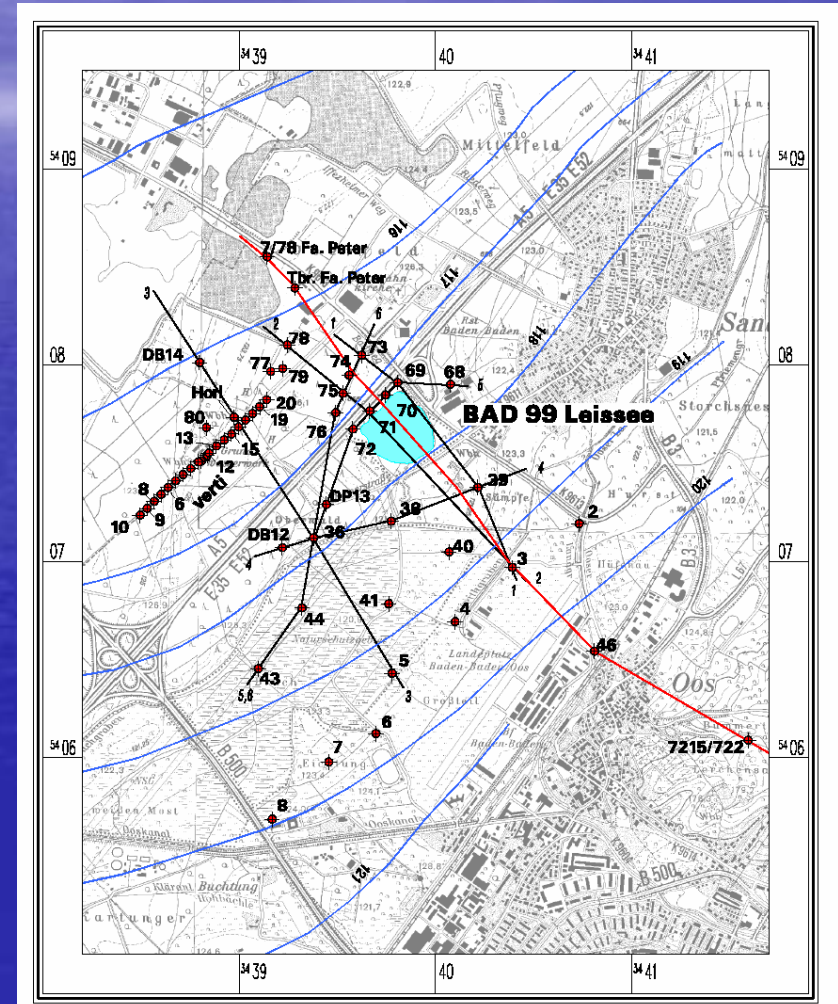


Mathematical modeling of environmental isotope data to estimate the catchment area of drinking water supply

Case study Lake Leis
(Rhine valley)

W. Stichler & P. Maloszewski (GSF)
B. Bertleff & R. Watzel (LGRB)

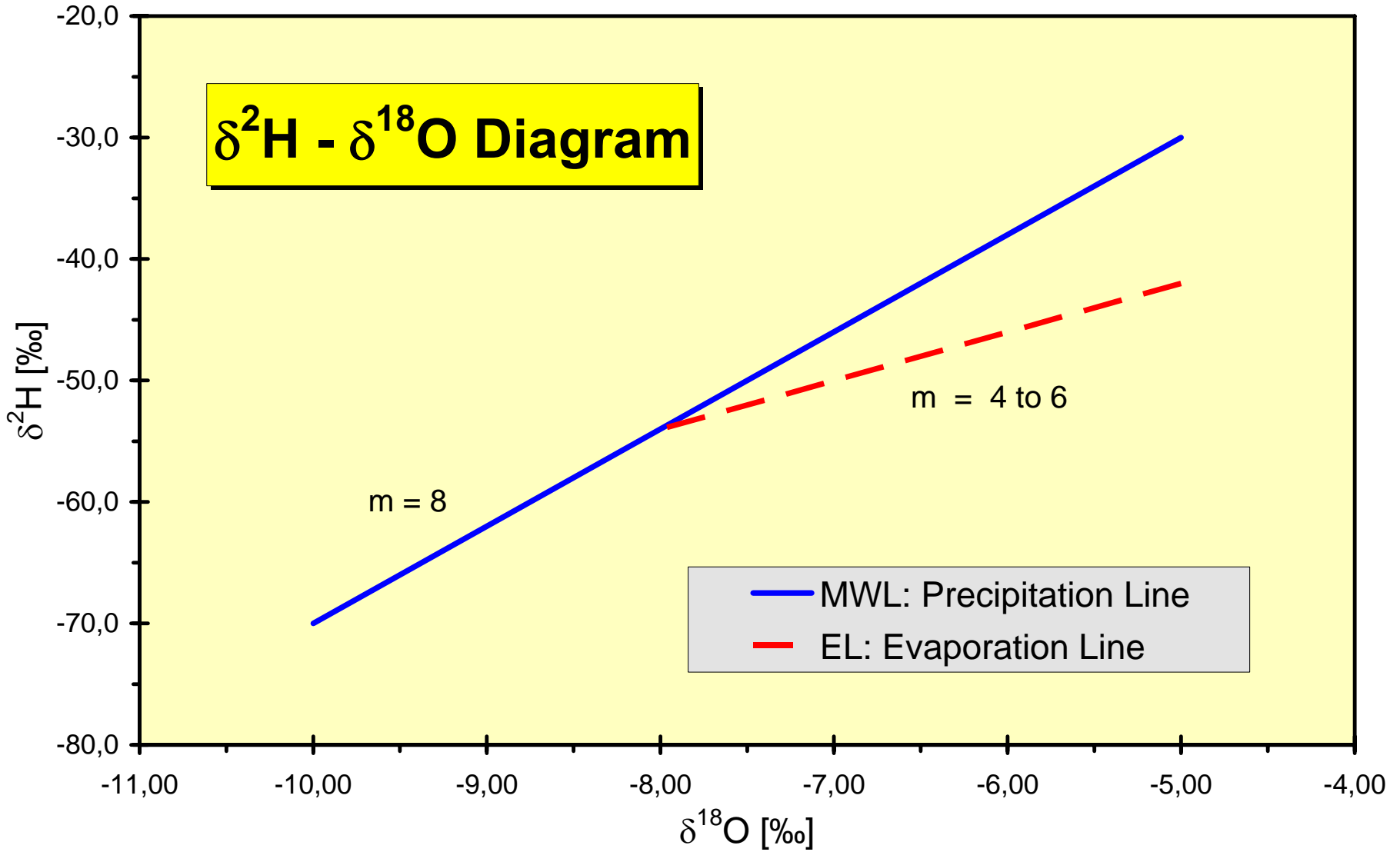
Interaction surface water – ground water







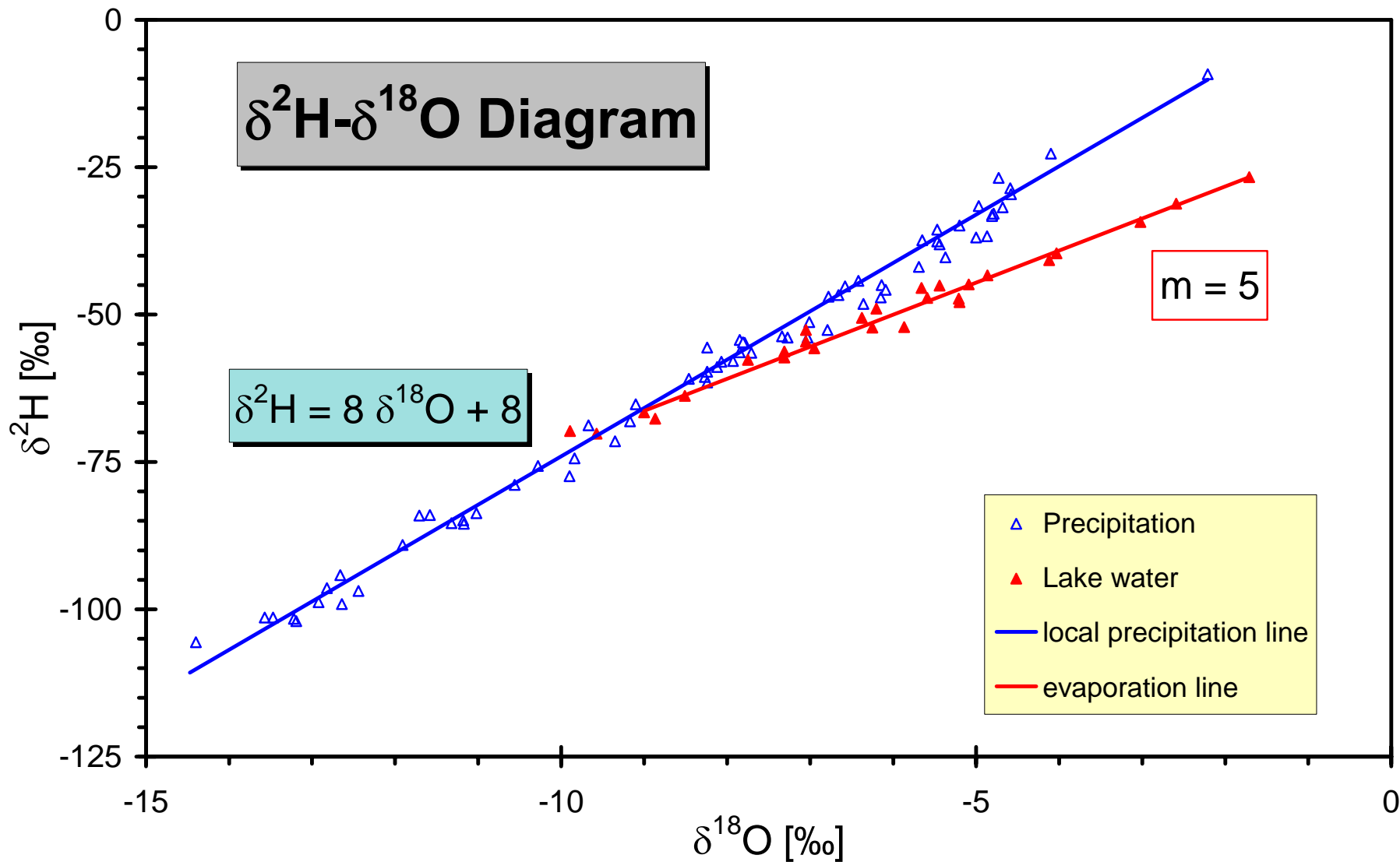
$\delta^2\text{H} - \delta^{18}\text{O}$ Diagram



