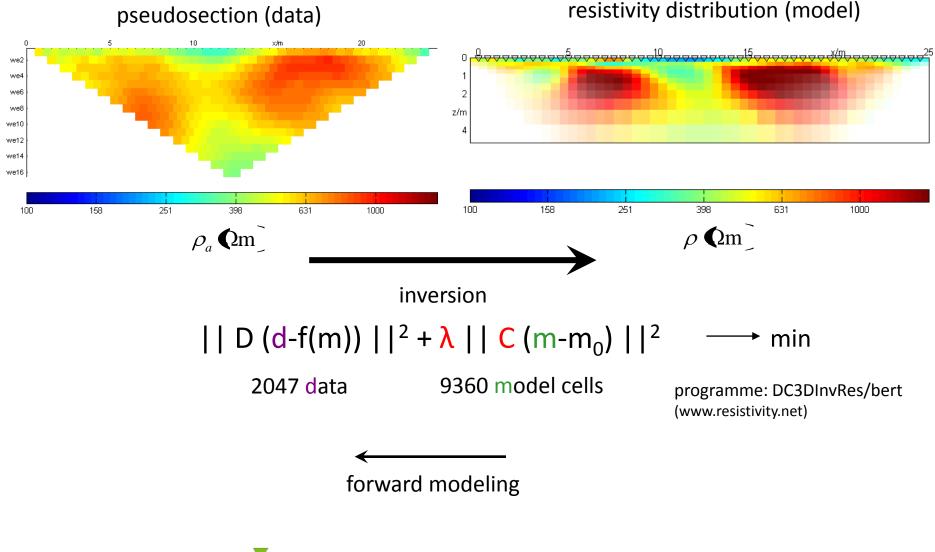
measurements visualised as: pseudo section







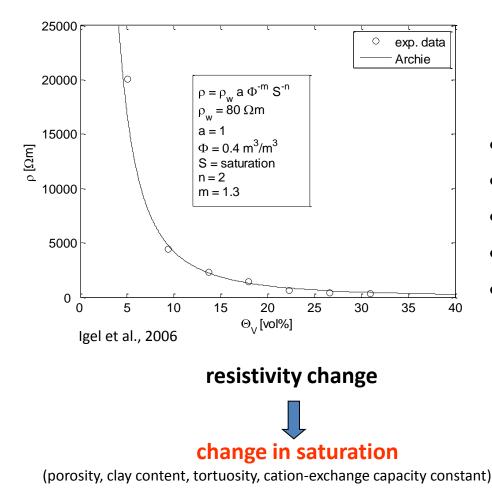
Data inversion







Petrophysical relationship between resistivity and water content



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$$\rho_t = \frac{a}{\Phi^m} \rho_w \frac{1}{S_w^n}$$

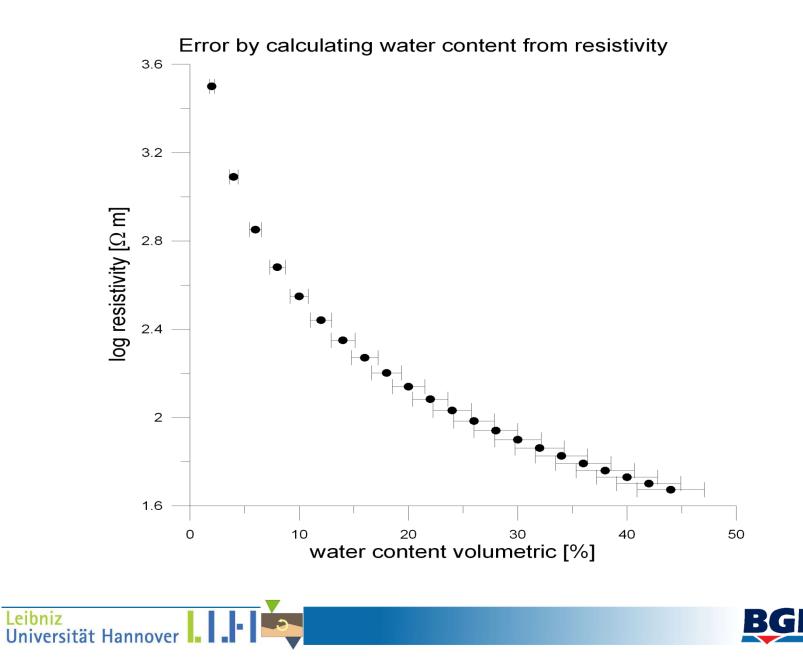
(Archie-equation)

- ρ_t : bulk resistivity
- Φ: porosity
- P_w : resistivity of the pore fluid
- S_w : saturation
- *m*, *n*, *a*: empirical constants

