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Bundesanstalt für Geowissenschaften und Rohstoffe

## Airborne Geophysical Investigations of CLIWAT Pilot Areas

## Survey Area <br> Vojens, Denmark, 2009

CLIMATE \& WATER
Interreg IVB Project:
CLIWAT - Adaptive and sustainable water management and protection of society and nature in an extreme climate


# BCR Bundesanstalt für Geowissenschaften und Rohstoffe Federal Institute for Geosciences and Natural Resources 

CLIWAT - Adaptive and sustainable water management and protection of society and nature in an extreme climate Survey Area Vojens, Denmark

2009

## Technical Report on the Interreg IVB Project

In Cooperation with Region Syddanmark Vejle, Denmark


## Region Syddanmark

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| Abbreviations |  |
| :---: | :---: |
| - | degree |
| ${ }^{\circ} \mathrm{C}$ | degree Celsius |
| ' | minute |
| " | second or inch |
| \% | per cent |
| 1-D | one-dimensional |
| a | aircraft background |
| A | amplitude of measured HEM components |
| $\mathrm{A}_{c}, \mathrm{~A}_{c}^{\prime}$ | amplitudes of calculated HEM components |
| $\mathrm{A}_{\mathrm{p}}$ | polynomial approximation of $\mathrm{A}_{c}^{\prime}(\delta)$ |
| Ah | ampere hours |
| agl | above ground level |
| asl | above mean sea level |
| $\alpha, \beta, \gamma, \mathrm{a}$ | stripping ratios |
| $\alpha_{e}, \beta_{e}, \gamma_{e}$ | height corrected stripping ratios |
| $\alpha_{0}$ | complex wave number |
| b | cosmic stripping factor |
| bgl | below ground level |
| BGR | Bundesanstalt für Geowissenschaften und Rohstoffe |
| Bi | Bismut |
| $\mathrm{B}_{\mathrm{n}}$ | layer admittance |
| C | concentration |
| $\mathrm{C}_{0}$ | element concentration at ground |
| $\mathrm{C}_{\mathrm{H}}$ | element concentration in presence of vegetation |
| CF | compact flash |
| ch | channel number |
| $\mathrm{c}_{1}$ | effective cable length |
| cps | counts per second |
| Cs | Cesium |
| © | copyright |
| $\mathrm{d}_{\mathrm{a}}$ | apparent depth |
| $\mathrm{D}_{\mathrm{a}}$ | apparent distance |
| DC | direct current |
| DEM | digital elevation model |
| DGPS | Differential Global Positioning System |
| DK | Denmark |
| DVD | Digital Versatile Disc |
| $\delta$ | inverse relative skin depth ( $=\mathrm{h} / \mathrm{p}$ ) |
| $\delta_{\text {p }}$ | polynomial approximation of $\delta\left(\varepsilon_{c}\right)$ |
| $\delta_{\text {T }}$ | residual (magnetics) |

$\Delta h \_l \quad$ reduced laser altitude
$\Delta \mathrm{I} \quad$ zero-level error of in-phase component
$\Delta \mathrm{Q} \quad$ zero-level error of quadrature component
$\Delta \mathrm{T} \quad$ anomalies of the total magnetic field
$\Delta \mathrm{T}_{\mathrm{f}} \quad$ anomalies of the total magnetic field, high-pass filtered
$\Delta \mathrm{V} \quad$ diurnal (magnetic) variations
E east
E energy
E ground level exposure rate
e base of the natural logarithm ( $1 / \mathrm{e} \approx 0.37$ )
eTh equivalent concentration of Thorium
eU equivalent concentration of Uranium
EM electromagnetic(s)
ERDF European Regional Development Fund
EU European Union
$\varepsilon \quad$ ratio of measured HEM components (= Q/I)
$\varepsilon_{c} \quad$ ratio of calculated HEM components ( $=\mathrm{Q} / \mathrm{I}$ )
$\varepsilon_{0} \quad$ permittivity of air: $8.854 \times 10^{-12} \mathrm{As} / \mathrm{Vm}$
$\varepsilon_{\mathrm{n}} \quad$ layer permittivity
f frequency
F IRGF
FAS Fugro Airborne Surveys
FFT Fast Fourier Transform
ft feet
G gain constant
GBA Geologische Bundesanstalt
GPS Global Positioning System
h bird altitude
$\mathrm{H} \quad$ thickness of vegetation
HCP horizontal coplanar
$h_{e} \quad$ effective height
$\mathrm{h}_{0} \quad$ nominal survey height
HEM helicopter-borne electromagnetic(s)
HMG helicopter-borne magnetic(s)
HRD helicopter-borne radiometric(s)
h_GPS GPS-Höhe
h_l laser altitude
$h_{-} r, h_{r} \quad$ radar altitude
Hz hertz
i counter
I in-phase component (real part) of the HEM data

| $\mathrm{I}_{\text {c }}$ | calculated in-phase value |
| :---: | :---: |
| IAEA | International Atomic Energy Association |
| IAGA | International Association of Geomagnetism and Aeronomy |
| IGRF | International Geomagnetic Reference Field |
| $\mathrm{J}_{0}$ | Bessel function of first kind and zero order |
| K | degree Kelvin |
| K | Potassium |
| keV | kilo electron volts |
| kg | kilogram |
| kHz | kilohertz |
| km | kilometre |
| km/h | kilometres per hour |
| 1 | litre |
| $\log$ | logarithm |
| $\lambda$ | wave number |
| m | metre |
| MeV | mega electron volts |
| $\mu$ | attenuation coefficient (vegetation or height) |
| $\mu_{0}$ | permeability of air: $4 \pi \times 10^{-7} \mathrm{Vs} / \mathrm{Am}$, |
| $\mu_{\mathrm{n}}$ | layer permeability |
| $\mu \mathrm{R} / \mathrm{h}$ | microroentgens per hour |
| n | number of frequencies |
| N | north |
| $\mathrm{n}, \mathrm{N}$ | raw, corrected count rate |
| NaI | sodium iodide |
| NASVD | noise adjusted singular value decomposition |
| NL | non-linear |
| NL | The Netherlands |
| nT | nanotesla |
| $\mathrm{N}_{\mathrm{m}}{ }^{\prime}$ | observed count rate at STP effective height |
| $\mathrm{N}_{S}$ | corrected count rate at nominal survey height |
| $\mathrm{N}_{\mathrm{x}}$ | background and STP corrected count rates ( $\mathrm{x}=\mathrm{K}, \mathrm{U}, \mathrm{Th}$ ) |
| $\mathrm{N}_{\mathrm{x} \text { (corr) }}$ | stripping corrected count rates ( $\mathrm{x}=\mathrm{K}, \mathrm{U}, \mathrm{Th}$ ) |
| $\Omega \mathrm{m}$ | ohm metre ( $\mathrm{Ohm}^{*} \mathrm{~m}$ ) |
| p | skin depth |
| P | barometric pressure |
| $\mathrm{P}_{0}$ | barometric pressure at sea level |
| PDF | Portable Document Format |
| ppm | parts per million |
| $\pi$ | Pi (=3.14159265...) |
| Q | quadrature or out-of-phase component (imaginary part) of the |


| $\mathrm{Q}_{\text {c }}$ | calculated quadrature value |
| :---: | :---: |
| r | distance parameter |
| $\mathrm{R}_{1}$ | reflexion factor |
| $\mathrm{r}_{1}$ | conversion factor |
| $\rho$ | resistivity |
| $\rho_{0}$ | resistivity of air: > $10^{8} \Omega \mathrm{~m}$ |
| $\rho_{a}$ | apparent resistivity |
| S | south |
| S | sensitivity |
| s | second |
| STE | standard error |
| STP | standard pressure and temperature |
| t | thickness (of a model layer) |
| t | time variable |
| T | air temperature |
| $\mathrm{T}_{0}$ | temperature at freezing point of water on Kelvin scale |
| T, TMI | total magnetic field intensity |
| tanh | hyperbolic tangent |
| TC | total count |
| Th | Thorium |
| Tl | Thallium |
| $\mathrm{t}_{1}$ | life time |
| $\mathrm{T}_{\mathrm{LP}}$ | low pass cut-off period |
| U | Uranium |
| USA | United States of America |
| USB | Universal Serial Bus |
| UTC | Coordinated Universal Time |
| UTM | Universal Transverse Mercator Projection |
| V | volt |
| VCX | vertical coaxial |
| VRS | vertical resistivity section |
| W | west |
| WFD | Water Framework Directive |
| WGS | World Geodetic System |
| $\omega$ | circular frequency |
| X, Y, Z | Cartesian coordinates, Z depth axis |
| Z | relative secondary magnetic field |
| $\mathrm{z}^{*}$ | centroid depth |

## 1. Summary

Climate change simulations indicate a sea-level rise and increasing rainfall in the North Sea region leading to higher groundwater levels and a forced outwash of nutrients and pollutants from industrial areas, agriculture and landfills. CLIWAT (climate \& water) is a transnational Interreg project in the North Sea region funded by the European Union with partners from Belgium, The Netherlands, Germany and Denmark. The goal of the project is to determine the effects of a possible climate change on groundwater systems, surface water and the fresh-saltwater boundary in the North Sea and Baltic Sea region.

Geological and geophysical measurements were carried out in the seven pilot areas of the project. In order to map the existing groundwater structures with airborne geophysical methods the German Federal Institute for Geosciences and Natural Resources (BGR) conducted three surveys in Zeeland, Friesland (both NL) and Vojens (DK). One of these pilot areas covers parts of northern Schleswig and southern Jutland on both sides of the Danish-German border. The target of this pilot area, the Hørløkke landfill, is located to the west of the town of Vojens in the northern part of the pilot area. The public landfill covers an area of about $12,000 \mathrm{~m}^{2}$ where approximately $65,000 \mathrm{~m}^{3}$ of household and industrial waste were filled in from 1968 to 1972. As the water divide is just east of the landfill it is suspected that a rising groundwater table will divert the plume of contaminants to the east - directly towards the town of Vojens.

By request of the Danish project partner Region Syddanmark a helicopter-borne survey of the area around Hørløkke landfill was conducted by the BGR airborne group in June 2009. The airborne survey comprises a 6.5 km by 8 km wide area ranging from $9^{\circ} 11^{\prime} 49^{\prime \prime} \mathrm{E}$ to $9^{\circ} 19^{\prime} 24^{\prime \prime} \mathrm{E}$ and $55^{\circ} 13^{\prime} 50$ " N to $55^{\circ} 17^{\prime} 20^{\prime \prime} \mathrm{N}$. With 6 flights $68 \mathrm{E}-\mathrm{W}$ profile lines and $16 \mathrm{~N}-\mathrm{S}$ tie lines were flown, totalling about 687 line-km. The nominal flight-line spacing was 80 m for the profile lines and 500 m for the tie lines.

The BGR helicopter-borne geophysical system includes six-frequency electromagnetics (HEM), magnetics (HMG) and radiometrics (HRD). The electromagnetic system provides information about the distribution of electrical conductivity in the earth down to a maximum depth of 150 m . The intensity of the earth's total magnetic field is measured with a magnetometer. Magnetic anomalies may have deep sources as well as shallow ones. The intensity of the gamma radiation is registered by a gammaray spectrometer. The radiation measured is mainly emitted from the elements thorium, uranium, and potassium. The origin of this radiation is normally close to the earth's surface.

The helicopter-borne system consists of the BGR helicopter, the geophysical equipment and electronic equipment for navigation. The HEM and HMG sensors, the GPS antenna and a laser altimeter are installed inside a towed tube, called bird. The navigation instruments and the gamma-ray spectrometer are mounted in the helicopter. A ground base station records the time-variant data required to correct the airborne data.

The survey altitudes of the sensors are normally $30-40 \mathrm{~m}$ for electromagnetics and magnetics and $70-80 \mathrm{~m}$ for gamma-ray spectrometry. HEM and HMG data are recorded 10 times per second during a survey flight and HRD data are recorded once per second. At an aircraft speed of about 140$150 \mathrm{~km} / \mathrm{h}$, this leads to mean sampling intervals of about 4 m and 40 m , respectively

The collected geophysical data and the corresponding positioning data are stored on a CF card during the flight. The digital data are checked immediately after the flight. Further processing of all survey data, including the data of the simultaneously operating base station which records the variations of the total magnetic intensity and the variations of the atmospheric pressure, take place in the field and finally at BGR in Hanover.

This "Technical Report" describes the survey operations and the survey equipment used, as well as the data processing and the presentation of the results as vertical resistivity sections and thematic maps. The processed data, the thematic maps and the vertical sections are stored on a DVD, accompanying this report.

Following parameters are displayed on a topographic map at a scale of 1:25,000:

- actual flight lines,
- topographic elevations,
- apparent resistivities at six frequencies ( $387,1,820,5,403,8,389,41,430$ and $133,200 \mathrm{~Hz}$ ),
- centroid depths at six frequencies ( $387,1,820,5,403,8,389,41,430$ and $133,200 \mathrm{~Hz}$ ),
- resistivities at $1,3,5,10,20,30,50,70,90$ and 120 m below ground level,
- anomalies of the total magnetic field,
- anomalies of the total magnetic field, high-pass filtered, anthropogenic anomalies removed,
- concentration of potassium,
- equivalent concentration of thorium,
- equivalent concentration of uranium,
- total count,
- exposure rate.

Cross-sections based on resistivity-depth 1-D inversion models (vertical resistivity sections) are displayed along all flight lines at a horizontal scale of 1:25,000 with a vertical exaggeration of 10 .

## 2. Introduction

Climate change simulations indicate a sea-level rise and increasing rainfall in the North Sea region. This will lead to higher groundwater levels and a forced outwash of nutrients and pollutants from industrial areas, agriculture and landfills (http://cliwat.eu/). The climate changes will affect the assessment of suitable industrial and agricultural development areas due to changes in the shape of the local waterworks catchments areas. Rise in groundwater level will challenge the construction business and it will be necessary to come up with new standards. It will also change the available groundwater resource and pattern of stream flow between summer and winter (reduced potential for irrigation from water table aquifers interacting with streams).

CLIWAT (climate \& water) is a transnational project funded by the Interreg IVB North Sea Region Programme of the European Regional Development Fund (ERDF) with partners from four participating countries of the European Union (EU): Belgium (Ghent University), The Netherlands (Deltares/TNO, VITENS, Provincie Fryslân, Wetterskip Fryslân), Germany (LIAG, LLUR, SEECON, BGR) and Denmark (Region Midtjylland, GEUS, Region Syddanmark, Environment Centre Aarhus, Environment Centre Ribe, Aarhus University, Municipality of Horsens).

The goal of the project is to determine the effects of a possible climate change on groundwater systems, surface water and the fresh-saltwater boundary in the North Sea and Baltic Sea region. The effect of the increased flux from agricultural and industrial land sites and landfills on groundwater quality in relation to indicators in the EU Water Framework Directive (WFD) has to be investigated as well as the impact on waterworks and important ground water aquifers near the coastlines. Also open question are the potential towards more accessible water in the hydrological system, the assessment of the consequences due to the increased recharge to groundwater systems and how to manage and solve the upcoming challenges for the construction business, for drainage and changes in conditions for biological/chemical decomposition in the soil.

Therefore geological and geophysical measurements were carried out in the seven pilot areas of the project (Fig. 1):

A: Belgische Middenkust, Belgium,
B: Zeeland, The Netherlands,
C: Terschelling and Northern Friesland, The Netherlands,
D: Borkum, Germany,
E: Schleswig and Southern Jutland, Germany and Denmark,
F: Egebjerg, Denmark,
G: Aarhus river, Denmark.


Fig. 1: Regions (green) funded by the Interreg IVB North Sea Region Programme of the European Regional Development Fund (ERDF) and project areas A-G. Red numbers indicate the BGR airborne survey areas.

One of these pilot areas $(\mathbf{E})$ is located on the Danish-German border covering parts of northern Schleswig and southern Jutland (Sønderjylland). The Hørløkke landfill to the west of the town of Vojens is one of the targets belonging to this pilot area (Fig. 2). The public landfill covers an area of about $12,000 \mathrm{~m}^{2}$ where approximately $65,000 \mathrm{~m}^{3}$ of household and industrial waste were filled in from 1968 to 1972. The leaching of contaminants is a severe problem as the water divide is just east of the landfill and it is suspected that a rising groundwater table will divert the plume to the east directly towards the town of Vojens.

A helicopter-borne survey of the area around Hørløkke landfill was conducted by the airborne group of the German Federal Institute for Geosciences and Natural Resources (BGR) in June 2009. The

Danish project partner Region Syddanmark, who is responsible for the coordination of the measurements and the interpretation of the diverse data sets of the area, requested the airborne survey.