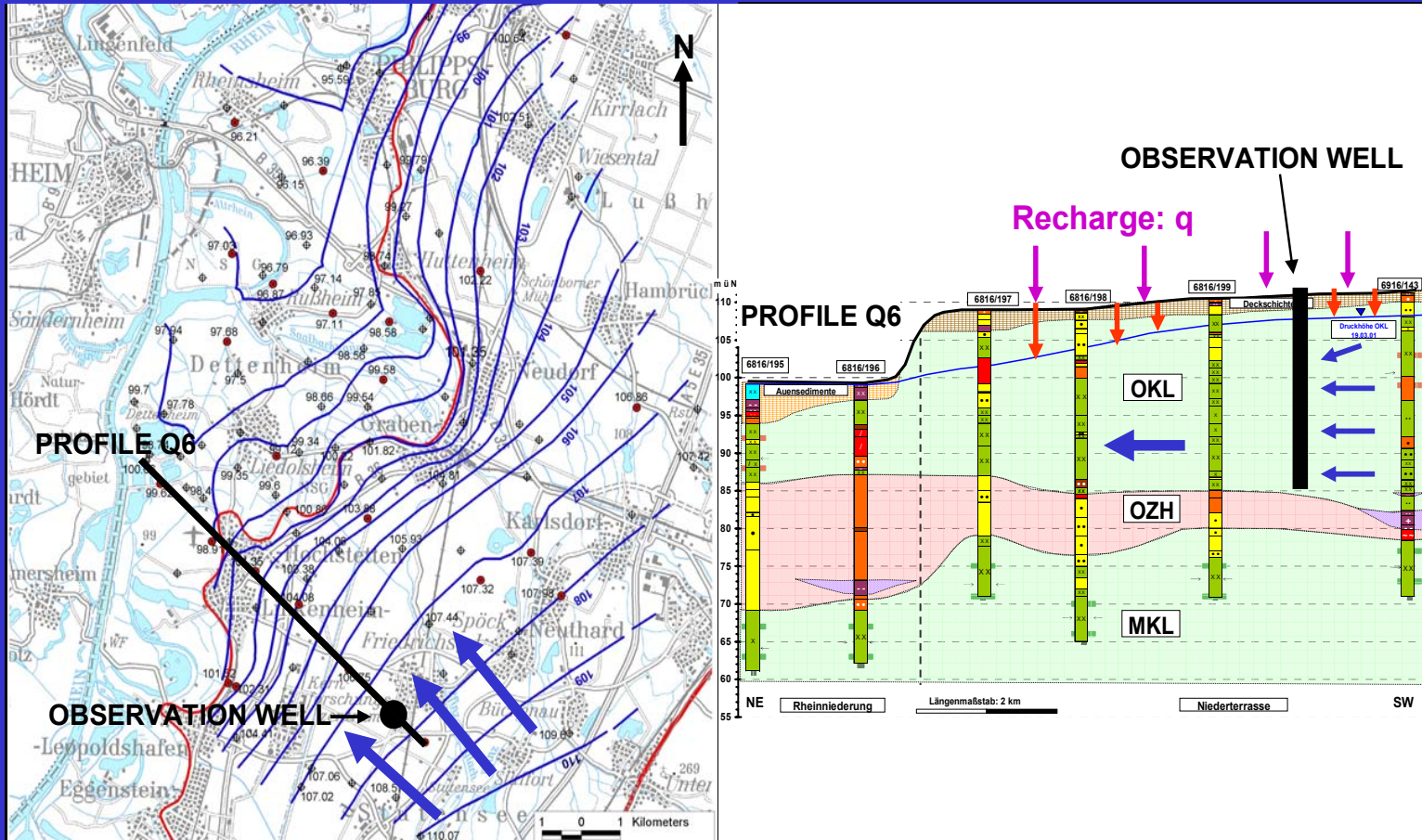
$\delta^2\text{H}-\delta^{18}\text{O}$ Diagram