 $\frac{\rho_1}{\rho_0} = \left(\frac{S_0}{S_1}\right)'$

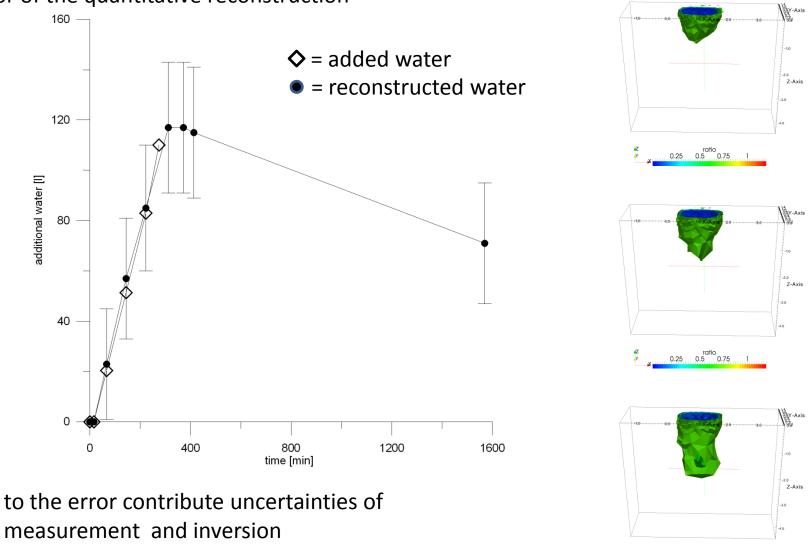
 $\Delta S = \left(\sqrt{\frac{\rho_0 * k}{\rho_1}} - 1\right) * S_0 \qquad \qquad k = \frac{\rho_{w1}}{\rho_{w0}}$





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Error of the quantitative reconstruction



Archie parameters m and n



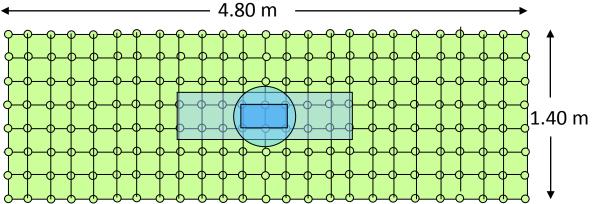
fluid conductivity



ratio 0.5 0.75

0.25

Two infiltration experiments in loess (sprinkling, ponded infiltration)









infiltration area:

1m x 0.4m / 80 l in 48 h (4.2 mm/h) 0.4 m diameter/ 25l in 24 min

array 150 electrodes (vertical or horizontal) dipole-dipole + wenner configuration 580 /1560 measurements within 5/15 min

excavation of central areas measurements of water content (TDR) and matric potential day(s) after infiltration.





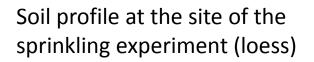


0-13 cm: Ap 1 13-30 cm: Ap 2 30-53 cm: Al (?) 53-77 cm: Bt 77-86 cm: Bv/Cv 86-103 cm: Cv (?)

lay out of the 150 electrodes

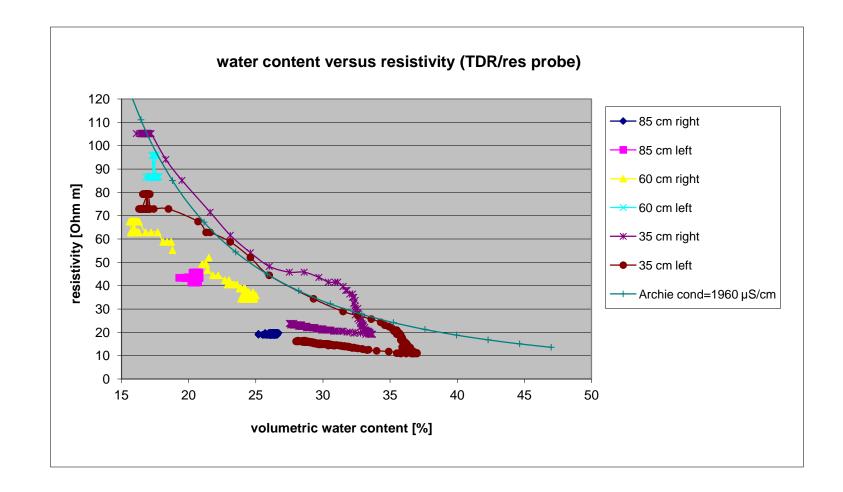
3D array, 4.8 x 1.4 m electrode distance 20 (40) cm infiltration area: 1 x 0.4 m brilliant blue tracer 80 I within 48 hrs