This "Technical Report" describes the survey operations and the survey equipment in use, as well as the data processing and the presentation of the results as vertical resistivity sections and thematic maps. The processed data, the thematic maps and the vertical sections are stored on a DVD accompanying this report.

## 3. Survey Area

The Vojens survey area around Hørløkke landfill is situated to the northwest of the town of Vojens, Denmark. It comprises a nominal 6.5 km by 8 km wide area ranging from $9^{\circ} 11^{\prime} 49^{\prime \prime} \mathrm{E}$ to $9^{\circ} 19^{\prime} 24^{\prime \prime} \mathrm{E}$ and $55^{\circ} 13^{\prime} 50^{\prime \prime} \mathrm{N}$ to $55^{\circ} 17^{\prime} 20^{\prime \prime} \mathrm{N}$. The actual survey area differs from the planned one due to restrictions by the Danish Civil Aviation Administration. As it was not allowed to overfly densely inhabited areas or buildings without obtaining the owner's permission, the actual flight lines deviate from the nominally straight lines where homesteads, villages and towns exist. A map of the projected survey area (small red dots) and its actual realization (bold red dots) is shown in Fig. 2, which also shows the boundary (dashed black line) of the 1: 25,000 topographic map used to present the geophysical results.


Fig. 2: Vojens survey area (red dots), projected area (red dotted line) and the frame of the map sheet (black dashed line).

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An area of approximately $54 \mathrm{~km}^{2}$ was surveyed with 6 flights on June 23-25, 2009. There were 68 E-W profile lines and $16 \mathrm{~N}-\mathrm{S}$ tie lines flown, totalling about 687 line-km. The nominal flight-line spacing was 80 m for the profile lines and 500 m for the tie lines. The survey flights commenced from Vojens airport ( 42 m asl). The survey parameters are given in Table 1.

Table 1: Survey parameters for the Vojens survey area.

| BGR area number | Vojens (DK) 134 |
| :--- | :---: |
| Field period | June 23-25, 2009 |
| Size of survey area | $54 \mathrm{~km}^{2}$ |
| Total length of survey lines | 687 km |
| Number of survey flights | 6 |
| Flight numbers | $13401-13406$ |
| Mean flight altitude of the EM sensor above ground | 40 m |
| Speed during survey flight | $140 \mathrm{~km} / \mathrm{h}$ |
| Number of profile-line flights | 4 |
| Number of profile lines | 68 |
| Profile-line lengths | $4-12 \mathrm{~km}$ |
| Profile-line directions (angle to N) | $90^{\circ}$ |
| Profile-line spacing | 80 m |
| Number of tie-line flights | 2 |
| Number of tie lines | 16 |
| Tie-line lengths | $3-8 \mathrm{~km}$ |
| Tie-line directions (angle to N) | $0^{\circ}$ |
| Tie-line spacing | 500 m |
|  |  |

The lines flown primarily northwards or eastwards are normally given an even profile number, while the ones flown in the opposite directions are odd numbered. The profile lines have the extension ". 1 " (after the profile number) or ". 2 " or " .3 " for repeated lines, and the tie lines have the extension " .9 " or ". 8 " for repeated lines. Details of the survey flights are given in Appendix I.

The average altitude of the helicopter was 40 m above ground level within the survey area. During a survey flight, particularly before the first and after the last profile, the altitude was increased to $>350 \mathrm{~m}$ to check the calibration of the HEM system far from any disturbing influences.

The base station recording the magnetic variations was located on the airport at $55^{\circ} 13^{\prime} 39^{\prime \prime} \mathrm{N}$, $9^{\circ} 16^{\prime} 46^{\prime \prime} \mathrm{E}$ and 43 m asl.

## 4. Airborne Geophysical System

BGR's airborne geophysical system simultaneously records the electromagnetic, magnetic, and gamma-ray spectrometry data. The geophysical instrumentation, the navigation and positioning systems, the digital recording units, as well as other equipment needed for the survey flights are integrated in one measuring system carried by a Sikorsky S-76B helicopter (Fig. 3).


Fig. 3: Principal sketch of the BGR airborne geophysical system.

### 4.1. The Helicopter

The helicopter, a Sikorsky S-76B (see Table 2), was purchased in 1986 by the Federal Ministry for Economic Cooperation and Development and assigned to BGR, mainly for technical cooperation projects.

Table 2: Technical specifications of the BGR helicopter $D-H B G R$

| Helicopter |  |
| :--- | :--- |
| Type | Sikorsky S-76B (Manufacturer: Sikorsky, USA) |
| Year of manufacture | 1986 |
| Engines | 2 turbines Pratt \& Whitney PT6B-36A <br> with 1033 SHP (shaft horse power) for each |
| Maximum gross weight | 11,700 pounds (5,363 kg) |
| Maximum payload | 3,300 pounds $(1,500 \mathrm{~kg})$ |
| Maximum flight duration | $2: 45$ hours |
| Fuel consumption per hour | $350-4001$ |

### 4.2. Measuring System

The airborne geophysical system (Table 3) is installed in the helicopter and in a towed tube, called bird. The navigation instruments and the gamma-ray spectrometer are mounted in the helicopter. The HEM and HMG sensors, the GPS antenna and a laser altimeter are installed inside the bird. This bird is towed by a 45 m long cable and its position is, depending on the flight speed, about 40 m beneath and little behind the helicopter. A ground base station records the time-variant data required to correct the airborne data.

The geophysical and recording systems are controlled by the HeliDas system that also assists the navigation during a survey flight. The operator as well as the navigator is able to check the flight data online as information about the flight path and selected data channels are displayed on tablet computers.

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Table 3: The geophysical survey system

| Geophysical systems |  |  |
| :--- | :--- | :--- |
| I. Six -frequency electromagnetic system (HEM) |  |  |
| Function | Investigation of the underground electric conductivity down to a <br> maximum depth of about 150 m |  |
| Manufacturer | Fugro Airborne Surveys (FAS), Canada |  |
|  | Type | RESOLVE, BKS36a (Bird 61) |
| II. Caesium magnetometer |  |  |
| Function | Recording of the total magnetic intensity of the earth |  |
| Manufacturer | Geometrics, USA |  |
|  | Type | G-822A |
|  | III. Gamma-ray spectrometer |  |
| Function | Recording of the energy spectrum of natural and man-made gamma |  |
| radiation within a range of 0 to 3 MeV |  |  |

### 4.3. Electromagnetics

A sinusoidal current flow through a transmitter coil at a discrete frequency generates the primary magnetic field. At a distance greater than about 2 m this field is very similar to a field of a magnetic dipole located in the centre of the transmitter coil. The resulting eddy currents in the subsurface generate a secondary magnetic field that depends on the frequency used and the conductivity distribution. The difference of the fields picked up by the receiver coil and a bucking coil, which is used to cancel out the dominating primary field, is related to the primary magnetic field at the receiver coil, i. e., the quantity measured is the relative secondary magnetic field in parts per million (ppm). Due to a small phase shift between the primary and the secondary field, the relative secondary magnetic field is a complex quantity with in-phase and out-of-phase (quadrature) components.

The HEM system, RESOLVE manufactured by Fugro Airborne Surveys, utilises six individual coil systems consisting of transmitter, receiver, bucking and calibration coils. The transmitter and receiver coils have a diameter of about half a metre and a distance of about 8 m . The orientation of five transmitter-receiver coil systems is horizontal coplanar (HCP) what is suitable for groundwater exploration purposes as the induced currents are predominantly flowing horizontally resolving layered structures best. In addition, a vertical coaxial coil (VCX) system is used in order to better locate vertical structures such as fault or fracture zones. The coil systems are housed by a 10 m long tube.

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Table 4: HEM system parameters (Bird 61)

| Frequency <br> $[\mathrm{Hz}]$ | Coil separation <br> $[\mathrm{m}]$ | Coil orientation | Denotation <br> FAS | Denotation <br> BGR |
| :---: | :---: | :---: | :---: | :---: |
| 387 | 7.938 | horizontal coplanar | EM_3 | 1. frequency |
| 1,820 | 7.931 | horizontal coplanar | EM_5 | 2. frequency |
| 5,403 | 9.055 | vertical coaxial | EM_6 | 3. frequency |
| 8,389 | 7.925 | horizontal coplanar | EM_2 | 4. frequency |
| 41,430 | 7.912 | horizontal coplanar | EM_1 | 5. frequency |
| 133,200 | 7.918 | horizontal coplanar | EM_4 | 6. frequency |

Small coils placed in the centre of each receiver coil are used for calibration. The calibration factors necessary to convert the measured signals to ppm values were provided by the manufacturer. The inphase and quadrature components of the relative secondary magnetic fields are used to derive the three-dimensional distribution of the electrical conductivity - or its inverse, the resistivity - in the subsurface. Horizontal resolution and vertical resolution are achieved by moving the system and using different system frequencies, respectively. Due to the skin-effect (high frequency currents are flowing on top of a perfect conductor) the penetration depths of the electromagnetic fields increase with decreasing frequency and conductivity. The frequencies used range from 387 Hz to 133 kHz enabling exploration depth ranges of about 1-30 m in a very conductive host such as saltwater saturated sediments and $5-150 \mathrm{~m}$ in a rather resistive host such as freshwater saturated sandy sediments.

The HEM system is not only sensitive to the electrically conductive subsurface but also to anthropogenic objects like, e. g., buildings, metallic bodies, and electrical installations, which have influence on the data measured, particularly at lower frequencies. As the helicopter itself is such an object, the HEM system is towed at a sufficiently large distance (about 40 m ) underneath the helicopter.

### 4.4. Magnetics

A highly sensitive caesium vapour magnetometer installed in the bird is used to measure the total intensity of the earth's magnetic field (unit Nanotesla, nT). The function of a caesium magnetometer is based on the measurement of the so-called Larmor frequency that occurs in a special, optically pumped system in the sensor. The frequency is directly proportional to the magnetic field intensity and can be determined with high precision and accuracy. The resolution of the instrument is 0.01 nT .

The magnetic field measured is composed of different parts. The Earth's main field, caused by sources in the Earth's core, varies between approximately $20,000 \mathrm{nT}$ in equatorial regions and $70,000 \mathrm{nT}$ at the poles. It is superimposed by the crustal magnetic field caused by rocks containing magnetised minerals. These produce anomalies in the range between less than one and up to several

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hundred nT. In populated areas, anthropogenic sources such as buildings, industrial plants, power lines, etc. can produce additional locally confined and sometimes strong magnetic anomalies. Finally, the magnetic field is subject to temporal changes due to fluctuations in the state of the ionosphere and magnetosphere. These diurnal variations are in the order of several tens of nT.

In order to record the diurnal variations, a magnetic base station (Table 5) is operated. The station, also equipped with a caesium magnetometer, is installed close to the area of investigation in a magnetically undisturbed place. Data recorded by the base station during the survey are used to correct the total magnetic field measured during the flight. GPS time is used to synchronise the two data sets.

Table 5: Base station

| Base station |  |
| :--- | :--- |
| Magnetic base station | Recording of the variation of the total magnetic intensity (TMI) |
| Function | Base station: FAS, Canada <br> Magnetometer: Cs sensor H-8, SCINTREX, Canada |
| Manufacturer | CF1 Data Logger |
| Type |  |

### 4.5. Radiometrics

For geophysical investigations the count rates of the common terrestrial radioactive elements (or their isotopes and daughter products) Tl-208 (thorium series), Bi-214 (uranium series), K-40 (potassium) are of interest. Mapping of the distribution of these three elements in the ground are useful for geological investigations.

BGR uses a standard 256-channel spectrometer system consisting of four sodium iodide (NaI) crystals to detect the ground gamma radiation and one upward looking crystal to detect the radon radiation in the air. The spectrometer crystals are placed together in an aluminium box. Each crystal has a volume of approximately $41\left(0.1 \times 0.1 \times 0.4 \mathrm{~m}^{3}\right)$. Incident gamma radiation is absorbed by the crystals and transformed to light pulses that are converted to electric pulses using a photomultiplier tube. The amplitudes of the electric pulses are directly proportional to the energy of incident gamma radiation.

The spectrometer covers an energy spectrum from 0 to 3 MeV . Depending on their energy, the pulses are mapped into one of 255 energy channels. Channel 256 is reserved for recording cosmic radiation between 3 and 6 MeV . Spectra recorded by the system contain counts of gamma radiation collected and integrated over one second. Energy windows and channel ranges of the different radiation sources are listed in Table 6. The spectrometer is internally stabilised for possible drifts in gain. This is done independently for each of the four downward-looking crystals using the thorium peak. Shifts of the thorium peak $(2.62 \mathrm{MeV})$ relative to the nominal value are identified and the gain of the photomultiplier tube of the respective crystal is corrected automatically. A caesium sample is used to stabilize the gain of the upward looking crystal.

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Table 6: Radiation sources and corresponding spectrometer parameters

| Radiation source | Energy window <br> in MeV | Peak energy <br> in MeV | Channel range |
| :---: | :---: | :---: | :---: |
| Total count | $0.41-2.81$ | - | $34-233$ |
| Potassium (K-40) | $1.37-1.57$ | 1.46 | $115-131$ |
| Uranium (Bi-214) | $1.66-1.86$ | 1.76 | $139-155$ |
| Thorium (Tl-208) | $2.41-2.81$ | 2.62 | $202-233$ |
| Cosmic radiation | $3.0-6.0$ | - | 255 |

### 4.6. Navigation and Positioning

The navigation system (Table 7) provides the pilot with all the information necessary to carry out a survey flight. Navigation software (LiNav von AG-NAV Inc.) calculates the coordinates of the starting and the end points of all survey lines from the coordinates of the corners of the survey area, the profile direction and the spacing of the flight lines. These coordinates are copied to the HeliDas system using a CF card or an USB stick. These profiles are displayed on the tablet computer with the line being flown highlighted.

Table 7: Navigation and positioning systems

| Systems for navigation and positioning |  |  |
| :---: | :---: | :---: |
|  | Navigation system |  |
|  | Function | On-line determination and display of the GPS navigational data required by the pilot during a survey flight; recording of the geographic position of the helicopter and its altitude above mean sea level |
|  | Manufacturer | Navigation computer and display: FAS, Canada GPS receiver: NovAtel, Canada |
|  | Type | Navigation computer: HeliDas GPS receiver: NovAtel OEMV-2-L1/L2 GPS antenna: NovAtel L1/L2 ANT-532-e |
| 홌 | Positioning system |  |
|  | Function | Determination and recording of the geographic position of the HEM bird and its altitude above mean sea level |
|  | Manufacturer | Position recording and display: FAS, Canada GPS receiver: CSI Wireless, Canada |
|  | Type | Position recording: HeliDas GPS receiver: DGPS MAX |

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The pilot obtains all information required to fly this profile as accurately as possible from a second display. The most important information is the lateral deviation from a line. The deviation appears digitally in metres, as well as on a bar diagram. The navigation computer receives information about the position of the helicopter from a GPS navigation receiver whose antenna is fixed outside on the helicopter. The error in the navigation data is less than 1-2 m.

The positioning system (Table 7) provides the coordinates of each geophysical measurement. A second GPS navigation receiver is used for this purpose, whose antenna is fixed inside the bird. The spatial positions of the sensors are determined from this positioning data. The error of the coordinates is also in the order of $1-2 \mathrm{~m}$.

A radar altimeter (Table 8) attached to the bottom of the helicopter determines its altitude above the ground or above obstacles (e. g., large stands of trees and buildings) with a precision of $\pm 3 \mathrm{~m}$. The altitude is needed to process the radiometric data. A barometric altimeter is used to determine the altitude of the helicopter above mean sea level, but this altimeter is employed only as a backup for the GPS receivers. Without a base station as reference the GPS measurements may have an error of some metres.

The altitude of the bird above the ground must be accurately known for the processing of the electromagnetic data and to generate a digital terrain model. A laser altimeter (Table 8) inside the bird provides this altitude with a precision of $\pm 0.2 \mathrm{~m}$. A further advantage of the laser altimeter, in addition to its precision, is the focused laser beam, which when above a forest often allows the distance to the surface to be determined and not only to the treetops, as it is the case with the radar altimeter.

Table 8: Altimeters

| Altimeters |  |  |
| :---: | :---: | :---: |
| $\begin{aligned} & \text { 흔 } \\ & \frac{0}{0} \\ & \frac{1}{0} \end{aligned}$ | Radar Altimeter |  |
|  | Function | Recording of the altitude of the helicopter above ground level |
|  | Manufacturer | Sperry, USA |
|  | Type | AA-200 |
|  | Barometric Altimeter |  |
|  | Function | Recording of the altitude of the helicopter above mean sea level |
|  | Manufacturer | Rosemount, USA |
|  | Type | 1241A5B |
| 읐 | Laser Altimeter |  |
|  | Function | Precise recording of the altitude of the HEM bird above ground |
|  | Manufacturer | Riegl, Austria |
|  | Type | LD90-3800VHS |

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The digital elevation model is derived from the GPS elevation of the HEM bird in $m$ asl minus the laser altitude. Without a base station as reference for the GPS measurements, and thus, the topographic elevations may have an error of some metres.