$$\delta^2\text{H} = 8 \delta^{18}\text{O} + 8$$

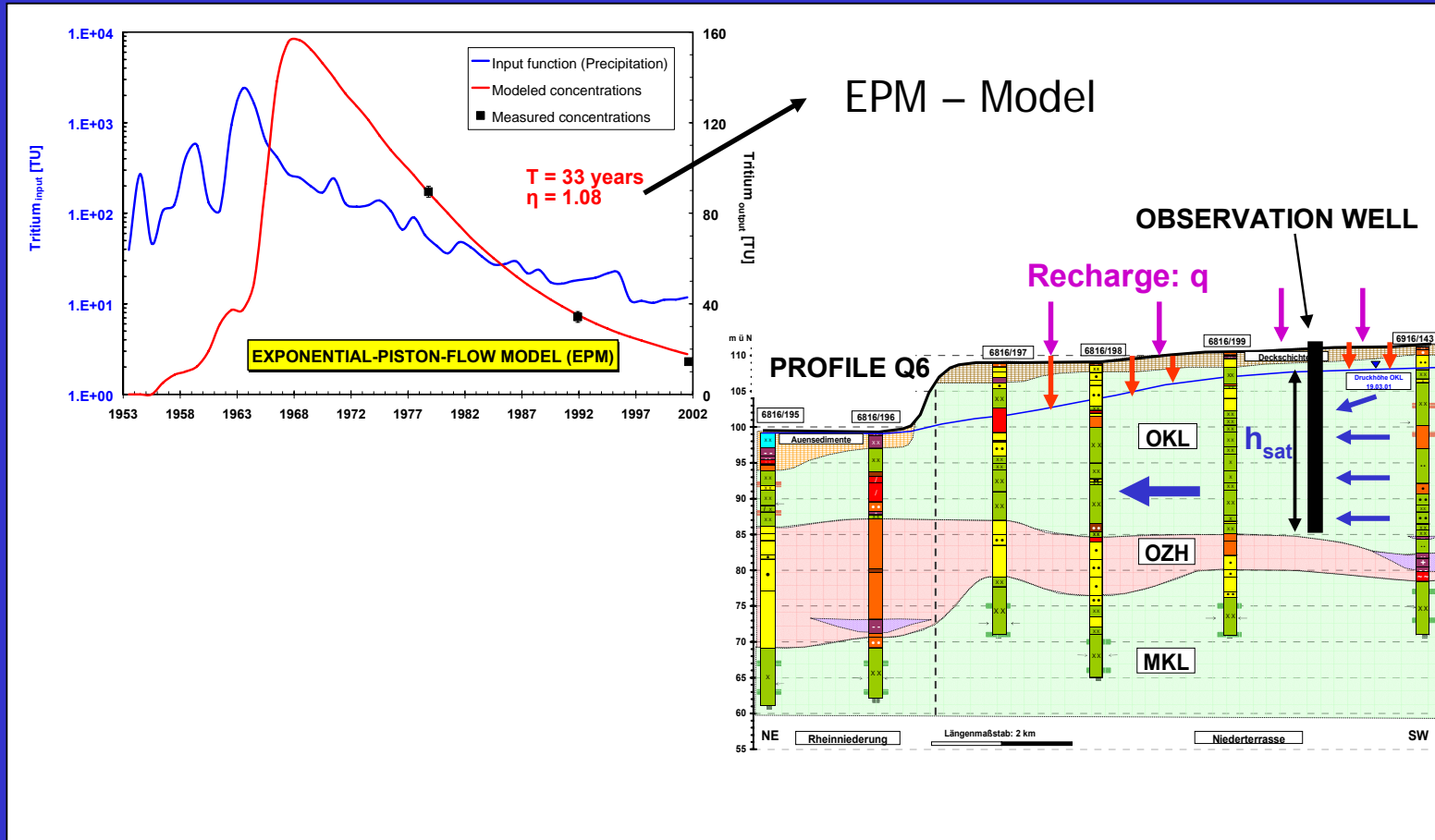
$m = 5$



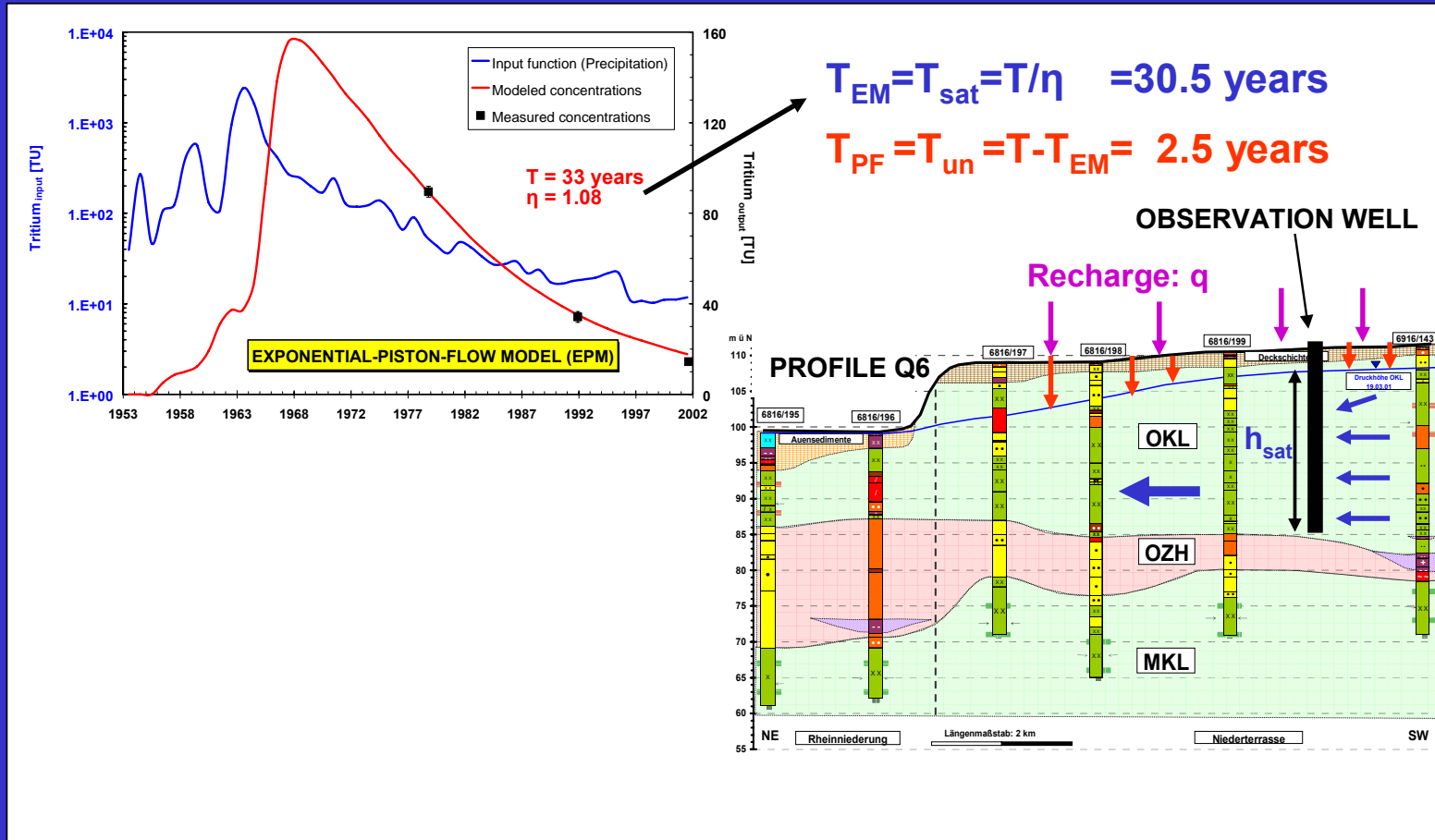
Estimation of local recharge



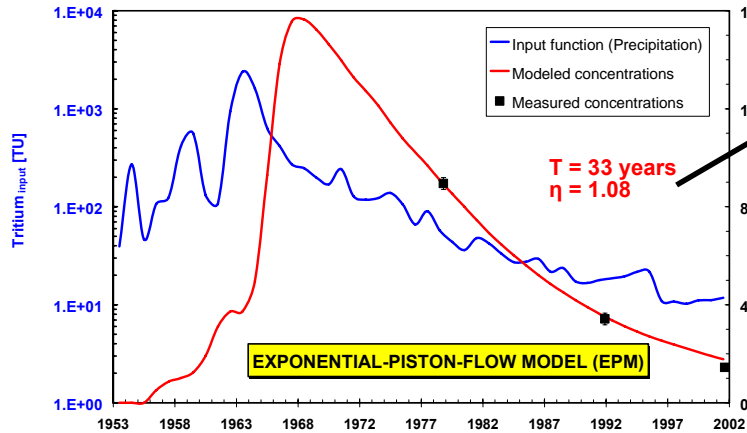
Estimation of local recharge



Estimation of local recharge



Estimation of local recharge

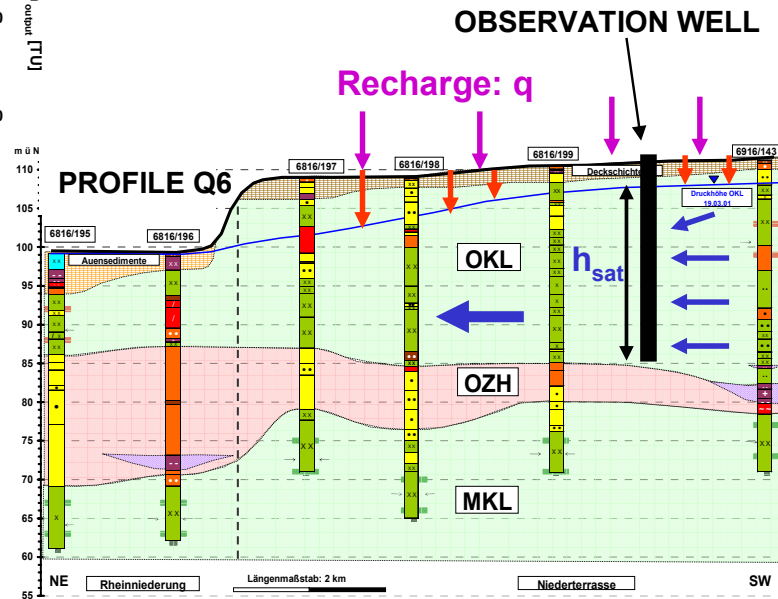


$$T_{EM} = T_{sat} = T / \eta = 30.5 \text{ years}$$

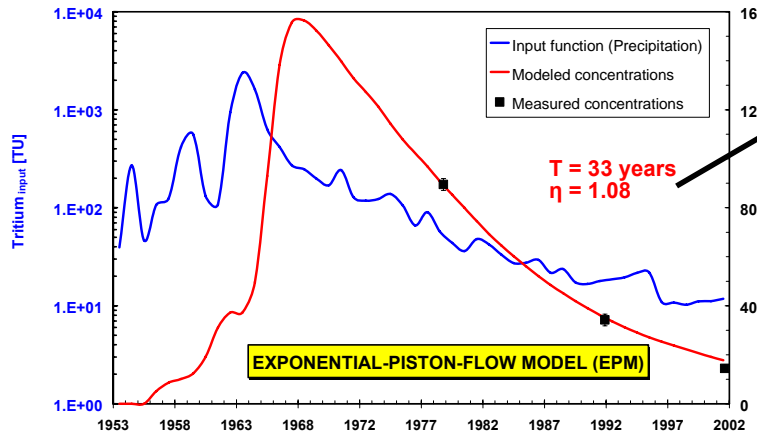
$$T_{PF} = T_{un} = T - T_{EM} = 2.5 \text{ years}$$