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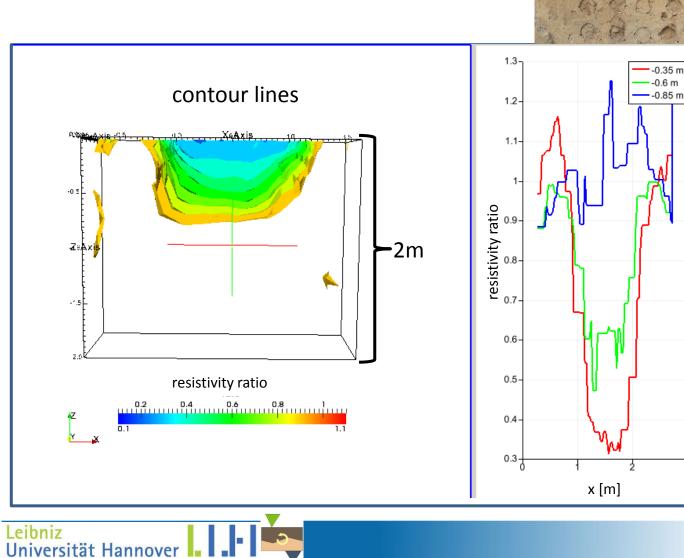


Results of the TDR measurements at different depths and positions







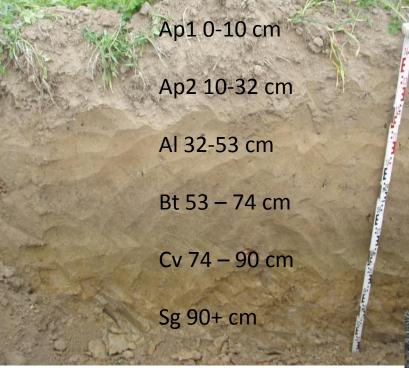


Section through the central sprinkling area

ERT inversion result central section 3 days after start of sprinkling (48 hrs)

1m





Soil near the test site, second profile

lay out of the 6 vertical electrodes

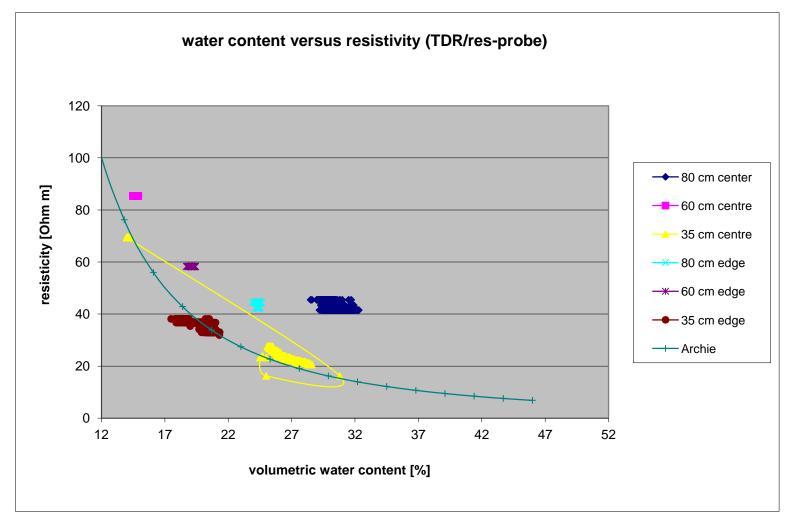
1.3 m long, 25 electrode rings each diameter of the ponding ring: 40 cm brilliant blue tracer25 l within 24 minutes







Results of the TDR measurements at different depths and positions

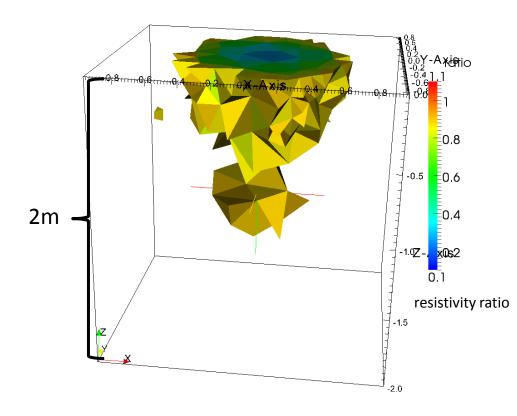


not yet corrected acc. Bechthold et al., SSSAJ: Volume 74: Number 2 • March–April 2010





Inversion result after 34 minutes









Conclusion

- ✓ ERT is a wonderful method to observe flow processes minimal invasively
- \checkmark ERT is easy, everybody can use it, and it is relatively cheap
- ✓ Flow paths can be visualised, but mind the resolution constraints
- ✓ Quantitative interpretation in unsaturated condition is still a challenge
- ✓ Try to avoid conductive tracers

Outlook

For quantitative assessment the change in pore water conductivity needs to be observed independently