### 4.7. Data Acquisition and Recording

The HeliDas system stores all the data digitally on CF card during a survey flight (Table 9). The data sets are ready for processing with GEOSOFT OASIS montaj. The most important data channels are also displayed on the tablet computers to enable continual checking of the data during the flight. Immediately after a flight, the digital data are copied to a field computer and checked more accurately in order to obtain an impression of the geophysical results and to detect any problems with the survey system.

Table 9: Data acquisition and recording systems

| $\begin{aligned} & \frac{2}{\dddot{\prime}} \\ & \frac{0}{0} \\ & \frac{3}{0} \\ & \hline \mathbf{y} \end{aligned}$ | Data acquisition and recording systems |  |
| :---: | :---: | :---: |
|  | Function | Digitizing of the analogue signals, buffering of all digital data; flight path and displaying of selected data channels; storage of position and field data on CF card ready for processing with GEOSOFT OASIS montaj |
|  | Manufacturer | FAS, Canada |
|  | Type | HeliDas |

### 4.8. Video System

A video camera (Table 10) is mounted in the bottom of the helicopter. Two monitors, one in the cockpit and one in the operator's rack, allow monitoring of the bird at take-off and landing as well as during the flight.

The video recording of the flight path is used to locate sources of anomalous or disturbed data on the ground. The flight path video can be correlated directly with the digital data.

Table 10: Video system

| Video system |  |  |
| :---: | :---: | :---: |
|  | Function | Recording of the flight track and monitoring of the movements of the HEM bird during take-off, landing and flight |
|  | Manufacturer | Colour camera: Sony, Japan Video recorder: AXI, Sweden |
|  | Type | Colour camera: DC372P <br> Video server: AXIS 241S |

### 4.9. Additional Equipment

The 28 V DC on-board voltage of the helicopter is smoothly buffered by a 24 Ah battery and connected to a central power unit. From there it is distributed to the individual components of the system with fuses built-in to protect devices from overvoltage.

Control and recording units of the airborne geophysical system are mounted in a 19" rack. Shock absorbers between the base of the rack and a wood board which is firmly screwed to the floor of the helicopter minimize the transfer of vibrations originating from the rotor.

Table 11: Additional equipment

| Central power unit |  |
| :--- | :--- |
| Function |  |
|  | 28 V DC on-board voltage of the helicopter buffered by a 24 Ah <br> buffer battery and connected to a central power unit |

## 5. Processing and Presentation of the Survey Data

The general objectives of the data processing may be summarized as follows:

- quality control of the measured data;
- conversion of the field data into physical parameters;
- presentation of the results as maps and vertical sections.


### 5.1. General Processing Steps

The airborne geophysical data are copied from the CF card to field computers directly after a survey flight in order to save the data and to check them for plausibility and for correctness. Using the software GEOSOFT OASIS montaj, the primary field data processing steps are conducted automatically, followed by a pre-processing of all survey data in order to display preliminary results.

The final data processing starts with the processing of the position data:

- coordinate transformation;
- correction of altitude data of the helicopter and the bird.

The following processing steps are valid for all methods:

- removal of spiky data;
- reduction of high-frequency noise by digital filtering;
- conversion of the data to the desired geophysical parameters;
- fixing of the ends of the profiles;
- merging the flight-line data sets to area data sets;
- levelling of the data;
- storage of the final survey data and geophysical parameters;
- production of maps and vertical sections (only HEM).

The field data processing and the calculation of the physical parameters for each method are described in more detail in the following chapters. GEOSOFT OASIS montaj is used throughout if not otherwise noted.

### 5.2. Position Data

### 5.2.1. Coordinates

The coordinates of the helicopter and the bird recorded during the survey flight refer to the WGS 84 geographic coordinate system. These geographic coordinates are transformed to local Cartesian coordinates. False coordinates are corrected und gaps are interpolated.

All survey results refer to UTM WGS 84 coordinates ( $9^{\circ}$ meridian, zone 32 N ).

### 5.2.2. Radar Altitude

The radar altitude data measured in feet at the helicopter ( $h{ }_{-} r_{\text {mess }}$ ) have to be transformed to metres above ground level ( m agl). For the purpose of comparison with the laser altitude data of the bird (h_l), the radar altitudes are also referred to the bird altitude (h_r)

$$
h \_r[m]=h \_r_{\text {mess }}[\text { feet }] * 0,3048[\mathrm{~m} / \text { feet }] * r_{1}-c_{1}[\mathrm{~m}],
$$

where
$\mathrm{h} \_\mathrm{r}[\mathrm{m}]=$ adjusted radar altitude (unit: m agl),
$h \_r_{\text {mess }}[f e e t]=$ radar altitude (unit: feet) measured by the altimeter,
$r_{1} \quad=$ conversion factor (gradient),
$c_{1} \quad=$ effective cable length (offset).
For this, the effective cable length, i. e., the distance between the helicopter and the bird, has to be estimated and subtracted. The effective cable length can be derived from the differences of the GPS elevations of the helicopter and the bird. Alternatively, the effective cable length and the conversion factor $r_{1}$ of laser and radar altitudes are obtained by linear regression.

For the correction of the radar altitude data $r_{1}=1.04$ and $c_{1}=44 \mathrm{~m}$ were used.

### 5.2.3. Laser Altitude

The laser altimeter data representing the bird altitude - as well as the radar altimeter data - may have gaps and outliers which have to be corrected by elimination and interpolation procedures. The movement of the bird causes attitudes (pitch and roll) deviating from the normal case and, thus, laser altitudes which are normally higher than the actual bird altitude. The mean pitch angle of about $6^{\circ}$ is corrected by applying the corresponding cosine function. The roll angle is generally not known. Thus, after identification by comparison with the radar altitudes of the bird, strongly affected laser altitudes have to be eliminated and interpolated afterwards.

The measurements of the laser altimeter data may be affected by the tree canopy or other reflectors. Thus, the distance between the helicopter containing the radiometric system or the bird containing the electromagnetic and magnetic systems and the ground level is often not correctly measured resulting in laser altitudes which are too low.

The affected laser altitudes (h_l) are corrected with the help of a combination of several checks and filter techniques (Table 12). The first step is to reduce the effect of strong gradients in the laser altitudes due to rapid changes in bird or topographic elevation. A base line derived by applying a low-
pass filter to the laser altitude data is subtracted from the laser altitude data to calculate reduced laser altitude values ( $\Delta \mathrm{h} \_\mathrm{l}$ ). Remaining outliers are removed by applying a very short non-linear filter. In order to identify and eliminate those segments where trees or other obstacles exist, two procedures are applied to the reduced laser altitude data:
a) Noise filter, followed by non-linear and low-pass filters applied to the noise channel ( $\Delta \mathrm{h} \mathrm{l}_{\mathrm{noise}}$ ), and a high-noise threshold of 0.4 m ;
b) Maximum filter and difference threshold of 2 m of filtered ( $\Delta \mathrm{h}_{-} \mathrm{l}_{\max }$ ) and unfiltered ( $\Delta \mathrm{h} \_\mathrm{l}$ ) data.

The gaps of eliminated data are filled in with slightly shifted maximum values representing the corrected reduced laser altitudes. As the maximum values may be too high, their levels are shifted to the levels on both side of each gap. Finally, the corrected values are low-pass filtered ( $\Delta \mathrm{h} \mathrm{l}_{\mathrm{kor}}$ ) and the base line is added again to get the corrected laser altitude values ( $h l_{\mathrm{k}} \mathrm{k}_{\mathrm{kor}}$ ). This procedure is able to eliminate all effects caused by single or small groups of trees. The effect of broad and densely wooded areas, however, is not always removed sufficiently and has to be corrected manually.

Table 12: Filter parameters for the removal of the tree-canopy effect

| Type of filter | Filter parameters | Channel |
| :---: | :---: | :---: |
| Low pass | Cut-off period: $5 \mathrm{~s}(\approx 200 \mathrm{~m}$ ) | h_l |
| Non linear | Window length: 1 point ( $\sim 5 \mathrm{~m}$ ), tolerance: 1.0 | $\Delta \mathrm{h}$ _1 |
| Noise (normal distr.) | Window length: 7 points ( $\sim 28 \mathrm{~m}$ ) | $\Delta \mathrm{h}$ _1 |
| Non linear | Window length: 3 points ( $\approx 15 \mathrm{~m}$ ), tolerance: 1.0 | $\Delta h_{-} 1_{\text {noise }}$ |
| Low pass | Cut-off period: $1 \mathrm{~s}(\approx 40 \mathrm{~m})$ | $\Delta \mathrm{h}_{-} \mathrm{l}_{\text {noise }}$ |
| Threshold | Cut-off value ( $\Delta \mathrm{h}_{\text {_ }} \mathrm{noises}$ ): 0.4 m | $\Delta \mathrm{h}$ _1 |
| Maximum | Window length: 21 points ( $\sim 84 \mathrm{~m}$ ) | $\Delta \mathrm{h}$ _l |
| Threshold | Cut-off value ( $\Delta \mathrm{h} \mathrm{l}_{\text {max }}-\Delta \mathrm{h} \_$l): 2 m | $\Delta \mathrm{h}$ _1 |
| Low pass | Cut-off period: $3 \mathrm{~s}(\approx 120 \mathrm{~m}$ ) | $\Delta \mathrm{h} 1_{\text {kor }}$ |

### 5.2.4. Topographic Elevation

The topographic relief derived by the difference of the GPS based bird elevation and the corrected laser altitude

$$
\text { topo }[\mathrm{m} \text { asl }]=\mathrm{h} \_ \text {GPS }[\mathrm{m} \text { asl }]-\mathrm{h} \_\mathrm{l}_{\mathrm{kor}}[\mathrm{~m}]
$$

is used to derive a digital elevation model of the survey area. As the tree canopy effect causes laser altitudes which are too low, the topographic elevations are too high. Therefore, the topographic values are also useful to identify and manually correct the laser altitude for remaining tree canopy effects, particularly if external topographic data as reference are available.

In order to remove line effects the topographic elevation data were levelled with respect to the digital elevation model provided by the project partner.

### 5.3. Processing of the Electromagnetic Data

The processing of the measured $I$ and $Q$ values (in ppm), i. e., the real part (in-phase or $0^{\circ}$-phase) and the imaginary part (out-of-phase, quadrature or $90^{\circ}$-phase) of the relative secondary field requires several steps:

- application of calibration factors;
- zero-level and drift correction;
- data correction;
- transformation to half-space parameters;
- correction of man-made effects;
- levelling;
- interpolation and smoothing.

While the half-space parameters, apparent resistivity and the centroid depth, are individually derived from secondary field values for each frequency, the final resistivity models are calculated at each survey point by 1-D inversion of the data of all (or selected) frequencies.

### 5.3.1. Calibration of the HEM System

The HEM system was calibrated on highly resistive ground by the manufacturer in Mountsburg Conservation Area, Canada. After adjusting the phase with the help of a ferrite rod, well-defined external calibration coils were used to derive the ppm values of the internal calibration coils. These calibration factors are used to convert the voltages measured during a survey flight to ppm values representing the secondary magnetic fields (Table 13).

Table 13: Calibration factors of the HEM system

| Frequency <br> [Hz] | Calforation factors Fugro <br> I [ppm] <br> Q [ppm] |  | Calibration factors BGR <br> Ippm] |  |
| :---: | :---: | :---: | :---: | :---: |
| 387 | -205.3 | -205.3 | -209.8 | -210.8 |
| 1,820 | -175.4 | -174.7 | -174.7 | -174.3 |
| 5,403 | 76.6 | 76.8 | 81.9 | 81.2 |
| 8,389 | -144.4 | -144.2 | -209.4 | -198.8 |
| 41,430 | -667.3 | -665.2 | -657.4 | -664.9 |
| 133,200 | -1404.2 | -1406.4 | -685.5 | -911.0 |

At the beginning of each survey flight and at high flight altitude, phase and gain of the EM system are adjusted automatically for each frequency using internal calibration coils. Due to instrumental drift, the calibration has to be checked several times during the flight. The calibration signals caused by internal calibration coils are compared with known calibration signals and phase shifts and gain correction factors are applied to the data.

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As a mutual coupling with the subsurface during the ground calibration procedure and technical changes of the system caused modified calibration factors, a flight over highly conductive North Sea water in February 2009 was used to check the calibration values. The evaluation of this data set yielded a set of phase and gain corrections being enormous particularly for the 8.3 and 133 kHz frequency data (Table 13).

A further check of the calibration factors during the Zeeland survey in August 2009 yielded an updated set of mean phase and gain corrections ( $8.3 \mathrm{kHz}: 2 \%$ gain, $41 \mathrm{kHz}: 0.5^{\circ}$ phase and $133 \mathrm{kHz}: 7^{\circ}$ phase and $16 \%$ gain). In addition, a fine tuning of the phase and gain correction values was necessary for the highest frequency data of the Vojens survey.

### 5.3.2. Zero-Level and Drift Correction

The signals measured by the receivers may still contain some non-compensated parts of the primary fields generated by the transmitters. These so called zero levels may also have thermal drift. The zero levels of the HEM data are generally determined at high flight altitudes ( $>350 \mathrm{~m}$ ) several times during a survey flight as the ground response is negligible at this altitude, i. e., the secondary field should be close to zero. Zero-level reference points are set at such high-altitude profile segments, preferably where the signal is not noisy. The zero level is obtained individually for each data channel by linear interpolation of the picked values at adjacent zero level reference points.

This procedure enables to remove the long-term, quasi-linear drift. Short-term variations, however, caused by temperature changes due to altitude variations, which occur particularly in the highestfrequency data, cannot be corrected successfully by this procedure. Therefore, additional reference points - also along the profiles at normal survey flight altitude - have to be determined where the secondary fields are small but not negligible. At these locations, the estimated half-space parameters are used to calculate the expected secondary field values, which then serve as local reference levels (Siemon, 2009). As this drift correction procedure is often not sufficient, statistical levelling procedures have to be applied in addition (see Section 5.3.6).

### 5.3.3. Data Correction

Noise from external sources (e. g., from radio transmitters, power lines, sferics, built-up areas, streets, railway tracks) is eliminated from the HEM data by appropriate filtering or interpolation. All those field values (I or Q) are automatically eliminated which fall below the relative standard error (rel. STE = STE/Mean) of the field values within a given data window. The field values are smoothed using a combination of non-linear (Naudy \& Dreyer, 1968) and low-pass filters to exclude outliers and to suppress high-frequency noise, respectively. Due to frequency dependent data qualities the data channels are treated individually (Table 14).

Induction effects from buildings and other electrical installations (see Section 5.3.5) or effects from strongly magnetized underground sources are normally not erased from the data during the initial stage of data processing.