(1) knowing $\Theta = 0.08$, $h_{un} = 3.2 \text{ m}$

$$q = h_{un} \cdot \Theta / T_{PF} = 105 \text{ mm/a}$$



Estimation of local recharge



$$T_{EM} = T_{sat} = T / \eta = 30.5 \text{ years}$$

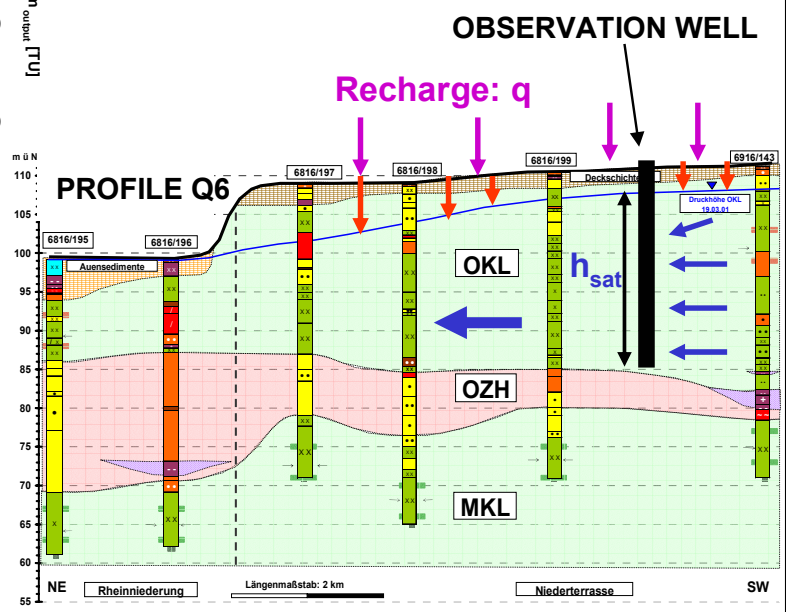
$$T_{PF} = T_{un} = T - T_{EM} = 2.5 \text{ years}$$

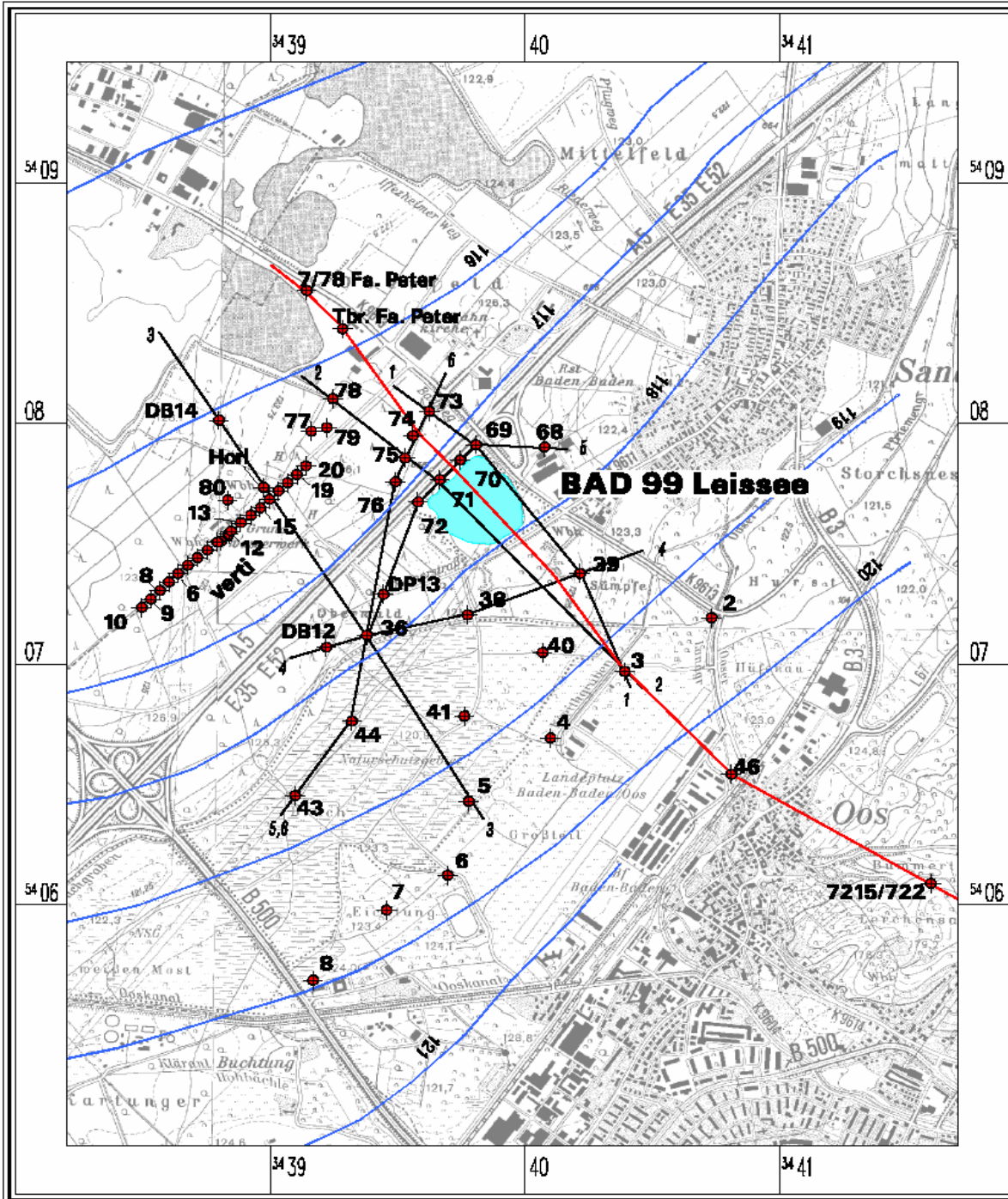
(1) knowing $\Theta=0.08$, $h_{un}=3.2 \text{ m}$

$$q = h_{un} \cdot \Theta / T_{PF} = 105 \text{ mm/a}$$

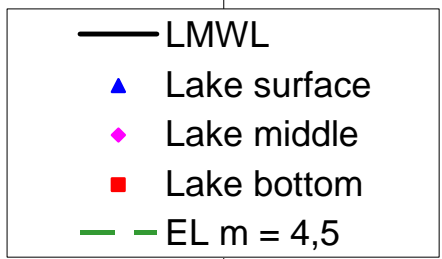
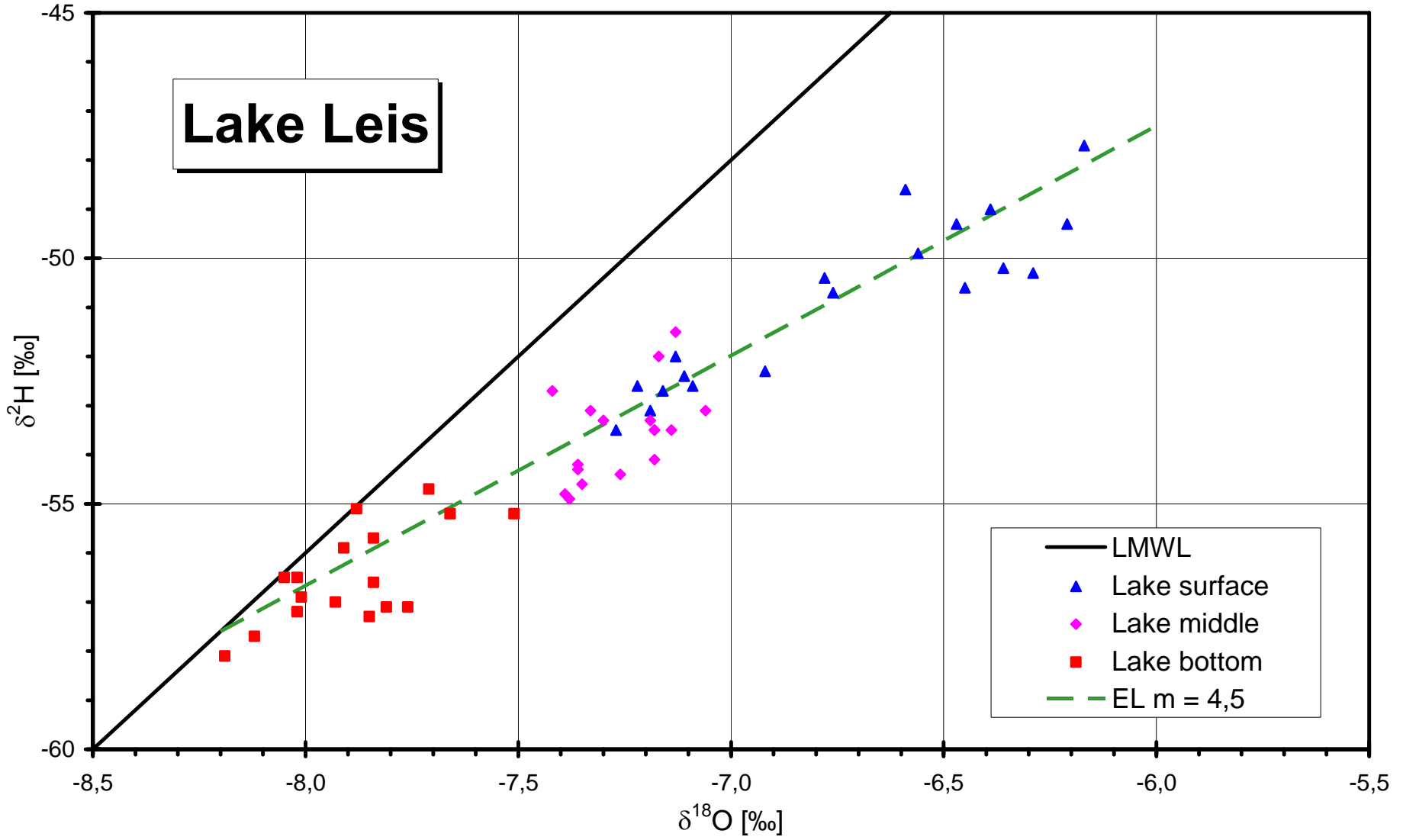
(2) estimating $n=0.2$

$$h_{sat} = q \cdot T_{EM} / n = 16 \text{ m}$$

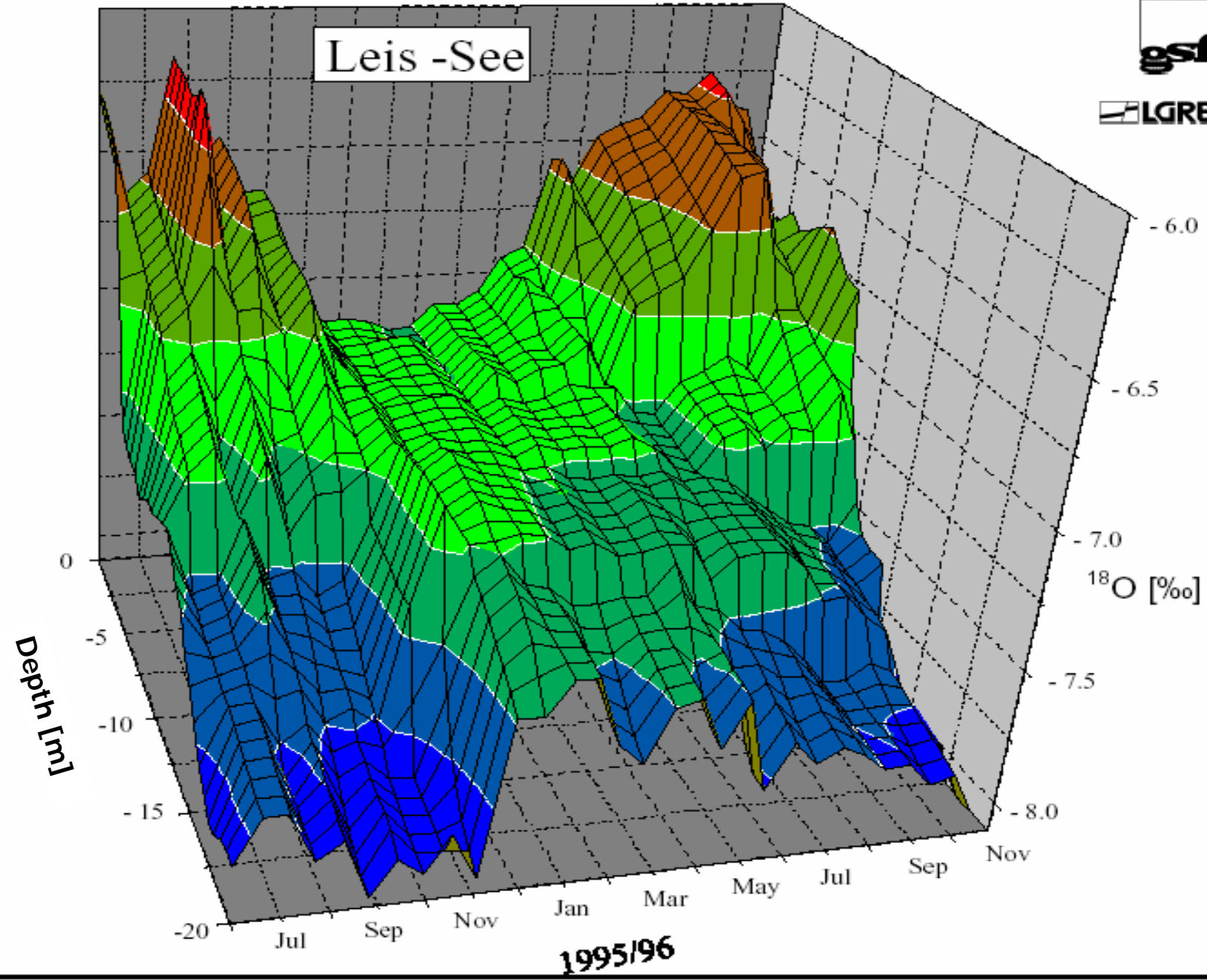




Lake Leis

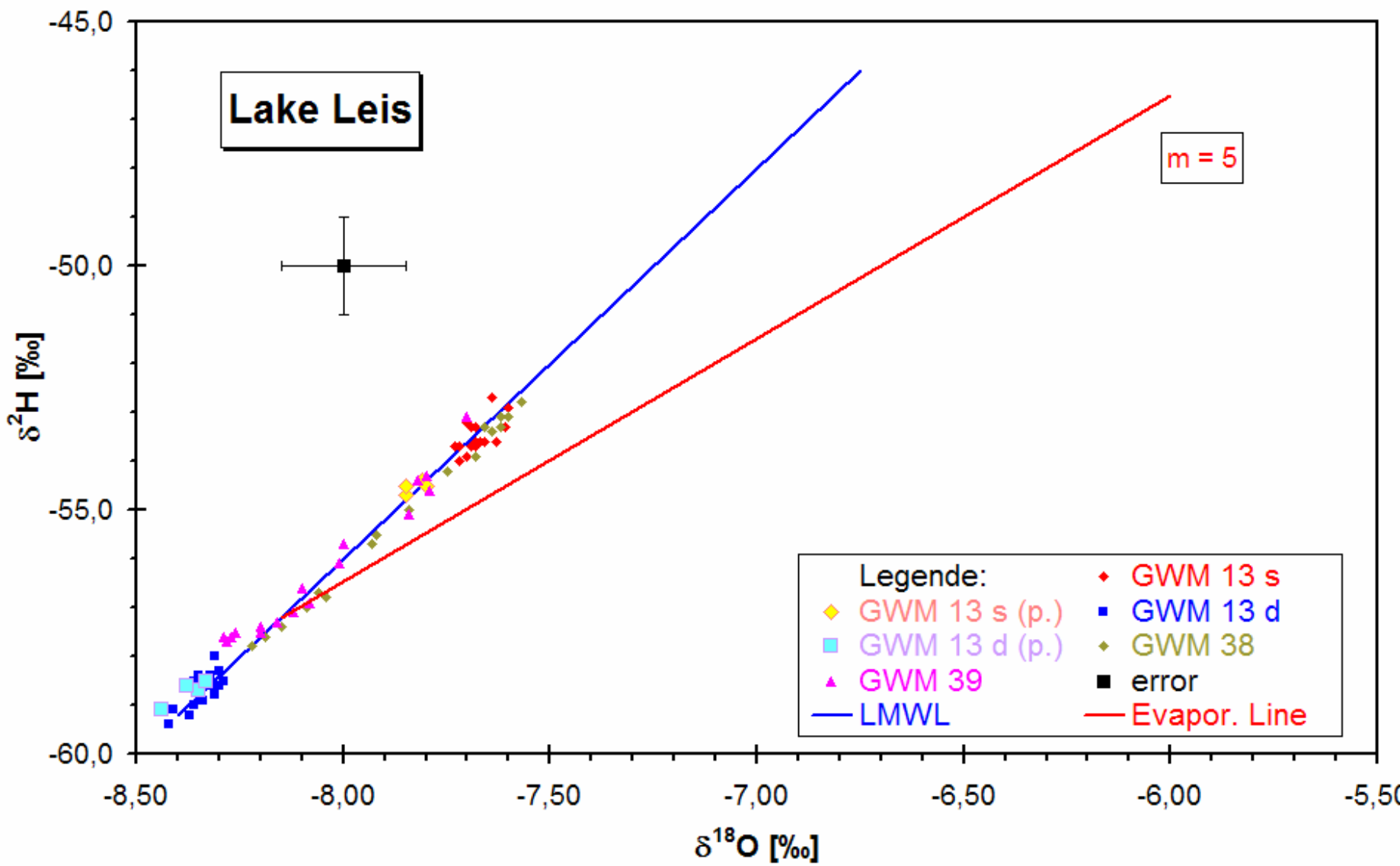


Leis -See