Table 14: Filter parameters for HEM data processing

| Frequency <br> [Hz] | Mean / STE <br> [Values] | Threshold (I/Q) <br> of rel. STE | NL filter <br> Values/Tolerance | LP filter <br> TLP [Values] |
| :---: | :---: | :---: | :---: | :---: |
| 387 | $75 / 25$ | $0.06 / 0.04$ | $20 / 3.0$ | 40 |
| 1,820 | $75 / 25$ | $0.05 / 0.05$ | $20 / 2.0$ | 30 |
| 5,403 | $75 / 20$ | $0.05 / 0.05$ | $10 / 2.0$ | 30 |
| 8,389 | $75 / 15$ | $0.05 / 0.05$ | $10 / 2.0$ | 30 |
| 41,430 | $75 / 15$ | $0.05 / 0.05$ | $5 / 2.0$ | 30 |
| 133,200 | $75 / 15$ | $0.04 / 0.02$ | $5 / 2.0$ | 30 |

### 5.3.4. Conversion of the Secondary Field Values to Half-Space Parameters

The relative secondary magnetic field $Z=\left(I_{c}, Q_{c}\right)$ for a horizontal-coplanar (HCP) coil pair with a coil separation $r$ and at an altitude $h$ above the surface are calculated by (e.g. Ward \& Hohmann, 1988)

$$
\mathrm{Z}=\mathrm{r}^{3} \int_{0}^{\infty} \mathrm{R}_{1}(\mathrm{f}, \lambda, \rho, \mu, \varepsilon) \frac{\lambda^{3} \mathrm{e}^{-2 \alpha_{0} \mathrm{~h}}}{\alpha_{0}} \mathrm{~J}_{0}(\lambda \mathrm{r}) \mathrm{d} \lambda
$$

where $\alpha_{0}{ }^{2}=\lambda^{2}-\omega^{2} \mu_{0} \varepsilon_{0}+i \omega \mu_{0} / \rho_{0}$ with $\mu_{0}=4 \pi * 10^{-7} \mathrm{Vs} / \mathrm{Am}, \varepsilon_{0}=8.854 * 10^{-12} \mathrm{As} / \mathrm{Vm}$ and $\rho_{0}>10^{8} \Omega \mathrm{~m}$, $J_{0}$ is Bessel functions of first kind and zero order, and $R_{1}$ is the complex reflection factor containing the material parameters of the subsurface. This complex integral is evaluated numerically using fast Hankel transforms (e. g. Anderson, 1989, Johansen \& Sørensen, 1979). A similar formula exists for a coaxial coil (VCX) configuration yielding smaller ppm values (VCX $\approx-0.25 * \mathrm{HCP}$ ). Following Weidelt (1991) the reflection factor $\mathrm{R}_{1}$ for a N -layer half-space model is derived by a recurrence formula

$$
\mathrm{R}_{1}=\frac{\mathrm{B}_{1}-\alpha_{0} \mu / \mu_{0}}{\mathrm{~B}_{1}+\alpha_{0} \mu / \mu_{0}}
$$

with

$$
\begin{gathered}
B_{n}=\alpha_{n} \frac{B_{n+1}+\alpha_{n} \tanh \left(\alpha_{n} t_{n}\right)}{\alpha_{n}+B_{n+1} \tanh \left(\alpha_{n} t_{n}\right)} \quad n=1,2, \ldots, N-1 \text { and } B_{N}=\alpha_{N} \\
\alpha_{n}=\sqrt{\lambda^{2}-\omega^{2} \varepsilon_{n} \mu_{n}+i \omega \mu_{n} / \rho_{n}} \quad n=1,2, \ldots, N
\end{gathered}
$$

where $\rho_{\mathrm{n}}, \varepsilon_{\mathrm{n}}, \mu_{\mathrm{n}}$ and $\mathrm{t}_{\mathrm{n}}$ are resistivity, dielectric permittivity, magnetic permeability and thickness of the $\mathrm{n}^{\text {th }}$ layer, respectively ( $\mathrm{t}_{\mathrm{N}}$ is assumed to be infinite). As magnetic effects and displacement currents are negligible, i. e., $\mu_{\mathrm{n}}=\mu_{0}$, and $\varepsilon_{\mathrm{n}}=\varepsilon_{0}$ only resistivities and depths are taken into account (Fig. 4).

Calculated secondary field values $I_{c}$ and $Q_{c}$ (in ppm ) are used to convert the calibrated measured values (I and Q) to the parameters of a homogeneous half-space (Siemon, 2001),

- apparent resistivity $\rho_{\mathrm{a}}[\Omega \mathrm{m}]$ and
- apparent distance $\mathrm{D}_{\mathrm{a}}[\mathrm{m}]$ from the sensor to the top of the conducting half-space, individually for each frequency.

For this, the reduced amplitude $\mathrm{A}_{c}^{\prime}=(\mathrm{h} / \mathrm{r})^{3} * \mathrm{~A}_{c}$ with $\mathrm{A}_{c}=\left(\mathrm{I}_{c}{ }^{2}+\mathrm{Q}_{c}{ }^{2}\right)^{1 / 2}$ and the ratio $\varepsilon_{c}=\mathrm{Q}_{c} / \mathrm{I}_{c}$ are calculated for an arbitrary half-space as a function of the ratio $\delta=\mathrm{h} / \mathrm{p}$ of sensor altitude h and skin depth $p=503.3 *\left(\rho_{a} / f\right)^{1 / 2}$.

The half-space parameters are then derived for each pair of measured secondary field values from the functions $\mathrm{A}_{c}^{\prime}(\delta)$ and $\delta\left(\varepsilon_{c}\right)$ approximated by polynomials $\left(\mathrm{A}_{\mathrm{p}}^{\prime}(\delta)\right.$ and $\left.\delta_{\mathrm{p}}(\varepsilon)\right)$ :

$$
D_{a}=r\left(A_{p}^{\prime}\left(\delta_{p}(\varepsilon) / A\right)\right)^{1 / 3} \quad \text { and } \quad \rho_{\mathrm{a}}=0.4 \pi^{2} f\left(D_{a} / \delta_{p}(\varepsilon)\right)^{2} .
$$

The calculated distance $\mathrm{D}_{\mathrm{a}}$ may differ from the observed HEM sensor altitude (in m above ground level), i. e., the top of the conducting half-space model needs not to coincide with the surface of the earth as determined by the altimeters. The difference between the two quantities is defined as the apparent depth $d_{a}=D_{a}-h$. If $d_{a}$ is positive, a resistive cover is assumed above the half-space. If $d_{a}$ is negative, a conductive cover is assumed.

In addition to the apparent resistivity $\rho_{a}$ and apparent distance $D_{a}$, the centroid depth $z^{*}=d_{a}+p / 2$ is determined (Siemon, 2001). The centroid depth is a measure of the mean penetration of the induced underground currents. The resulting sounding curves, $\rho_{\mathrm{a}}\left(\mathrm{z}^{*}\right)$, provide the initial approximation of the vertical resistivity distribution.

## Model 1: Homogeneous half-space



Model 2: Layered half-space


Fig. 4: HEM inversion based on a homogeneous half-space or a layered half-space
The actual approach for calculating the half-space parameters differs from that described by Siemon (2001) as the field values are calculated more accurately, particularly at higher frequencies, and the polynomial approximation of the functions $\mathrm{A}^{\prime}(\delta)$ und $\delta(\varepsilon)$ are optimised for each individual frequency.

The half-space parameters are checked for plausibility, i. e., high altitude (h>100 m) and extreme ( $\rho_{a}>1000 \Omega m, d_{a}>100 m$ ) values have been eliminated, before they are used for further processing.

### 5.3.5. Effect of Anthropogenic Influences on the HEM Data

In addition to the geogenic contribution to the secondary fields measured over densely populated areas, there is often an anthropogenic contribution from buildings and electrical installations etc. Generally, these have little influence on the HEM data and the data can be corrected using the standard data processing tools. In some cases, e. g., large buildings with a high metal content, the anthropogenic components in the HEM data are no longer negligible. Furthermore, external electromagnetic fields exist close to power lines, electric railway tracks or built-up areas which are able to substantially affect the HEM measurements. These man-made effects appear particularly in the lower frequency data because the geogenic contribution to the secondary fields is comparatively smaller at lower than at higher frequencies and, thus, the anthropogenic contribution, which is rather frequency independent, may dominate.

The anthropogenic influence lowers the calculated resistivity and associated depth. Thus, low resistivity and depth pattern on maps and sections often correlate with man-made effects such as villages or streets. These man-made effects can be detected in the HEM data due to their typical shape or by correlation with magnetic data. Topographic or Google Earth maps of the survey area, an analysis of the video records or an on-site inspection can help identify such effects.

A manual correction of man-made effects is very time consuming as each HEM channel of each survey line has to be examined individually. Therefore, a semi-automatic filter procedure has been developed and integrated into GEOSOFT OASIS montaj software. This procedure uses the gridded data of the half-space parameters apparent resistivity and apparent depth. These grids are inspected (once or several times) for each individual frequency for anomalous data. Minimum and/or maximum anomalies are detected when the differences of the grid values and their corresponding median values, which are calculated in circular areas shifted over the grid, exceed a given threshold (Ta-
ble 15).
Table 15: Filter parameters for semi-automatic identification of man-made effects

| Frequency <br> $[\mathrm{Hz}]$ | Radius <br> $[\mathrm{m}]$ <br> $\log \rho_{\mathrm{a}} / \mathrm{d}_{\mathrm{a}}$ | $\log \rho_{\mathrm{a}} / \mathrm{d}_{\mathrm{a}}$ | Threshold <br> $\log \rho_{\mathrm{a}} / \mathrm{d}_{\mathrm{a}}$ | Number of <br> passes <br> $\log \rho_{\mathrm{a}} / \mathrm{d}_{\mathrm{a}}$ |
| :---: | :---: | :---: | :---: | :---: |
| 387 | $100 / 10$ | $0.17 / 3$ | $4 / 1$ | Type of <br> ann. $/$ Min. <br> 1,820 |
| $50 / 10$ | $0.15 / 5$ | $1 / 1$ | Min. / Both |  |
| 5,403 | $50 / 10$ | $0.10 / 5$ | $1 / 1$ | Min. / Min. |
| 8,389 | $50 / 10$ | $0.10 / 6$ | $1 / 1$ | Min. / Min. |
| 41,430 | $50 / 5$ | $0.15 / 6$ | $2 / 1$ | Min. / Min. |
| 133,200 | $50 / 5$ | $0.35 / 6$ | $7 / 1$ | Min. / Min. |

A topographic map and a Google Earth map are used to check whether the corresponding data segments are affected due to man-made sources and - if necessary - the data are reinstalled in manually

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selected areas. In order to close the remaining data gaps one can either apply gridding and resampling tools on the grids or use a spline interpolation along each survey line. Afterwards the HEM data are recalculated from the corrected half-space parameters. The measured HEM data are replaced by the calculated HEM data where the semi-automatic procedure has cut the data out.

### 5.3.6. Statistical Levelling

In order to identify and to correct zero-level errors in the HEM data a grid based micro-levelling (Table 16) is applied to the half-space parameters ( $\log \rho_{a}$ and $d_{a}$ ) of the parallel survey lines. The resulting error grids are resampled along the survey lines and the smoothed (spline filter: smoothness $=1.0$, tension $=0.5$ ) error channels are subtracted from the half-space parameters.

Table 16: Filter parameters for micro-levelling of $\log \rho_{a}$ and $d_{a}$

| Frequency [Hz] | Butterworth (high pass) Cut-off value, degree of filter $\log \rho_{\mathrm{a}} / \mathrm{d}_{\mathrm{a}}$ | Directional Cosine (pass) Azimuth, degree of function $\log \rho_{a} / d_{a}$ |
| :---: | :---: | :---: |
| 387 | $333 \mathrm{~m}, 8 / 333 \mathrm{~m}, 6$ | $90^{\circ}, 2 / 270^{\circ}, 1$ |
| 1,820 | $500 \mathrm{~m}, 8 / 333 \mathrm{~m}, 6$ | $90^{\circ}, 2 / 270^{\circ}, 2$ |
| 5,403 | 333 m, 10 / $333 \mathrm{~m}, 6$ | $90^{\circ}, 2 / 270^{\circ}, 2$ |
| 8,389 | 333 m, $10 / 400 \mathrm{~m}, 8$ | $90^{\circ}, 2 / 270^{\circ}, 2$ |
| 41,430 | 333 m, 12 / $400 \mathrm{~m}, 8$ | $90^{\circ}, 2 / 270^{\circ}, 2$ |
| 133,200 | 333 m, 10 / $400 \mathrm{~m}, 8$ | $90^{\circ}, 2 / 270^{\circ}, 2$ |

Strong HEM anomalies are normally smoothed by the two-dimensional lateral filtering of the microlevelling procedure. Therefore, grids where the local anomalies have been removed beforehand are used for micro-levelling, resulting in rather smooth apparent resistivity and apparent depth maps.

The tie lines are levelled afterwards using the levelled line grids as reference. The smoothed (lowpass filter: cut-off period $=50 \mathrm{~s}$, wave length $\approx 2000 \mathrm{~m}$ ) differences of levelled and unlevelled halfspace parameters are used to correct the tie-line data.

The levelled half-space parameter values are then converted to secondary field values ( $\mathrm{I}_{c}, \mathrm{Q}_{c}$ ) which are compared with the corresponding unlevelled values. Selected parts of the differences of the levelled and unlevelled values ( $\Delta \mathrm{I}=\mathrm{I}-\mathrm{I}_{c}, \Delta \mathrm{Q}=\mathrm{Q}-\mathrm{Q}_{c}$ ) are strongly smoothed using a non-linear filter and a smoothing spline interpolation. The selection is based on constant (data noise, system altitude) and dynamic ( $\mathrm{I}_{\text {spline }}, \mathrm{Q}_{\text {spline }}$ ) threshold values (Table 17). These interpolated smoothed differences are assumed to characterize the zero-level errors and they are used to correct the HEM data without losing details (Siemon, 2009). The levelling is done prior to the 1-D inversion of the HEM data.

Table 17: Filter parameters for the levelling of HEM data

| Type of filter | Filter parameters | Channel |
| :---: | :---: | :---: |
| Threshold | Cut-off value $\left(h_{-} l_{\text {kor }}\right): 300 \mathrm{~m}$ | $\Delta \mathrm{I}, \Delta \mathrm{Q}$ |
| Threshold | Cut-off value $\left(\mathrm{I}_{\text {noise }}, \mathrm{Q}_{\text {noise }}\right): 0.05$ | $\Delta \mathrm{I}, \Delta \mathrm{Q}$ |
| B-Spline | Smoothness $=1.0$, tension $=0.2$ | $\mathrm{I}, \mathrm{Q}$ |
| Non linear | Window length: 50 points $(\approx 200 \mathrm{~m})$, tolerance: 3.0 | $\Delta \mathrm{I}, \Delta \mathrm{Q}$ |
| B-Spline | Smoothness $=0.95$, tension $=0.5$ | $\Delta \mathrm{I}, \Delta \mathrm{Q}$ |

### 5.3.7. 1-D Inversion of the HEM Data

The model parameters of the 1-D inversion are the resistivities $\rho$ and thicknesses $t$ of a layered halfspace (Fig. 4), where the thickness of the underlying half-space is assumed to be infinite. Marquardt's inversion procedure is used (Sengpiel \& Siemon, 2000), which requires a starting model. This starting model is derived from the apparent resistivity vs. centroid depth values $\left(\rho_{\mathrm{a}}, \mathrm{z}^{*}\right)_{\mathrm{i}}, \mathrm{i}=$ 1,...,n(Fig. 5).


Fig. 5: Construction of starting models derived from apparent resistivity $\rho_{a}$, centroid depth $z^{*}$, and apparent depth $d_{a}$ of a five-frequency HEM data set

The standard model (Siemon, 2006) contains as many layers as frequencies. The layer resistivities are set equal to the apparent resistivities, the layer boundaries are chosen as the logarithmic mean of each two neighbouring centroid depth values. The standard starting model is constructed with respect to the number of frequencies used, i. e., the number of layers is given and the model layers correspond to the apparent resistivities and centroid depths of each frequency. The use of this kind of stating model enables the highest resolution, but also the highest sensitivity to calibration errors. Therefore, a modified starting model is constructed having an arbitrary number of layers. The resistivities and depths of the first and last layers are derived from the apparent resistivities and centroid depths of the highest and lowest frequencies, respectively. These confining layer boundaries can be shifted upwards or downwards. The thicknesses of the intermediate layers increase linearly and the resistivities are picked from an apparent resistivity sounding curve at the corresponding layer centres (on a $\log$ scale). Optionally, a resistive cover layer may be used. The thickness of the cover layer

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is derived from the apparent depth $\mathrm{d}_{\mathrm{a}}$ of the highest frequency. If this apparent depth value is less than a minimum layer thickness value, this value (e. g. 0.5 m ) is used.

As the laser altitude measurement are affected by the attitude of the HEM system resulting in measured bird altitudes that may be too high, the inversion may derive cover layers being to conductive. To overcome this difficulty, the bird altitude has to be variable. For this, a non-conducting cover layer is introduced representing a variable portion of the bird altitude. During the inversion the thickness of this layer (starting value: 5 m ) is optimised.

The inversion procedure is stopped when a given threshold is reached. This threshold is defined as the differential fit of the modelled data to the measured HEM data. Normally a $10 \%$ threshold is used; i. e., the inversion stops when the enhancement of the fit is less than $10 \%$.

The data of 5.5 kHz frequency were not used for inversion as they are obtained with a vertical coaxial coil system being sensitive to steeply dipping conductors (and also to external sources) whereas all others are obtained with horizontal coplanar systems.

The survey data were inverted with both types of starting models having five-layers and a highly resistive cover layer. A comparison with borehole data showed that the use of the modified starting model provided the best results.

### 5.3.8. Presentation of the Results

The HEM results are presented on maps and vertical resistivity sections (VRS). The maps are produced for the half-space parameters, apparent resistivity and centroid depth. Resistivities at ten depth levels (1-120 m below surface) are picked from 1-D inversion models and mapped. All the maps prepared from the results of this survey are listed in Section 6.3.

All data points used for the production of apparent resistivity and centroid depth maps are drawn as small black dots (flight lines). White dots mark areas of interpolated data. On the maps displaying resistivities at certain depths, the white dots inform about the number of interpolated data sets: the bigger the dot the more interpolated data were used for the inversion.