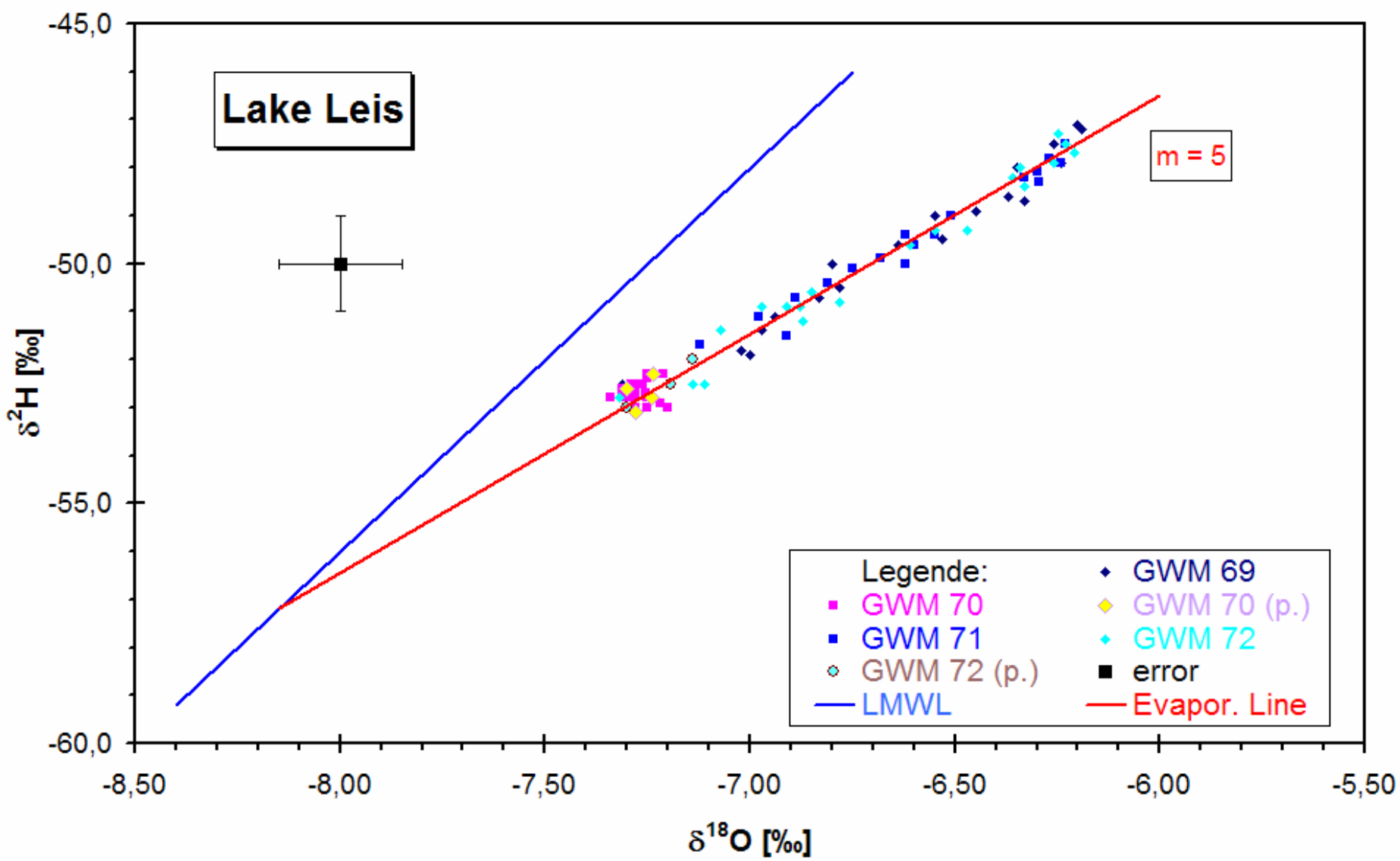


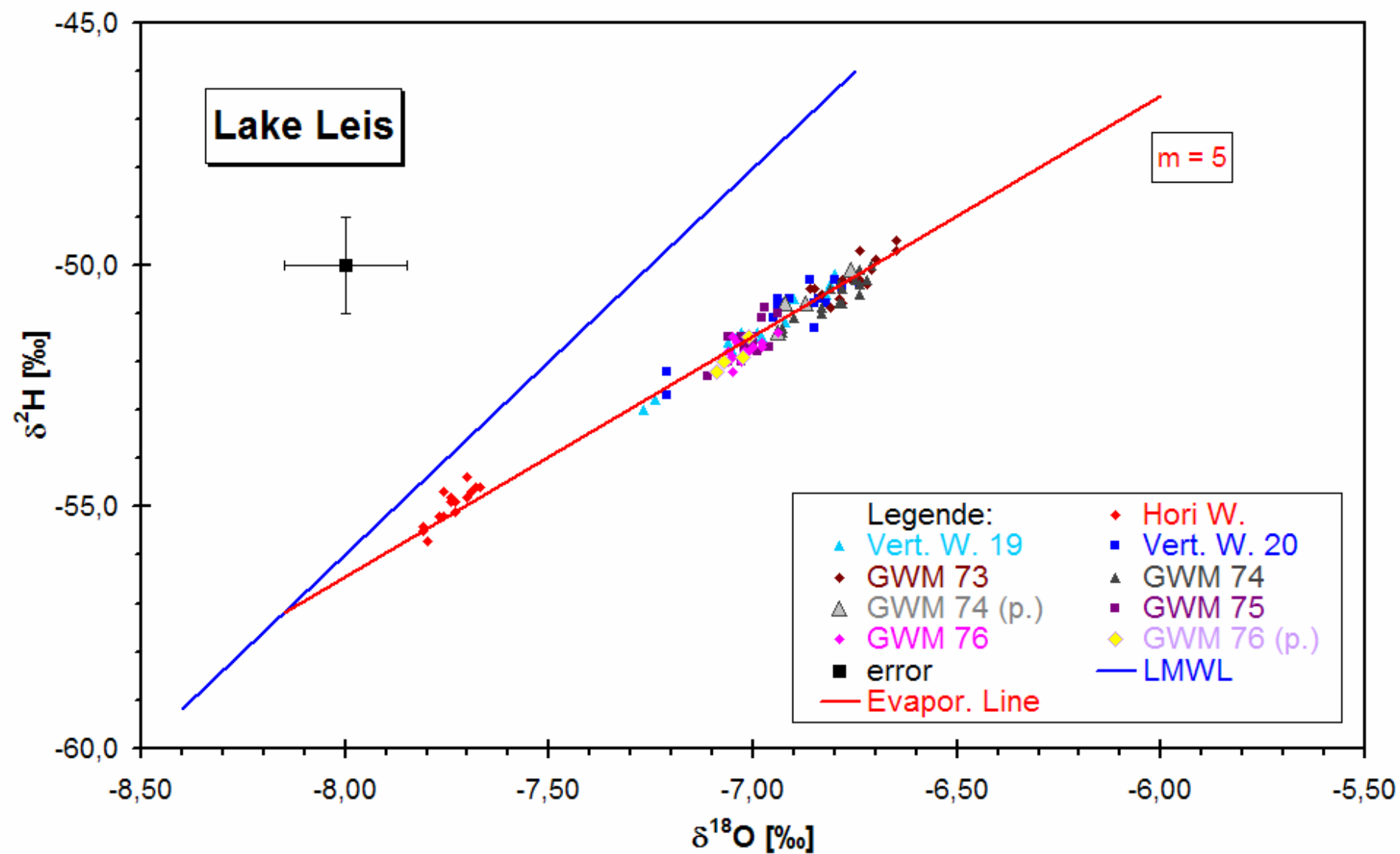
Lake Leis

m = 5

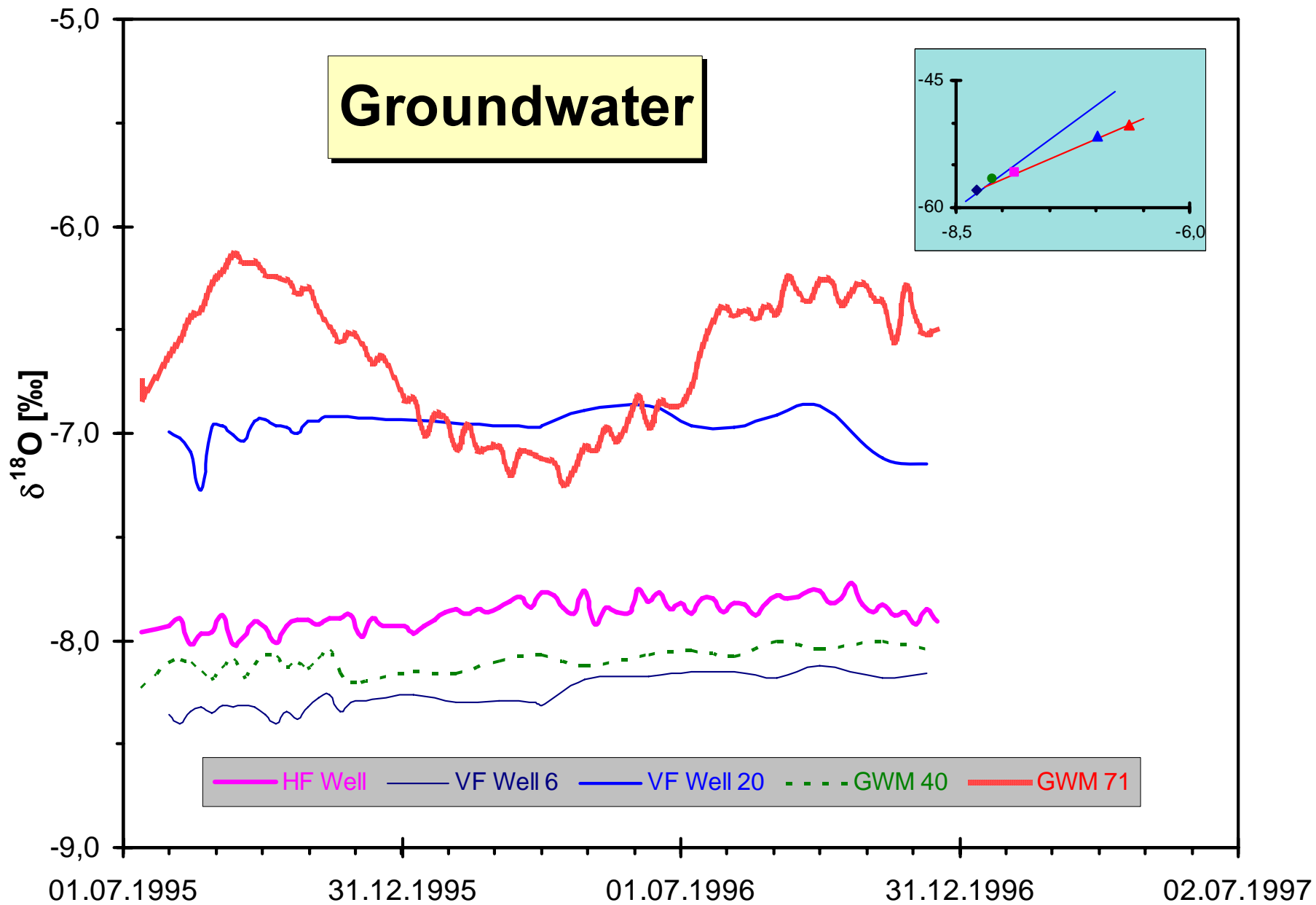


Lake Leis





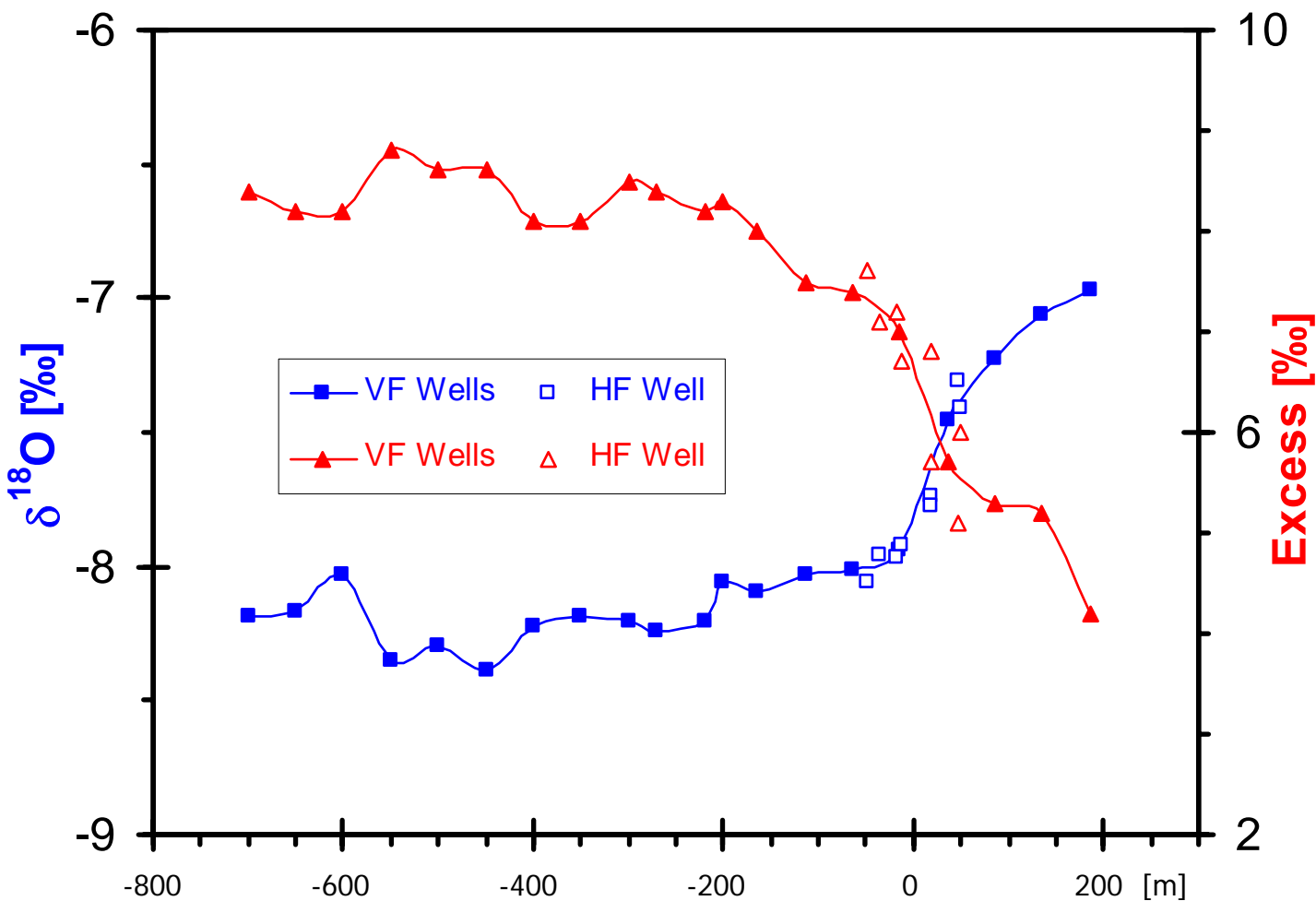
Groundwater



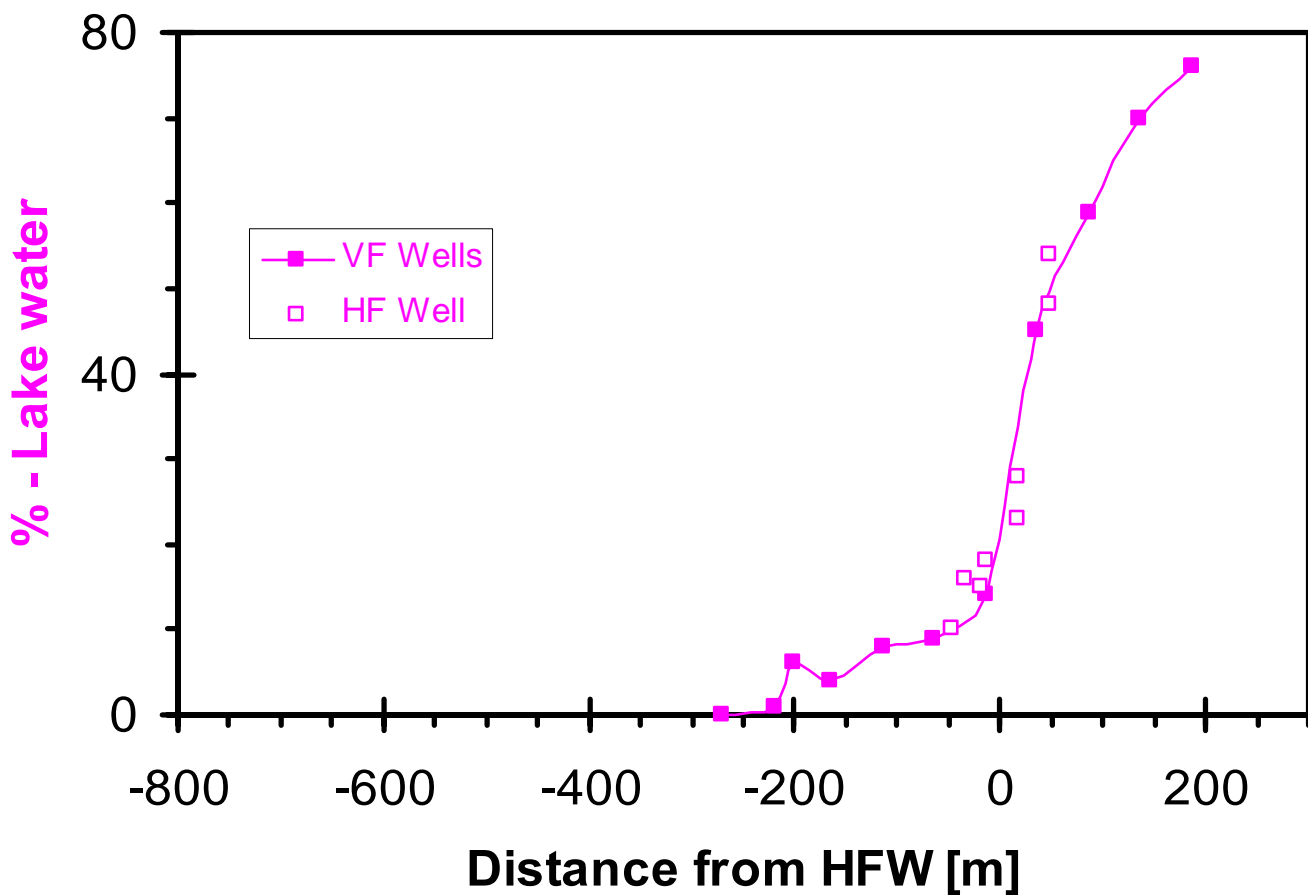
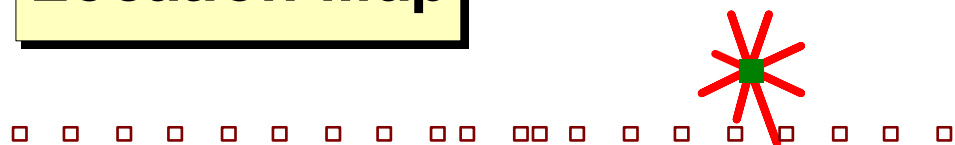
Location Map