The VRS, also based on the 1-D inversion results, are produced for each of the survey lines. These vertical sections are constructed by placing the resistivity models for each sounding point along a survey profile next to each other using the topographic relief as base line (in m asl). The thickness of the bottom layer (substratum) is derived from the corresponding resistivity, but the minimum thickness is 10 m . The altitude of the EM sensor, information about the data processing, the fitting error of the inversion, and the HEM data, which are described in a legend, are plotted above the resistivity models.

### 5.4. Processing of Magnetic Data

### 5.4.1. Magnetic Total Field

The earth's total magnetic field $T$ at a point $r$ and at a time $t$, e. g., measured with an airborne system, is the sum of the following parts:

$$
\mathrm{T}(\mathrm{r}, \mathrm{t})=\mathrm{F}(\mathrm{r})+\Delta \mathrm{V}(\mathrm{t})+\Delta \mathrm{T}(\mathrm{r})+\delta_{\mathrm{T}}(\mathrm{r}, \mathrm{t})
$$

where
$\mathrm{F}(\mathrm{r})=$ geomagnetic main field (IGRF = International Geomagnetic Reference Field),
$\Delta \mathrm{V}(\mathrm{t})=$ diurnal variations of the earth's magnetic field,
$\Delta T(r)=$ the crustal field in the survey area,
$\delta_{\mathrm{T}}(\mathrm{r}, \mathrm{t})=$ anthropogenic part of the magnetic field.
The anomalies of the crustal field $\Delta \mathrm{T}(\mathrm{r})$, caused by rock magnetization, are of interest. While the IGRF $F(r)$, which can be calculated from table values, and the diurnal variations $\Delta V(t)$, which are recorded at the base station, can be subtracted from the measured total field, the anthropogenic part $\delta_{\mathrm{T}}(r, \mathrm{t})$ cannot be quantified independently. Therefore, the derived $\Delta \mathrm{T}$ values contain both the geogenic part and the disturbing anthropogenic part. Anthropogenic sources are located at the earth's surface (e. g., buildings, power lines, industrial sites). They are mostly locally constrained and thus can be identified using maps and other sources of information.

### 5.4.2. IGRF

The IGRF (International Geomagnetic Reference Field) can be calculated for any point on and above the earth's surface at a specific time on the basis of spherical harmonic coefficients, which are updated every five years by the International Association of Geomagnetism and Aeronomy (IAGA, 1992). The geomagnetic main field values of the survey area were calculated for each point using the IGRF-10 model from 2005 (IAGA, 2005).

### 5.4.3. Diurnal Variations

The base station for recording the time variant parts of the total magnetic field, the diurnal variations, was placed on the airport at $55^{\circ} 13^{\prime} 39^{\prime \prime} \mathrm{N}, 9^{\circ} 16^{\prime} 46^{\prime \prime} \mathrm{E}$ and 43 m asl. $\Delta \mathrm{V}(\mathrm{t})$ values are calculated as the measured value minus the IGRF value for the respective time and place. Possible disturbances of the base station recordings are eliminated using despiking and low-pass (filter width: 20) filters.

### 5.4.4. Levelling

After subtraction of the main field and diurnal variations from the measured magnetic field values, a statistical levelling is performed. The differences at the intersections of the flight lines and the tie lines are determined and averaged for each flight. The averaged values are then used to correct level errors that may occur in case of changes in the setup of the airborne or base station magnetic sensors during the survey.

Remaining, mostly small level errors may occur, inter alia, as result of different flight directions (heading errors) and are eliminated in a subsequent micro-levelling process. Micro-levelling is based
on gridded line data in which level errors are identified using two-dimensional Butterworth highpass (cut-off value: 500 m , degree: 4 ) and directional Cosine FFT (azimuth: $90^{\circ}$, degree: 1) filters. Result of the filtering process is an error grid which is sampled along the flight lines. The sampled error values are heavily smoothed using a B-Spline filter (smoothness: 1.05 , tension: 0.5 ) and then subtracted from the original data. Gridding of the levelled data yields a $\Delta T$ grid that is virtually free of level errors. Finally, the tie-line data are fit to the levelled line-data grid by removing possible offsets and trends in their differences.

In (partly) populated areas, grids of $\Delta \mathrm{T}$ values are mostly dominated by high-amplitude anthropogenic anomalies. These anomalies act as a source of disturbance during the micro-levelling process as well as during the identification of weak geogenic magnetic anomalies. Therefore, a semi-automatic filter procedure is applied to the data prior to micro-levelling. The procedure detects anomalous data in the $\Delta \mathrm{T}$ grid. Anomalies are detected when the differences of the grid values and their corresponding median values, which are calculated in circular areas shifted over the grid, exceed a given threshold. Manual interaction in the detection process is possible. The resulting grid is, as far as possible, freed from anthropogenic anomalies and is used as input for the micro-levelling process. The anthropogenic regions blanked in the finally levelled data may be re-introduced by applying error values interpolated from neighbouring data sections to them.

The levelled data are used to produce two final $\Delta \mathrm{T}$ grids. One grid $(\Delta \mathrm{T})$ is based on all data including anthropogenic anomalies. A second grid $\left(\Delta \mathrm{T}_{\mathrm{f}}\right)$ is based on the data freed from anthropogenic anomalies, and, additionally, is high-pass filtered. The high-pass filter (cutoff wavelength: 5 km ) is applied in order to remove large regional trends in the magnetic field that may mask weak and small-scale geogenic magnetic anomalies. The resulting grid is slightly smoothed using a two-dimensional median filter with a radius two grid cells.

### 5.4.5. Presentation of the Results

The maps produced to display the magnetic anomaly data are listed in Section 6.3. All data points used for map production are drawn as small black dots (flight lines). White dots mark areas of interpolated data.

### 5.5. Processing of Gamma-Ray Spectrometry Data

The natural gamma radiation of rocks and soil is mainly generated by the radioelements potassium, uranium, and thorium. According to the recommendations of the IAEA (2003), the spectrometry data recorded in the aircraft have to be converted to equivalent ground concentrations of these elements. This requires some preparatory procedures regarding spectrometer calibration and a number of data processing steps listed below.

Spectrometer calibration:

- Determination of cosmic and aircraft background count rates by means of flights over extensive water bodies
- Determination of stripping ratios for Compton scattering correction using calibration pads
- Determination of height attenuation and sensitivity coefficients by means of flights over a calibration range
- Determination of vegetation attenuation coefficients

Data processing:

- Energy calibration
- Reduction of count rate statistical noise
- Determination of detector height above ground and effective height
- Live time correction
- Background correction
- Compton (stripping) correction
- Height-attenuation reduction
- Calculation of equivalent ground concentrations


### 5.5.1. Energy Calibration

The spectral stability of gamma spectrometers is not perfect. Due to temperature effects, the mapping of energy peaks to correct channel positions may drift slightly during a survey flight. Therefore, an energy calibration is applied to the recorded spectra during data post processing. The channelenergy mapping of a spectrometer can be expressed as follows:

$$
\mathrm{ch}=\mathrm{E} / \mathrm{G}+\mathrm{offs},
$$

where

```
ch = channel number,
E = energy in keV,
G = gain constant of spectrometer in keV/channel,
offs = channel offset.
```

A 256 -channel spectrometer has a nominal gain constant of $12.0 \mathrm{keV} /$ channel and an offset of 0 channels. In order to determine the actual gain and offset of the spectrometer used, mean spectra are calculated for each flight line. The positions of the known energy peaks in the mean spectra ( K , $\mathrm{U}, \mathrm{Th})$ can then be used to calculate actual gain and offset of the instrument during each of the analysed time windows (flight lines). Based on these values, the recorded spectra are re-mapped to a nominal $12 \mathrm{keV} /$ channel raster.

### 5.5.2. Reduction of Statistical Noise

Due to a relatively large distance between the sources of radiation at the earth's surface and the radiation detector in the helicopter, count rates in airborne gamma-ray surveys are generally low. This results in a high portion of statistical noise present in the recorded spectra and, consequentially, also present in the calculated ground concentrations of radioelements. Therefore, a method for noise reduction developed by Hovgaard \& Grasty (1997) is applied to the data. The NASVD method (noise adjusted singular value decomposition) is based on a statistical analysis of all spectra recorded in a survey area and a reconstruction of noise reduced spectra using singular value decomposition routines. The procedure results in smoothed spectra reconstructed from five principal components, from which the count rates for the energy windows of interest (see Table 6) are determined. Furthermore, an adaptive filter (Mathis, 1987) for smoothing the count rate channels (filter width: 10) is applied.

### 5.5.3. Detector Height above Ground and Effective Height

Knowledge of the distance between the source of radiation (on the ground) and the detector (in the helicopter) is crucial for inferring ground concentrations of radioelements from airborne radiometric data correctly. The helicopter system used by BGR is equipped with two altimeters: a radar altimeter in the helicopter and a laser altimeter in the EM bird. Generally, the radar altimeter data are used to determine the detector's height above ground because it is installed on the same platform. However, the data from the laser altimeter are more precise and also contain information on the presence and eventually the thickness of vegetation on the ground beneath the system. Whereas, in forested areas, the radar altimeter detects the tree canopy and underestimates the ground distance, laser altimeter data allow both for estimation of the true ground distance and the vegetation height, which is required for an application of a vegetation attenuation correction (see Section 5.5.8). Therefore, a fusion of data from the radar and laser altimeters is performed, resulting in a more accurate estimation of the detector height above ground and an estimation of vegetation height.

In order to apply the radiometric analysis techniques, it is necessary to convert actual environmental conditions of the survey to standard conditions. This includes the adjustment of the measured ground clearance to standard temperature and pressure (STP conditions). The adjusted ground clearance value called "effective height" has the same mass of STP air between the ground and the helicopter as the actual one during data acquisition. The adjustment is applied according to IAEA (2003):

$$
\mathrm{h}_{\mathrm{e}}=\left(\mathrm{h}_{\mathrm{r}} * \mathrm{P} * \mathrm{~T}_{0}\right) /\left(\mathrm{P}_{0^{*}}\left(\mathrm{~T}+\mathrm{T}_{0}\right)\right),
$$

where
$h_{e}=$ effective height above ground level at STP [m],
$h_{r}=$ helicopter height above ground, determined from corrected radar altimeter data [m],
$\mathrm{T}_{0}=273.15 \mathrm{~K}$; freezing point of water on Kelvin scale,
$\mathrm{T}=$ air temperature $\left[{ }^{\circ} \mathrm{C}\right]$,
$\mathrm{P}_{0}=101.325 \mathrm{kPa}$; mean air pressure at sea level,
$\mathrm{P}=$ barometric pressure $[\mathrm{kPa}]$.

### 5.5.4. Live Time Correction

Gamma-ray spectrometers need a certain amount of time to process a pulse detected by the system. During that time, further incoming pulses are rejected. The amount of time the system is able to detect pulses ("live time") is recorded by the system. Due to the statistical nature of gamma radiation a correction of measured count rates in order to obtain count rate values for a nominal 1 s integration interval (IAEA, 2003) is easily achieved by using the following formula:

$$
\mathrm{N}=\mathrm{n} * 10^{3} / \mathrm{t}_{1},
$$

with
$\mathrm{N}=$ corrected count rate,
$\mathrm{n}=$ raw count rate,
$\mathrm{t}_{1}=$ system live time in milliseconds.

### 5.5.5. Background Radiation Correction

Cosmic radiation background is caused by high-energy ( $>3 \mathrm{MeV}$ ) cosmic ray particle interaction with the atmosphere. Another source of background radiation is the immanent radioactivity of the helicopter and its equipment. Background radiation distorts the measurements of geogenic radiation and has to be corrected for. The required correction coefficients are determined by means of flights over extensive water bodies at altitudes between 100 and 3500 m . The background correction is applied according to the following formula:

$$
\mathrm{N}=\mathrm{a}+\mathrm{b} * \mathrm{C}
$$

where
$\mathrm{N}=$ combined cosmic and aircraft background for each channel,
a = aircraft background for each channel,
$\mathrm{b}=$ cosmic rate stripping factor for each channel,
C = low-pass filtered cosmic channel (> 3 MeV ) count.
The values $a$ and $b$ were determined using data from test flights at different altitudes over the North Sea in 2008. For each channel K, U, Th, and TC (total count) a linear regression of the count rates for different altitude intervals and the filtered cosmic channel count rates revealed values for $a$ and $b$. The values are listed in Table 18.

Table 18: Aircraft background and cosmic stripping factors

| Channel | Aircraft background <br> a [cps] | Cosmic stripping factor <br> b |
| :---: | :---: | :---: |
| TC | 31.09 | 0.722 |
| K | 5.51 | 0.041 |
| U | 0.48 | 0.033 |
| Th | 0.33 | 0.041 |

### 5.5.6. Compton Correction

Compton scattering leads to certain amounts of radiation from one energy window being scattered into other energy windows. For example, some amount of thorium radiation will be scattered into lower energy windows such as uranium and potassium. The removal of these effects (Compton correction) is done using so-called stripping ratios. These coefficients describe the magnitudes of scatter between the energy windows of interest. They were determined in 2008 using portable calibration pads (Grasty et al., 1991) and are listed in Table 19.

Table 19: Stripping ratios

|  | Stripping ratio | Value |
| :---: | :---: | :---: |
| $\mathrm{Th} \rightarrow \mathrm{U}$ | $\alpha$ | 0.2485 |
| $\mathrm{Th} \rightarrow \mathrm{K}$ | $\beta$ | 0.3852 |
| $\mathrm{U} \rightarrow \mathrm{K}$ | $\gamma$ | 0.6599 |
| $\mathrm{U} \rightarrow \mathrm{Th}$ | a | 0.0395 |

The values of $\alpha, \beta, \gamma$ increase with altitude of the helicopter above ground level and have to be corrected on the base of STP equivalent altitude according to the following factors (see IAEA, 2003):

$$
\begin{aligned}
& \alpha_{e}=\alpha+0.00049 * h_{e} \\
& \beta_{e}=\beta+0.00065 * h_{e} \\
& \gamma_{e}=\gamma+0.00069 * h_{e}
\end{aligned}
$$

with
$h_{e}=$ equivalent height above ground level at STP in metres.
To obtain the net count rates of the particular energy windows, the stripping ratios are applied to the data:

$$
\begin{gathered}
\mathrm{N}_{\mathrm{Th} \text { (corr) }}=\left(\mathrm{N}_{\mathrm{Th}}-\mathrm{a} \mathrm{~N}_{\mathrm{U}}\right) /(1-\mathrm{a} \alpha) \\
\mathrm{N}_{\mathrm{U} \text { (corr) }}=\left(\mathrm{N}_{\mathrm{U}}-\alpha \mathrm{N}_{\mathrm{Th}}\right) /(1-\mathrm{a} \alpha) \\
\mathrm{N}_{\mathrm{K}(\text { corr })}=\mathrm{N}_{\mathrm{K}}-\beta \mathrm{N}_{\mathrm{Th} \text { (corr) }}-\gamma \mathrm{N}_{\mathrm{U} \text { (corr) }}
\end{gathered}
$$

where $\mathrm{N}_{\mathrm{Th}}, \mathrm{N}_{\mathrm{K}}, \mathrm{N}_{\mathrm{U}}$ represent the background and STP corrected count rates, $\mathrm{N}_{\mathrm{Th}(\text { corr })}, \mathrm{N}_{\mathrm{U}(\text { (corr) }}, \mathrm{N}_{\mathrm{K}(\text { corr) }}$ are the stripping corrected count rates, and $\alpha, \beta, \gamma$, a are the STP corrected stripping ratios. No Compton correction is applied to the total radiation values (see IAEA, 2003).

### 5.5.7. Height-Attenuation Reduction

The intensity of gamma radiation measured in airborne surveys varies approximately exponentially with height. In order to estimate count rates at a nominal survey height of 80 m , the following formula is used:

$$
\mathrm{N}_{\mathrm{s}}=\mathrm{N}_{\mathrm{m}} * \mathrm{e}^{-\mu\left(\mathrm{h}_{\mathrm{o}}-\mathrm{h}_{\mathrm{e}}\right)},
$$

where
$\mu=$ window attenuation coefficient (per metre),
$\mathrm{N}_{\mathrm{m}}=$ observed count rate at STP effective height $\mathrm{h}_{\mathrm{e}}$,
$\mathrm{N}_{\mathrm{s}}=$ corrected count rate for the nominal survey height $\mathrm{h}_{0}$.
The values (Table 20) were determined from data acquired at different heights over the Allentsteig (Austria) calibration range in 2003.

Table 20: Height attenuation coefficients

| Window | Height attenuation coefficient $\boldsymbol{\mu}$ <br> (per metre at STP) |
| :---: | :---: |
| K | 0.007733 |
| U | 0.008132 |
| Th | 0.005784 |
| TC | 0.006468 |

### 5.5.8. Radioelement Concentrations and Exposure Rate

IAEA (2003) recommends converting the count rates for the three radioelements into surface concentrations and exposure rates at ground level. The advantage is that the results of measurements with different instruments (e. g. with different crystal volumes) can be compared with each other. Conversion between count rates and concentrations is done using sensitivity coefficients (Tab. 21):

$$
\mathrm{C}=\mathrm{N}_{\mathrm{s}} / \mathrm{S},
$$

with
$\mathrm{C}=$ element concentration ( K in \%, eU in ppm, eTh in ppm),
$\mathrm{N}_{\mathrm{s}}=$ count rate for each window (after height attenuation and stripping),
$S=$ broad source sensitivity for the spectral window.
The calculated concentrations are expressed as equivalent concentrations eU and eTh (in ppm) and as concentrations of K (in \%).

Table 21: Sensitivity coefficients

| Sensitivity |  |
| :--- | :--- |
| $1 \% \mathrm{~K}$ | $=28.42 \mathrm{cps}$ |
| 1 ppm eTh | $=1.96 \mathrm{cps}$ |
| 1 ppm eU | $=2.92 \mathrm{cps}$ |

The sensitivities (Table 21) were determined over the Allentsteig (Austria) calibration range. Concentrations calculated this way refer to an infinitely extended and permanently radiating plane. They may differ from the actual concentrations of the elements at ground surface, especially in areas of irregularly distributed radiation sources and under wet conditions. Furthermore, the presence of atmospheric radon may vary considerably during a survey. Radon can spoil radiometric data, in par-
ticular uranium concentrations inferred from count rates, because its radiation is detected in the uranium energy window. Presently, there is no correction of the effect of radon radiation on airborne gamma-ray measurements implemented in our radiometric data processing routines. Absolute values of uranium concentrations indicated in the maps are therefore to be regarded with caution.

The calculated concentrations will also be erroneous in areas where vegetation, mostly in the form of trees in forested areas, absorbs part of the radiation from the ground. A vegetation correction can be applied to the data assuming that the attenuation of gamma radiation varies exponentially with vegetation thickness and vegetation thicknesses are known:

$$
C_{H}=C_{0} * e^{-\mu \cdot H},
$$

where
$C_{0}=$ element concentration at the ground,
$C_{H}=$ element concentration determined in the presence of vegetation,
$\mathrm{H}=$ thickness of vegetation,
$\mu=$ linear attenuation coefficient of vegetation.
Values for $\mu$ (Table 22) were determined empirically using extensive data sets acquired over northern Germany. Vegetation thicknesses were inferred from laser altimeter data. In addition to the three radioelements, an attenuation coefficient for the total counts energy window was determined empirically.

Table 22: Linear attenuation coefficients $\mu$ for vegetation.

| Element | $\mu$ |
| :---: | :---: |
| K | 0.012 |
| U | 0.008 |
| Th | 0.011 |
| TC | 0.006 |

The ground level exposure rate is calculated as a function of the K , U , and Th concentrations after application of the vegetation correction:

$$
\mathrm{E}=1.505 * \mathrm{~K}+0.653 * \mathrm{eU}+0.287 * \mathrm{eTh},
$$

with
$\mathrm{E}=$ ground level exposure rate $[\mu \mathrm{R} / \mathrm{h}]$
using the following conversions (IAEA, 2003):
$1 \% \mathrm{~K}=1.505 \mu \mathrm{R} / \mathrm{h}$,
$1 \mathrm{ppm} \mathrm{eU}=0.653 \mu \mathrm{R} / \mathrm{h}$,
1 ppm eTh $=0.287 \mu \mathrm{R} / \mathrm{h}$.

### 5.5.9. Data Levelling and Smoothing

Prior to the vegetation correction and the determination of exposure rates, a statistical levelling of the concentration and total count data is performed. The differences at the intersections of the flight lines and the tie lines are determined and averaged for each flight. The averaged values are then used to correct level errors that may occur in case of environmental changes (e. g. humidity, radon abundance) during the survey.

Remaining, mostly small level errors are eliminated in a subsequent micro-levelling process based on gridded line data in which level errors are identified using two-dimensional Butterworth high-pass (cut-off value: 300 m , degree: 8 ) and directional Cosine FFT (azimuth: $90^{\circ}$, degree: 1) filters. Result of the filtering process is an error grid which is sampled along the flight lines. The sampled error values are heavily smoothed using a B-Spline filter (smoothness: $0.65 / 0.5$, tension: $0.5 / 0.5$ for lines / tie lines) and then subtracted from the original data. Gridding of the levelled data yields grids that are virtually free of level errors. Finally, the tie-line data are fit to the levelled line data-grid. This is done by calculating the difference (error) between the values of the levelled line-data grid and the tie-line data, spline smoothing the error and subtracting it from the tie-line data.

Grids of the finally levelled data are slightly smoothed using a two-dimensional median filter of radius 1 (for potassium, thorium and total count) or 2 (for uranium) grid cells. The filtered grids are sampled along the flight path and the sampled data are used as input for vegetation attenuation correction and exposure rate calculations (see previous chapter).

### 5.5.10. Presentation of the Results

The results of the gamma-ray survey are presented as maps of the equivalent concentrations of the radioelements potassium, uranium, and thorium, total counts, and the ground level exposure rate. The maps produced to display the radiometric data are listed in Section 6.3. All data points used for map production are drawn as small black dots (flight lines). White dots mark areas of interpolated data.

Due to technical problems during one flight (13402) no data on the southernmost E-W lines (50.168.1) were available and, thus, the final grids differ from the grids of the other methods.

## 6. Cartographic Work

### 6.1. Topographic Map

A topographic map was produced as the base map for all thematic maps displaying the airborne geophysical results. A scale of 1:25,000 was chosen for the survey area. An UTM coordinate grid, based on the WGS 84 ellipsoid, is included on the topographic maps. Table $\mathbf{2 3}$ contains the corner coordinates of the map sheet.

Table 23: Coordinates of the corners of the 1:25,000 Vojens topographic map sheet

| Map corners | Geographic coordinates <br> (WGS 84) |  | UTM WGS 84 coordinates <br> (Zone 32N) |  |
| :---: | :---: | :---: | :---: | :---: |
|  | Easting | Northing | Easting | Northing |
| SW | $9^{\circ} 09^{\prime} 26^{\prime \prime}$ | $55^{\circ} 13^{\prime} 03$ | 510000 | 6119000 |
| NW | $9^{\circ} 09^{\prime} 27^{\prime \prime}$ | $55^{\circ} 18^{\prime} 26^{\prime \prime}$ | 510000 | 6129000 |
| NE | $9^{\circ} 21^{\prime} 44^{\prime \prime}$ | $55^{\circ} 18^{\prime} 25^{\prime \prime}$ | 523000 | 6129000 |
| SE | $9^{\circ} 21^{\prime} 41^{\prime \prime}$ | $55^{\circ} 13^{\prime} 01^{\prime \prime}$ | 523000 | 6119000 |

The map is based on the »Topographic Map of Denmark 1:25,000«, © Kort \& Matrikelstyrelsen, København. The following map sheets were used:

## 1212 IV NØ, 1212 IV NV, 1212 IV SØ, 1212 IV SV.

The map has a digitally constructed border and tick marks indicating coordinates in the WGS 84 coordinate system. The grey-shading of the topography of the thematic map has a screen density of $50 \%$ of the original digital topographic map.

### 6.2. Map Production with GEOSOFT and GIS Software

The grids for the geophysical thematic maps were produced using the software package GEOSOFT OASIS montaj 7.2. Table $\mathbf{2 4}$ shows the grid parameters used for the Vojens survey.

Table 24: Grid parameters

| Parameter | Value |
| :---: | :---: |
| Gridding method | Minimum curvature |
| Grid size [m] | 20 |
| Search radius [m] | 20 |
| Internal tension (0-1) | 0 |
| Cell extend beyond data | 5 |
| Log option | $\log \rho$ (else linear) |

The final maps including geophysical, topographical and legend information are prepared using the program ESRI ArcGIS 9.3.1. A special plug-in provided by GEOSOFT for ArcGIS (available on DVD or http://www.geosoft.com/resources/releasenotes/plugins/arcGISplugin.asp) is necessary to import and display the GEOSOFT grids as a layer in ArcMap. Adobe Acrobat 9.3 is used for preparing the PDF documents.

### 6.3. Thematic Maps

Coloured geophysical thematic maps (Table 25, Appendix IV) were produced at a scale of 1:25,000 for each parameter of interest.

HEM: Apparent resistivities and centroid depths
at $387 \mathrm{~Hz}, 1,820 \mathrm{~Hz}, 5,403 \mathrm{~Hz}, 8,389 \mathrm{~Hz}, 41,430 \mathrm{~Hz}$, and $133,200 \mathrm{~Hz}$;
Resistivities at 1, 3, 5, 10, 20, 30,50, 70,90 and 120 m below ground level (bgl);
HMG: Anomalies of the total magnetic field;
Anomalies of the total magnetic field, high-pass filtered, anthropogenic anomalies removed;
HRD: Equivalent concentrations of the radioelements potassium, uranium, and thorium, total count and ground level exposure rate.

The digital topographic map was used as base map. The surveyed flight lines are plotted in black/white containing information about the quality of the data. In addition, flight-line and elevations maps were produced.

The flight-line maps show the position of the surveyed profiles on the topographic maps. The corresponding line number is shown at the end of a profile at which the flight for that profile commenced. Positions of selected time marks (records), e. g., every 100th, are marked with an " $\times$ ". Every tenth plotted time mark is labelled with its number. The flight-line maps permit fast and easy correlation of data from profiles and vertical sections and their position in the survey area.

Digital elevation models (DEM) are derived from the corrected and levelled difference of bird elevation laser altitude. The elevation map also contains the topographic base map and the flight lines.

## 7. Archiving

All data sets and plots are stored on DVD and archived at BGR section B 2.1-Geophysical Exploration - Resources and Near Surface Processes. The data formats of processed data are described in Appendix II. A technical report, the vertical sections, and the thematic maps (as PDF files) are stored together with the final data (ASCII-coded in GEOSOFT-XYZ format) on a DVD (Table 25). A copy of this DVD is attached to this report. The content is listed in Appendix III. Appendix IV and Appendix V contain copies of all maps and vertical resistivity sections, respectively, reduced to smaller scales fitting the A4 format of this report.

Table 25: Content of the $D V D$

| Directory |  | Description of content |
| :---: | :---: | :---: |
| \Adobe Acrobat |  | Adobe ${ }^{\otimes}$ Acrobat Reader in diverse versions for popular system software |
| $\backslash$ Report |  | Technical report of the project in PDF format |
| $\begin{aligned} & \text { !í } \\ & \text { !í } \end{aligned}$ | \HEM | ASCII file with all raw data (HEM134_RAW.xyz) <br> ASCII file with all processed data (HEM134_DAT.xyz) <br> ASCII file with all derived parameters (HEM134_APP.xyz) <br> ASCII file with results of the 1-D inversion (HEM134_INV.xyz) |
|  | \MAG | ASCII file with data of the total magnetic field, IGRF, base station data, diurnal variations etc. (HMG134.xyz) |
|  | \SCI | ASCII file with data of the equivalent concentrations of potassium, uranium and thorium and the total radiation (HRD134.xyz) |
|  | \HEM | Apparent resistivity maps and centroid depth maps at a scale of 1:25,000 for the frequencies $387 \mathrm{~Hz}, 1,820 \mathrm{~Hz}, 5,403 \mathrm{~Hz}, 8,389 \mathrm{~Hz}, 41,430 \mathrm{~Hz}$, $133,200 \mathrm{~Hz}$ in PDF format <br> Resistivity maps at a scale of $1: 25,000$ at $1,3,5,7,10,20,30,50,70,90$ and 120 m below surface based on five-layer inversion results in PDF format |
|  | \HMG | Magnetic anomalies maps at a scale of 1:25,000 in PDF format |
|  | \HRD | Maps of the equivalent concentrations of the radioelements potassium, uranium, and thorium, the total radiation and the ground level exposure rate at a scale of 1:25,000 in PDF format |
|  | $\backslash$ Flight lines | Flight line map with topography at a scale of 1:25,000 in PDF format |
|  | \DEM | Digital elevation model at a scale of 1:25,000 in PDF format |
|  | \ArcGIS | Map projects for ArcGIS 9.3.1 (*mxd) incl. legends (*bmp), Raster data TK 50 (GRID) and Geosoft-Plugin for ArcGIS |
| \Vertical sections |  | Vertical resistivity section based on five-layer inversion results for each profile of the survey area at a horizontal scale of 1:25,000 and at a vertical scale of 1:2,500 in PDF format |

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# BUNDESANSTALT FÜR GEOWISSENSCHAFTEN UND ROHSTOFFE BGR, HANOVER 



## Appendix I

## Survey Area 134 - Vojens

Airport: $\quad$ Vojens, Elevation: $140 \mathrm{ft} / 43 \mathrm{~m}$

## Survey parameters:

| Line separation: | Lines: 80 m | Tie lines: 500 m |  |
| :--- | :--- | :--- | :--- |
| Line direction: | Lines: | $90^{\circ}$ | Tie lines: |
| 0 |  |  |  |

Table A-1: Flight table

| Flight | Date | Time (UTC) <br> Start - End | Lines | Remarks |
| :---: | :---: | :---: | :---: | :---: |
| 13401 | 23.06.09 | 07:58-10:18 | $\begin{gathered} 1.1 \mathrm{~W} \\ 4.1 \mathrm{E} \\ 7.1 \mathrm{~W} \\ 10.1 \mathrm{E} \\ 13.1 \mathrm{~W} \\ 16.1 \mathrm{E} \\ 19.1 \mathrm{~W} \\ 22.1 \mathrm{E} \\ 25.1 \mathrm{~W} \\ 28.1 \mathrm{E} \\ 31.1 \mathrm{~W} \\ 34.1 \mathrm{E} \\ 37.1 \mathrm{~W} \\ 40.1 \mathrm{E} \\ 43.1 \mathrm{~W} \\ 46.1 \mathrm{E} \\ 49.1 \mathrm{~W} \\ 41.1 \mathrm{E} \end{gathered}$ | HELIDAS: SYS14; BKS36a; <br> EM: EM4 shortly out after line 49.1 <br> Magnetometer: ok <br> Spectrometer: ok <br> Video: <br> Weather: sunny, $12^{\circ} \mathrm{C}$, no wind <br> Line-km: 156 |


| 13402 | 23.06.09 | 13:39-15:28 |  | HELIDAS: SYS14; BKS36a; <br> EM: EM4 shortly out on and after line 57.1 <br> Magnetometer: ok <br> Spectrometer: recording of spectral data not ok <br> Video: <br> Weather: sunny, $26^{\circ} \mathrm{C}$, no wind <br> Line-km: 152 |
| :---: | :---: | :---: | :---: | :---: |
| 13403 | 24.06.09 | 7:33-9:53 | 48.1 E 47.1 W 45.1 E 41.2 W 44.1 E 45.2 W 42.1 E 39.1 W 38.1 E 35.1 W 36.1 E 33.1 W 32.1 E 29.1 W 30.1 E 27.1 W 26.1 E 23.1 W TL (34.3) | HELIDAS: SYS14; BKS36a; <br> EM: EM4 shortly out on and after line 23.1 <br> Magnetometer: ok <br> Spectrometer: ok <br> Video: <br> Weather: sunny, $15^{\circ} \mathrm{C}$, no wind <br> Line-km: 136 <br> Line 45.1 flown in opposite direction, repeated (45.2) |


| 13404 | 24.06.09 | 12:26-13:36 | 16.9 N <br> 15.9 S <br> 14.9 N <br> 13.9 S <br> 12.9 N <br> 11.9 S <br> 24.1 E <br> 21.1 W <br> 10.9 N | HELIDAS: SYS14; BKS36a; <br> EM: EM4 shortly out on tie line 11.9 System failure after line 21.1 at record 34480 (near radar station on airport) <br> Magnetometer: ok <br> Spectrometer: ok <br> Video: <br> Weather: sunny, $23^{\circ} \mathrm{C}$, no wind <br> Line-km: 65 |
| :---: | :---: | :---: | :---: | :---: |
| 13405 | 24.06.09 | 13:37-14:55 | 1.9 S 2.9 N 3.9 S 4.9 N 5.9 S 6.9 N 7.9 S 8.9 N 9.9 S 10.8 N $\mathrm{TL}(34.5)$ | HELIDAS: SYS14; BKS36a; <br> EM: ok <br> Magnetometer: ok <br> Spectrometer: ok <br> Video: <br> Weather: sunny, $23^{\circ} \mathrm{C}$, no wind <br> Line-km: 70 <br> Continuation of flight 13404 |
| 13406 | 25.06.09 | 7:46-9:52 | 20.1 E 17.1 W 18.1 E 15.1 W 14.1 E 11.1 W 12.1 E 9.1 W 8.1 E 5.1 W 6.1 E 3.1 W 2.1 E 41.3 W $\mathrm{TL}(34.6)$ | HELIDAS: SYS14; BKS36a; <br> EM: EM4 shortly out after line 15.1 <br> Magnetometer: ok <br> Spectrometer: ok <br> Video: <br> Weather: sunny, $18^{\circ} \mathrm{C}$, no wind <br> Line-km: 112 |

## Appendix II

## Final Data Format Description

## A) Electromagnetics

Description of the four ASCII-coded data files containing the final (levelled) data of a helicopter-borne electromagnetic (HEM) survey

## General HEADER:

/BGR HEADER (SHORT VERSION):
/
/AREANAME
/VOJENS
/AREACODE
/134
/C MERIDIAN, ZONE, REFERENCE SYSTEM
/ 932 WGS84
/ELLIPSOID FOR LON AND LAT
/WGS84
/BIRD
/61
/NUMFREQ
/ 6
/FREQUENCY
/ 387.00 $1820.00 \quad 5403.008389 .0041430 .00133200 .00$
/COILGEOMETRY
$\begin{array}{lllllll}\text { / } & 1.00 & 1.00 & 4.00 & 1.00 & 1.00 & 1.00\end{array}$
/COILSEPERATION

|  | 7.94 | 7.93 | 9.06 | 7.93 | 7.91 | 7.92 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |

/TOWCABLE
/ 40.00
/DUMMY
/ -999.990
/DECIMATIONVALUE
/ 1
/PRIVTEXT
(up to five lines of comment may be written here)

## 1) Raw data: HEM134_RAW.XYZ

Example:
/Unprocessed data
//Flight 13401
//Date 2009/06/23
Random 0


 $\begin{array}{lllllll}526901.006124502 .009 .423349 & 55.266256 & 2.00 & 75744.20 & 1611.10 & 463.2\end{array}$

| 490.78 | 537.91 | 289.92 | 6.58 | 0.01 | $\ldots$. | 1.94 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |

In this data file all secondary field values are stored in the order of the following description:

| Channel | Unit | Remarks |
| :---: | :---: | :---: |
| X | m | UTM easting in m (WGS 84, Zone 32N), these coordinates have a false easting of 500000 metres |
| Y | m | UTM northing in m (WGS 84, Zone 32N), these coordinates have no false northing |
| LON | - | geographic longitude, reference system WGS 84 |
| LAT | - | geographic latitude, reference system WGS 84 |
| RECORD |  | time mark increasing by 1 every 0.1 seconds |
| UTC_TIME | hhmmss.s | GPS time (UTC) |
| ALTR | ft | radar altimeter reading (helicopter) |
| ALTL_FP | m | laser altimeter reading (bird) |
| ZHG_BIRD_RAW | m | GPS elevation of the bird, reference system WGS 84 |
| ZHG_HELI_RAW | m | GPS elevation of the helicopter, reference system WGS 84 |
| ALTB | ft | barometric elevation of the helicopter |
| EM1I | ppm | raw value of the inphase component at the frequency $\mathrm{f}=41,430 \mathrm{~Hz}$ |
| EM1Q | ppm | raw value of the quadrature component at the frequency $\mathrm{f}=41,430 \mathrm{~Hz}$ |
| EM2I | ppm | raw value of the inphase component at the frequency $f=8,389 \mathrm{~Hz}$ |
| EM2Q | ppm | raw value of the quadrature component at the frequency $f=8,389 \mathrm{~Hz}$ |
| EM3I | ppm | raw value of the inphase component at the frequency $\mathrm{f}=387 \mathrm{~Hz}$ |
| EM3Q | ppm | raw value of the quadrature component at the frequency f $=387 \mathrm{~Hz}$ |
| EM4I | ppm | raw value of the inphase component at the frequency $f=133,200 \mathrm{~Hz}$ |
| EM4Q | ppm | raw value of the quadrature component at the frequency $f=133,200 \mathrm{~Hz}$ |
| EM5I | ppm | raw value of the inphase component at the frequency $f=1,820 \mathrm{~Hz}$ |
| EM5Q | ppm | raw value of the quadrature component at the frequency $f=1,820 \mathrm{~Hz}$ |
| EM6I | ppm | raw value of the inphase component at the frequency $f=5,403 \mathrm{~Hz}$, converted to horizontal coplanar |
| EM6Q | ppm | raw value of the quadrature component at the frequency $f=5,403 \mathrm{~Hz}$, converted to horizontal coplanar |


| EM1_FREQ | Hz | frequency of EM1 channels ( $\mathrm{f}=41,430 \mathrm{~Hz}$ ) |
| :---: | :---: | :---: |
| EM2_FREQ | Hz | frequency of EM2 channels ( $\mathrm{f}=8,389 \mathrm{~Hz}$ ) |
| EM3_FREQ | Hz | frequency of EM3 channels ( $\mathrm{f}=387 \mathrm{~Hz}$ ) |
| EM4_FREQ | Hz | frequency of EM4 channels ( $\mathrm{f}=133,200 \mathrm{~Hz}$ ) |
| EM5_FREQ | Hz | frequency of EM5 channels ( $\mathrm{f}=1,820 \mathrm{~Hz}$ ) |
| EM6_FREQ | Hz | frequency of EM6 channels ( $\mathrm{f}=5,403 \mathrm{~Hz}$ ) |
| CPPL |  | power-line detector |
| CPSP |  | sferics detector |

Remarks:
Lines starting with "/"
Lines starting with "//"
Lines starting with "Random"
flight number and date,
original flights.

Original vertical coaxial data are indicated by -0.25 (instead of 4.00 for converted data):
/COILGEOMETRY

| $/$ | 1.00 | 1.00 | -0.25 | 1.00 | 1.00 | 1.00 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |

General Remarks for the next three data sets:
Lines starting with "/"
Lines starting with "//"
Lines starting with "Line"
Lines starting with "Tie"
comment,
flight number and date
lines,
tie lines.
geozentrum hannover Survey Area Vojens, Denmark, 2009

## 2) Data: HEM134_DAT.XYZ

Example:
/Processing by A. Ullmann (BGR) using Oasis montaj
/Levelled data

| X Y | LON | LAT | RECORD | UTC | TOPO | H_RADAR | __LASE | BIRD_NN | H_BARO | REAL_1 | UAD |  | REAL_6 | QUAD_6 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| //Flight 13401 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| //Date 2009/06/23 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Line 1.1 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 5213556127061 | 9.336263 | 55.289514 | 6640 | 80848.0 | 50.42 | 101.28 | 97.1 | 1147.53 | 145.64 | 3.390 | 0.468 |  | 4.046 | 6.518 |
| 5213526127060 | 9.336212 | 55.289513 | 6641 | 80848.1 | 50.32 | 100.26 | 96.60 | 146.92 | 143.41 | 3.390 | 0.468 |  | 4.046 | 6.518 |
| 5213496127060 | 9.336161 | 55.289511 | 6642 | 80848.2 | 50.22 | 99.78 | 96.10 | 146.32 | 144.16 | 3.004 | 1.075 |  | 2.898 | 4.980 |

In this data file all necessary position parameters and secondary field values are stored in the order of the following description:

| Channel | Unit | Remarks |
| :---: | :---: | :---: |
| X | m | UTM easting in m (WGS 84, Zone 32N), these coordinates have a false easting of 500000 metres |
| Y | m | UTM northing in m (WGS 84, Zone 32N), these coordinates have no false northing |
| LON | ${ }^{\circ}$ | geographic longitude, reference system WGS 84 |
| LAT | 。 | geographic latitude, reference system WGS 84 |
| RECORD |  | time mark increasing by 1 every 0.1 seconds |
| UTC_TIME | hhmmss.s | GPS time (UTC) |
| TOPO | m | levelled topographic elevation (in metre above sea level), derived from the difference of bird elevation (BIRN_NN) and bird altitude (H_LASER) |
| H_RADAR | m | smoothed value of the radar altitude minus the effective cable length ( 40 m ) from the helicopter to the bird, corresponds to the bird altitude |
| H_LASER | m | smoothed value of the laser altimeter, corresponds to the bird |
| BIRD_NN | m | smoothed bird elevation (in m asl = metre above sea level), reference system: WGS84 |
| H_BARO | m | processed value of the barometric sensor minus the effective cable length ( 40 m ) from the helicopter to the bird |
| REAL_1 | ppm | processed value of the inphase component at the frequency $\mathrm{f}=387 \mathrm{~Hz}$ |
| QUAD_1 | ppm | processed value of the quadrature component at the frequency $f=387 \mathrm{~Hz}$ |
| REAL_2 | ppm | processed value of the inphase component at the frequency $f=1,820 \mathrm{~Hz}$ |
| QUAD_2 | ppm | processed value of the quadrature component at the frequency $f=1,820 \mathrm{~Hz}$ |
| REAL_3 | ppm | processed value of the inphase component at the frequency $\mathrm{f}=5,403 \mathrm{~Hz}$, converted to horizontal coplanar |
| QUAD_3 | ppm | processed value of the quadrature component at the frequency $\mathrm{f}=5,403 \mathrm{~Hz}$, converted to horizontal coplanar |
| REAL_4 | ppm | processed value of the inphase component at the frequency $f=8,389 \mathrm{~Hz}$ |
| QUAD_4 | ppm | processed value of the quadrature component at the frequency $\mathrm{f}=8,389 \mathrm{~Hz}$ |
| REAL_5 | ppm | processed value of the inphase component at the frequency $f=41,430 \mathrm{~Hz}$ |
| QUAD_5 | ppm | processed value of the quadrature component at the frequency $f=41,430 \mathrm{~Hz}$ |
| REAL_6 | ppm | processed value of the inphase component at the frequency $f=133,200 \mathrm{~Hz}$ |
| QUAD_6 | ppm | processed value of the quadrature component at the frequency $f=133,200 \mathrm{~Hz}$ |

## 3) Half-space parameters: HEM134_APP.XYZ

Example:
/Processing by A. Ullmann (BGR) using Oasis montaj
/Levelled data
/ X Y LON LAT RECORD UTC TOPO H_RADAR H_LASER BIRD_NN H_BARO RHOA_1 KDA_1 ZST_1... RHOA_6 KDA_6 ZST_6
//Flight 13401
//Date 2009/06/23
Line 1.1

| 521355 | 6127061 | 9.336263 | 55.289514 | 6640 | 80848.0 | 50.42 | 101.28 | 97.11 | 147.53 | 145.64 | 28.60 | 11.56 | 80.06 | $\ldots$ | 38.06 | -0.01 | 4.24 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 521352 | 6127060 | 9.336212 | 55.289513 | 6641 | 80848.1 | 50.32 | 100.26 | 96.60 | 146.92 | 143.41 | 28.51 | 11.40 | 79.79 | $\ldots$ | 40.19 | -0.04 | 4.33 |


| 521352 | 6127060 | 9.336212 | 55.289513 | 6641 | 80848.1 | 50.32 | 100.26 | 96.60 | 146.92 | 143.41 | 28.51 | 11.40 | 79.79 | $\ldots$ | 40.19 | -0.04 | 4.33 |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| 521349 | 6127060 | 9.336161 | 55.289511 | 6642 | 80848.2 | 50.22 | 99.78 | 96.10 | 146.32 | 144.16 | 28.44 | 11.24 | 79.54 | $\ldots$ | 42.17 | -0.09 | 4.39 |

In this data file all necessary position parameters and half-space parameters are stored in the order of the following description:

| Channel | Unit | Remarks |
| :---: | :---: | :---: |
| X | m | UTM easting in m (WGS 84, Zone 32N), these coordinates have a false easting of 500000 metres |
| Y | m | UTM northing in m (WGS 84, Zone 32N), these coordinates have no false northing |
| LON | ${ }^{\circ}$ | geographic longitude, reference system WGS 84 |
| LAT | 。 | geographic latitude, reference system WGS 84 |
| RECORD |  | time mark increasing by 1 every 0.1 seconds |
| UTC_TIME | hhmmss.s | GPS time (UTC) |
| TOPO | m | levelled topographic elevation (in metre above sea level), derived from the difference of bird elevation (BIRN_NN) and bird altitude (H_LASER) |
| H_RADAR | m | smoothed value of the radar altitude minus the effective cable length ( 40 m ) from the helicopter to the bird, corresponds to the bird altitude |
| H_LASER | m | smoothed value of the laser altimeter, corresponds to the bird |
| BIRD_NN | m | smoothed bird elevation (in m asl = metre above sea level), reference system: WGS84 |
| H_BARO | m | filtered value of the barometric sensor minus the effective cable length ( 40 m ) from the helicopter to the bird |
| RHOA_1 | $\Omega \mathrm{m}$ | apparent resistivity at the frequency $\mathrm{f}=387 \mathrm{~Hz}$ |
| KDA_1 | m | apparent depth at the frequency $\mathrm{f}=387 \mathrm{~Hz}$ |
| ZST_1 | m | centroid depth at the frequency $\mathrm{f}=387 \mathrm{~Hz}$ |
| RHOA_2 | $\Omega \mathrm{m}$ | apparent resistivity at the frequency $\mathrm{f}=1,820 \mathrm{~Hz}$ |
| KDA_2 | m | apparent depth at the frequency $\mathrm{f}=1,820 \mathrm{~Hz}$ |
| ZST_2 | m | centroid depth at the frequency $\mathrm{f}=1,820 \mathrm{~Hz}$ |
| RHOA_3 | $\Omega \mathrm{m}$ | apparent resistivity at the frequency $f=5,403 \mathrm{~Hz}$ |
| KDA_3 | m | apparent depth at the frequency $\mathrm{f}=5,403 \mathrm{~Hz}$ |
| ZST_3 | m | centroid depth at the frequency $\mathrm{f}=5,403 \mathrm{~Hz}$ |
| RHOA_4 | $\Omega \mathrm{m}$ | apparent resistivity at the frequency $\mathrm{f}=8,389 \mathrm{~Hz}$ |
| KDA_4 | m | apparent depth at the frequency $\mathrm{f}=8,389 \mathrm{~Hz}$ |
| ZST_4 | m | centroid depth at the frequency $f=8,389 \mathrm{~Hz}$ |
| RHOA_5 | $\Omega \mathrm{m}$ | apparent resistivity at the frequency $\mathrm{f}=41,430 \mathrm{~Hz}$ |
| KDA_5 | m | apparent depth at the frequency $f=41,430 \mathrm{~Hz}$ |
| ZST_5 | m | centroid depth at the frequency $\mathrm{f}=41,430 \mathrm{~Hz}$ |
| RHOA_6 | $\Omega \mathrm{m}$ | apparent resistivity at the frequency $\mathrm{f}=133,200 \mathrm{~Hz}$ |
| KDA_6 | m | apparent depth at the frequency $\mathrm{f}=133,200 \mathrm{~Hz}$ |
| ZST_6 | m | centroid depth at the frequency $f=133,200 \mathrm{~Hz}$ |

## 4) Inversion models HEM134_INV.XYZ

Example
/Processing by A. Ullmann (BGR) using Oasis montaj
/Levelled data
/ X Y Y $\quad$ Y
//Flight 13401

In this data file all necessary position parameters and inversion models are stored in the order of the following description:

| Channel | Unit | Remarks |
| :---: | :---: | :---: |
| X | m | UTM easting in m (WGS 84, Zone 32N), these coordinates have a false easting of 500000 metres |
| Y | m | UTM northing in m (WGS 84, Zone 32N), these coordinates have no false northing |
| LON | ${ }^{\circ}$ | geographic longitude, reference system WGS 84 |
| LAT | - | geographic latitude, reference system WGS 84 |
| RECORD |  | time mark increasing by 1 every 0.1 seconds |
| UTC_TIME | hhmmss.s | GPS time (UTC) |
| TOPO | m | levelled topographic elevation (in metre above sea level), derived from the difference of bird elevation (BIRN_NN) and bird altitude (H_LASER) |
| H_RADAR | m | smoothed value of the radar altitude minus the effective cable length ( 40 m ) from the helicopter to the bird, corresponds to the bird altitude |
| H_LASER | m | smoothed value of the laser altimeter, corresponds to the bird |
| BIRD_NN | m | smoothed bird elevation (in m asl = metre above sea level), reference system: WGS84 |
| H_BARO | m | filtered value of the barometric sensor minus the effective cable length ( 40 m ) from the helicopter to the bird |
| RHO_I_1 | $\Omega \mathrm{m}$ | resistivity of the top layer of a five-layer inversion model |
| D_I_1 | m | thickness of the top layer of a five-layer inversion model |
| RHO_I_2 | $\Omega \mathrm{m}$ | resistivity of the second layer of a five-layer inversion model |
| D_I_2 | m | thickness of the second layer of a five-layer inversion model |
| RHO_I_3 | $\Omega \mathrm{m}$ | resistivity of the third layer of a five-layer inversion model |
| D_I_3 | m | thickness of the third layer of a five-layer inversion model |
| RHO_I_4 | $\Omega \mathrm{m}$ | resistivity of the fourth layer of a five-layer inversion model |
| D_I_4 | m | thickness of the fourth layer of a five-layer inversion model |
| RHO_I_5 | $\Omega \mathrm{m}$ | resistivity of the fifth layer of a five-layer inversion model |
| QALL | \% | misfit of the inversion (L1 norm) |

The header contains following additional lines:
The header con
/IFREQUENCY
$\begin{array}{lllllll}/ & 1 & 1 & 0 & 1 & 1 & 1\end{array}$
/NUMLAYER
/ 5
/MUELAYER
/ 0
geozentrum hannover Survey Area Vojens, Denmark, 2009

## B) Magnetics

Description of the ASCII coded data file HMG134.XYZ containing the final (levelled) data of a helicopter-borne magnetic (HMG) survey
Example:

| x | Y | LON | LAT | RECORD | UTC_DATE | UTC_TIME | ALT_BIRD | H_RADAR_RAW | H_LASER_RAW | T_BASE_RAW | T_BASE_F | T_RAW | Delta_t | Delta_t_LEV | DELTA_T_FILT |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| //Flight 13401 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| //Date 2009/06/23 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Line 1.1 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 521355 | 612706 | 9.336263 | 55.289514 | 6640 | 20090623 | 80848.0 | 143.3 | 453.5 | 96.1 | 49881.29 | 49881.28 | 50058.42 | 160.60 | 159.62 | -1.33 |
| 521352 | 6127060 | 9.336212 | 55.289513 | 6641 | 20090623 | 80848.1 | 142.7 | 453.4 | 96.1 | 49881.29 | 49881.28 | 50058.34 | 160.58 | 159.59 | -1.32 |
| 521349 | 6127060 | 9.336161 | 55.289511 | 6642 | 20090623 | 80848.2 | 142.1 | 451.4 | 95.0 | 49881.29 | 49881.28 | 50058.27 | 160.55 | 159.56 | -1.31 |

In this data file all necessary position parameters and magnetic data are stored in the order of the following description:

| Channel | Unit | Remarks |
| :---: | :---: | :---: |
| X | m | UTM easting in m (WGS 84, Zone 32N), these coordinates have a false easting of 500000 metres |
| Y | m | UTM northing in m (WGS 84, Zone 32N), these coordinates have no false northing |
| LON | - | geographic longitude, reference system WGS 84 |
| LAT | - | geographic latitude, reference system WGS 84 |
| RECORD |  | time mark increasing by 1 every 0.1 seconds |
| UTC_DATE | yyyymmdd | date |
| UTC_TIME | hhmmss.s | GPS time (UTC) |
| ALT_BIRD | m | smoothed bird elevation (in m asl = metre above sea level), reference system: WGS84 |
| H_RADAR | m | smoothed value of the radar altitude minus the effective cable length ( 40 m ) from the helicopter to the bird, corresponds to the bird altitude |
| H_LASER | m | smoothed value of the laser altimeter, corresponds to the bird |
| T_BASE_RAW | nT | raw data of the magnetic field at the base station |
| T_BASE_F | nT | processed data of the magnetic field at the base station |
| T_RAW | nT | raw data of the magnetic field at the bird |
| DELTA_T | nT | anomalies of the magnetic field |
| DELTA_T_LEV | nT | levelled anomalies of the magnetic field |
| DELTA_T_FILT | nT | (high-pass) filtered anomalies of the magnetic field, corrected for man-made effects |

## Remarks:

Lines starting with "/" comment,
Lines starting with "//"
Lines starting with "Line"
Lines starting with "Tie"
flight number and date
lines,
tie lines.

## C) Radiometry

Description of the ASCII coded data file HRD134.XYZ containing the final (levelled) data of a helicopter-borne radiometric (HRD) survey
Example:


In this data file all necessary position parameters and radiometric data are stored in the order of the following description:

| Channel | Unit | Remarks |
| :---: | :---: | :---: |
| X | m | UTM easting in m (WGS 84, Zone 32N), these coordinates have a false easting of 500000 metres |
| Y | m | UTM northing in m (WGS 84, Zone 32N), these coordinates have no false northing |
| LON | 。 | geographic longitude, reference system WGS 84 |
| LAT | - | geographic latitude, reference system WGS 84 |
| RECORD |  | time mark increasing by 1 every 0.1 seconds |
| UTC_DATE | yyyymmdd | date |
| UTC_TIME | hhmmss.s | GPS time (UTC) |
| ALT_BIRD | m | smoothed bird elevation (in m asl = metre above sea level), reference system: WGS84 |
| H_RADAR_RAW | m | value of the radar altitude minus the effective cable length ( 40 m ) from the helicopter to the bird, corresponds to the bird altitude |
| H_LASER_RAW | m | value of the laser altimeter, corresponds to the bird |
| HAG | m | altitude of helicopter above ground level |
| PRESSURE | kPa | air pressure |
| TEMP | ${ }^{\circ} \mathrm{C}$ | air temperature |
| LIVE_T | ms | live time |
| COSMIC | cps | cosmic radiation > 3 MeV |
| TOT_RAW | cps | measured total count rate |
| POT_RAW | cps | measured potassium count rate |
| URA_RAW | cps | measured uranium count rate |
| THO_RAW | cps | measured thorium count rate |
| URAUP | cps | measured uranium count rate in upward looking crystal |


| TOT | cps | total count |
| :---: | :---: | :---: |
| POT | \% | potassium concentration on ground level |
| URA | ppm | equivalent uranium concentration ground level |
| THO | ppm | equivalent thorium concentration ground level |
| TOT_LEV | cps | levelled total count, corrected for the effect of vegetation |
| POT_LEV | \% | levelled potassium concentration on ground level, corrected for the effect of vegetation |
| URA_LEV | ppm | levelled equivalent uranium concentration on ground level, corrected for the effect of vegetation |
| THO_LEV | ppm | levelled equivalent thorium concentration on ground level, corrected for the effect of vegetation |
| EXPO | $\mu \mathrm{R} / \mathrm{h}$ | ground level exposure rate |
| Remarks: |  |  |
| Lines starting with "/" |  | mment, |
| Lines starting with "//" |  | ght number and date, |
| Lines starting with "Line" |  | es, |
| Lines starting with "Tie" |  |  |

## Appendix III

## DVD



Adobe - Adobe Reader herunterladen.URL
$\backslash$ Acrobat Reader\Linux\}
AdbeRdr9.3.4-1_i486linux_deu.bin
$\backslash$ Acrobat Reader $\backslash \mathrm{Mac} \backslash$
AdbeRdr930_de_DE_i386.pkg.zip
$\backslash$ Acrobat Reader $\backslash$ Windows
AdbeRdr934_de_DE.exe
\Data
\Data HEM
Format_description_HEM134.txt
HEM134_APP.XYZ
HEM134_DAT.XYZ
HEM134_INV.xyz
HEM134_RAW.XYZ
\Data HMG
Format_description_HMG134.txt
HMG134.XYZ
\Data HRD
Format_description_HRD134.txt HRD134.XYZ
$\backslash$ Maps
$\backslash M a p s \backslash A r c G i s$
134 Vojens apparent resistivity rhoa1.mxd 134 Vojens apparent resistivity rhoa2.mxd 134 Vojens apparent resistivity rhoa3.mxd 134 Vojens apparent resistivity rhoa4.mxd 134 Vojens apparent resistivity rhoa5.mxd 134 Vojens apparent resistivity rhoa6.mxd 134 Vojens centroid depth zst1.mxd 134 Vojens centroid depth zst2.mxd 134 Vojens centroid depth zst3.mxd 134 Vojens centroid depth zst4.mxd 134 Vojens centroid depth zst5.mxd 134 Vojens centroid depth zst6.mxd 134 Vojens DEM.mxd 134 Vojens exposure rate.mxd 134 Vojens flight lines.mxd 134 Vojens magnetic anomalies filtered.mxd 134 Vojens magnetic anomalies.mxd 134 Vojens Potassium.mxd 134 Vojens resistivity -001m.mxd 134 Vojens resistivity -003m.mxd 134 Vojens resistivity -005m.mxd 134 Vojens resistivity -010m.mxd 134 Vojens resistivity -020m.mxd 134 Vojens resistivity -030m.mxd 134 Vojens resistivity -050m.mxd 134 Vojens resistivity -070m.mxd 134 Vojens resistivity -090m.mxd 134 Vojens resistivity -120m.mxd

134 Vojens Thorium.mxd
134 Vojens total count.mxd
134 Vojens Uranium.mxd
$\backslash$ Maps $\backslash$ ArcGis $\backslash$ Legends
134 Vojens Legende apparent resistivity rhoa1.bmp
134 Vojens Legende apparent resistivity rhoa2.bmp
134 Vojens Legende apparent resistivity rhoa3.bmp
134 Vojens Legende apparent resistivity rhoa4.bmp
134 Vojens Legende apparent resistivity rhoa5.bmp
134 Vojens Legende apparent resistivity rhoa6.bmp
134 Vojens Legende centroid depth zst1.bmp
134 Vojens Legende centroid depth zst2.bmp
134 Vojens Legende centroid depth zst3.bmp
134 Vojens Legende centroid depth zst4.bmp
134 Vojens Legende centroid depth zst5.bmp
134 Vojens Legende centroid depth zst6.bmp
134 Vojens Legende DEM.bmp
134 Vojens Legende Exposure rate.bmp
134 Vojens Legende flight lines.bmp
134 Vojens Legende magnetic anomalies filtered.bmp
134 Vojens Legende magnetic anomalies.bmp
134 Vojens Legende Potassium.bmp
134 Vojens Legende resistivity -001m.bmp
134 Vojens Legende resistivity -003m.bmp
134 Vojens Legende resistivity -005m.bmp
134 Vojens Legende resistivity -010m.bmp
134 Vojens Legende resistivity -020m.bmp
134 Vojens Legende resistivity -030m.bmp
134 Vojens Legende resistivity -030m.tif
134 Vojens Legende resistivity -050m.bmp
134 Vojens Legende resistivity -070m.bmp
134 Vojens Legende resistivity -090m.bmp
134 Vojens Legende resistivity -120m.bmp
134 Vojens Legende Thorium.bmp
134 Vojens Legende Total count.bmp
134 Vojens Legende Uranium.bmp
$\backslash$ Maps $\backslash$ ArcGis $\backslash$ Oasis
EXPO.grd
expo.map
expo.map.xml
flightlines_P.map
flightlines_P.map.xml
hsp_P1.map
hsp_P1.map.xml
hsp_P2.map
hsp_P2.map.xml
hsp_P3.map
hsp_P3.map.xml
hsp_P4.map
hsp_P4.map.xml
hsp_P5.map
hsp_P5.map.xml
hsp_P6.map
hsp_P6.map.xml
mag.map
mag.map.xml
mag_filt.map
mag_filt.map.xml
mag_filt_P.map
mag_filt_P.map.xml
MAG_LEV.grd
MAG_LEV_CE_BP_M2_S.grd
mag_P.map
mag_P.map.xml
pot.map
pot.map.xml
POT_BIO.grd
rho_001m.grd
rho_001m.map
rho_001m.map.xml
rho_003m.grd
rho_003m.map
rho_003m.map.xml
rho_005m.grd
rho_005m.map
rho_005m.map.xml
rho_010m.grd
rho_010m.map
rho_010m.map.xml
rho_020m.grd
rho_020m.map
rho_020m.map.xml
rho_030m.grd
rho_030m.map
rho_030m.map.xml
rho_050m.grd
rho_050m.map
rho_050m.map.xml
rho_070m.grd
rho_070m.map
rho_070m.map.xml
rho_090m.grd
rho_090m.map
rho_090m.map.xml
rho_090m_P.map
rho_090m_P.map.xml
rho_120m.grd
rho_120m.map
rho_120m.map.xml
rho_120m_P.map
rho_120m_P.map.xml
rho_P.map
rho_P.map.xml
rhoa1.map
rhoa1.map.xml
rhoa1_LEV_IQ.GRD
rhoa2.map
rhoa2.map.xml
rhoa2_LEV_IQ.GRD
rhoa3.map
rhoa3.map.xml
rhoa3_LEV_IQ.GRD
rhoa4.map
rhoa4.map.xml
rhoa4_LEV_IQ.GRD
rhoa5.map
rhoa5.map.xml
rhoa5_LEV_IQ.GRD
rhoa6.map
rhoa6.map.xml
rhoa6_LEV_IQ.GRD
tho.map
tho.map.xml
THO_BIO.grd
tot.map
tot.map.xml
TOT_BIO.grd
ura.map
ura.map.xml
URA_BIO.grd
zst1.map
zst1.map.xml
zst1_LEV_IQ.GRD
zst2.map
zst2.map.xml
zst2_LEV_IQ.GRD
zst3.map
zst3.map.xml
zst3_LEV_IQ.GRD
zst4.map
zst4.map.xml
zst4_LEV_IQ.GRD
zst5.map
zst5.map.xml
zst5_LEV_IQ.GRD
zst6.map
zst6.map.xml
zst6_LEV_IQ.GRD
$\backslash$ Maps $\backslash$ ArcGis $\backslash$ Topography
1212_IV_NØ.tfw
1212_IV_NØ.tif
1212_IV_NV.tfw
1212_IV_NV.tif
1212_IV_SØ.tfw
1212_IV_S $\varnothing$.tif
1212_IV_SV.tfw
1212_IV_SV.tif
134 Vojens TK Schnitt - Kopie.tfw
134 Vojens TK Schnitt - Kopie.tif
134 Vojens TK Schnitt.tfw
134 Vojens TK Schnitt.tif
$\backslash \mathrm{Maps} \backslash \mathrm{DEM}$
134 Vojens DEM.pdf
\Maps $\backslash$ Flightlines
134 Vojens flight lines.pdf
$\backslash$ Maps $\backslash \mathrm{HEM}$
134 Vojens apparent resistivity rhoa1.pdf
134 Vojens apparent resistivity rhoa2.pdf
134 Vojens apparent resistivity rhoa3.pdf
134 Vojens apparent resistivity rhoa4.pdf
134 Vojens apparent resistivity rhoa5.pdf
134 Vojens apparent resistivity rhoa6.pdf
134 Vojens centroid depth zst1.pdf
134 Vojens centroid depth zst2.pdf
134 Vojens centroid depth zst3.pdf
134 Vojens centroid depth zst4.pdf
134 Vojens centroid depth zst5.pdf
134 Vojens centroid depth zst6.pdf
134 Vojens resistivity -001m.pdf
134 Vojens resistivity -003m.pdf

134 Vojens resistivity -005m.pdf 134 Vojens resistivity -010m.pdf 134 Vojens resistivity -020m.pdf 134 Vojens resistivity -030m.pdf 134 Vojens resistivity -050m.pdf 134 Vojens resistivity -070m.pdf 134 Vojens resistivity -090m.pdf 134 Vojens resistivity -120m.pdf
$\backslash$ Maps $\backslash H M G$
134 Vojens magnetic anomalies filtered.pdf
134 Vojens magnetic anomalies.pdf
$\backslash$ Maps $\backslash H R D$
134 Vojens exposure rate.pdf
134 Vojens Potassium.pdf
134 Vojens Thorium.pdf
134 Vojens total count.pdf
134 Vojens Uranium.pdf
$\backslash$ Report
Technical Report 134 Vojens.pdf
\VRS
VRS 1340011.PDF
VRS 1340019.PDF
VRS 1340021.PDF
VRS 1340029.PDF
VRS 1340031.PDF
VRS 1340039.PDF
VRS 1340041.PDF
VRS 1340049.PDF
VRS 1340051.PDF
VRS 1340059.PDF
VRS 1340061.PDF
VRS 1340069.PDF
VRS 1340071.PDF
VRS 1340079.PDF
VRS 1340081.PDF
VRS 1340089.PDF
VRS 1340091.PDF
VRS 1340099.PDF
VRS 1340101.PDF
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## Appendix IV

## Maps

(reduced to a scale of 1:75,000)


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# Alle anderen Karten und Vertikalsektionen sind in dieser Web-Fassung des Berichtes nicht enthalten. 

All other maps and vertical resistivity sections are not included in this web edition of the report.