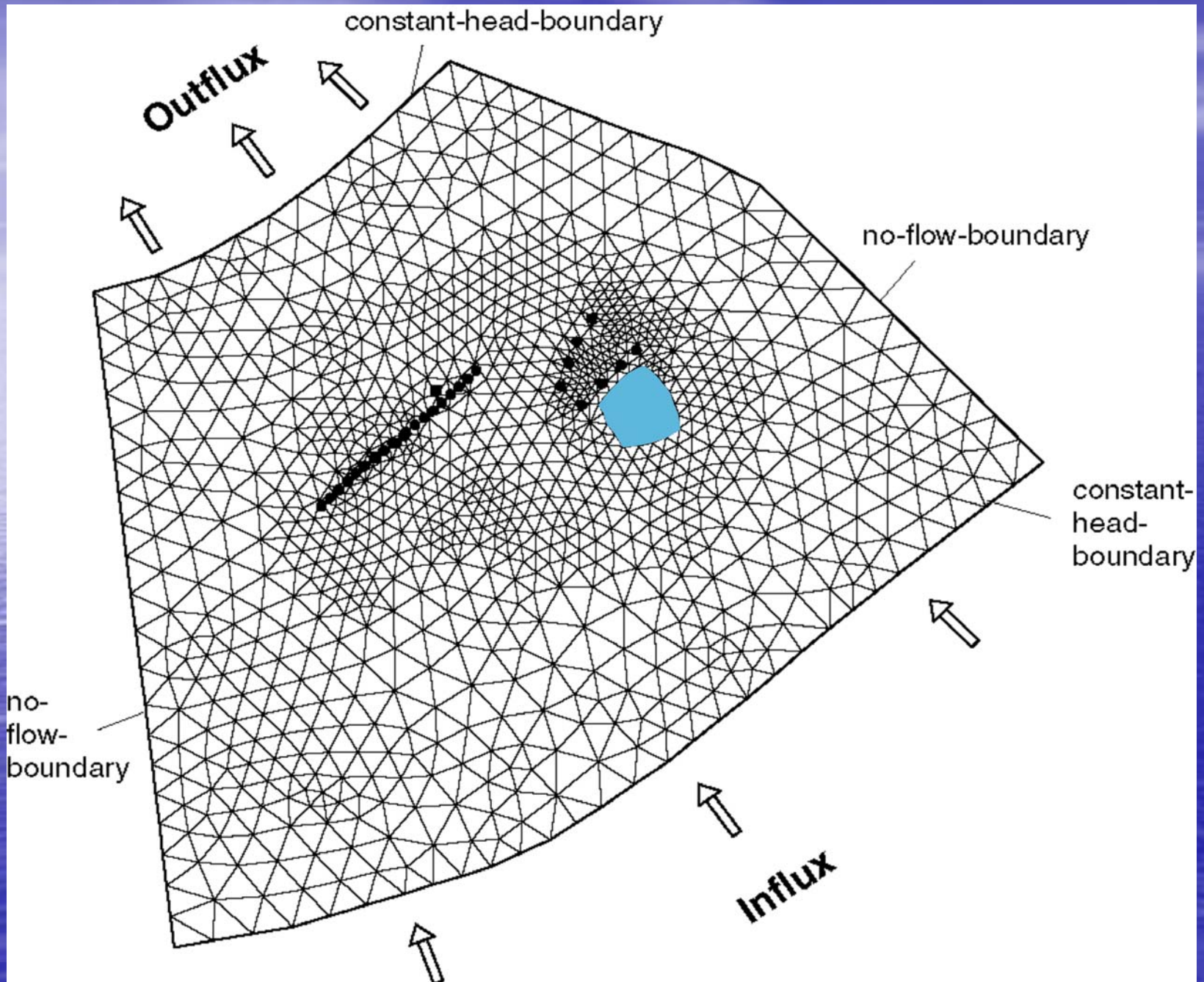


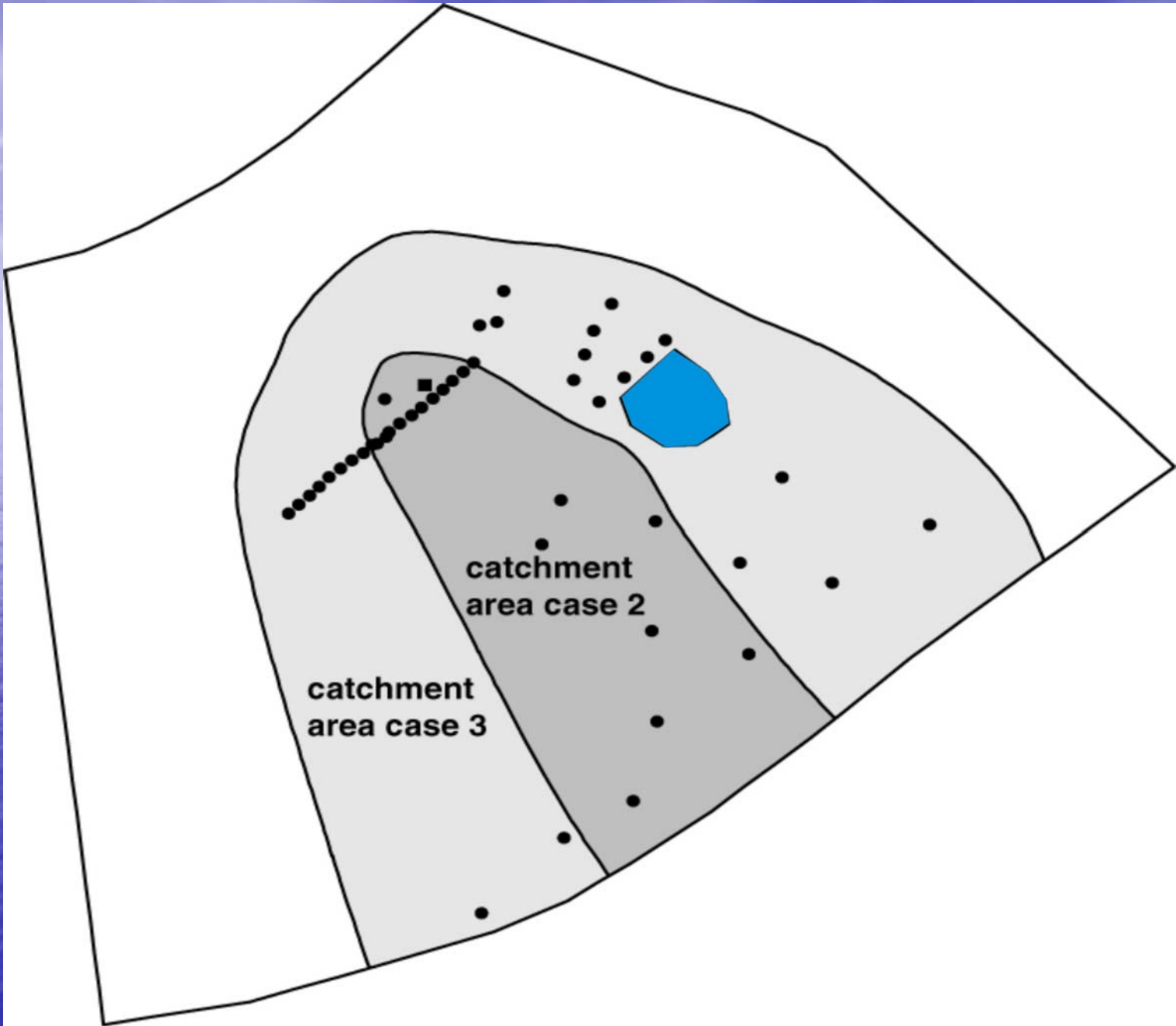
□ VF Wells ■ HF Well — Single Tubes

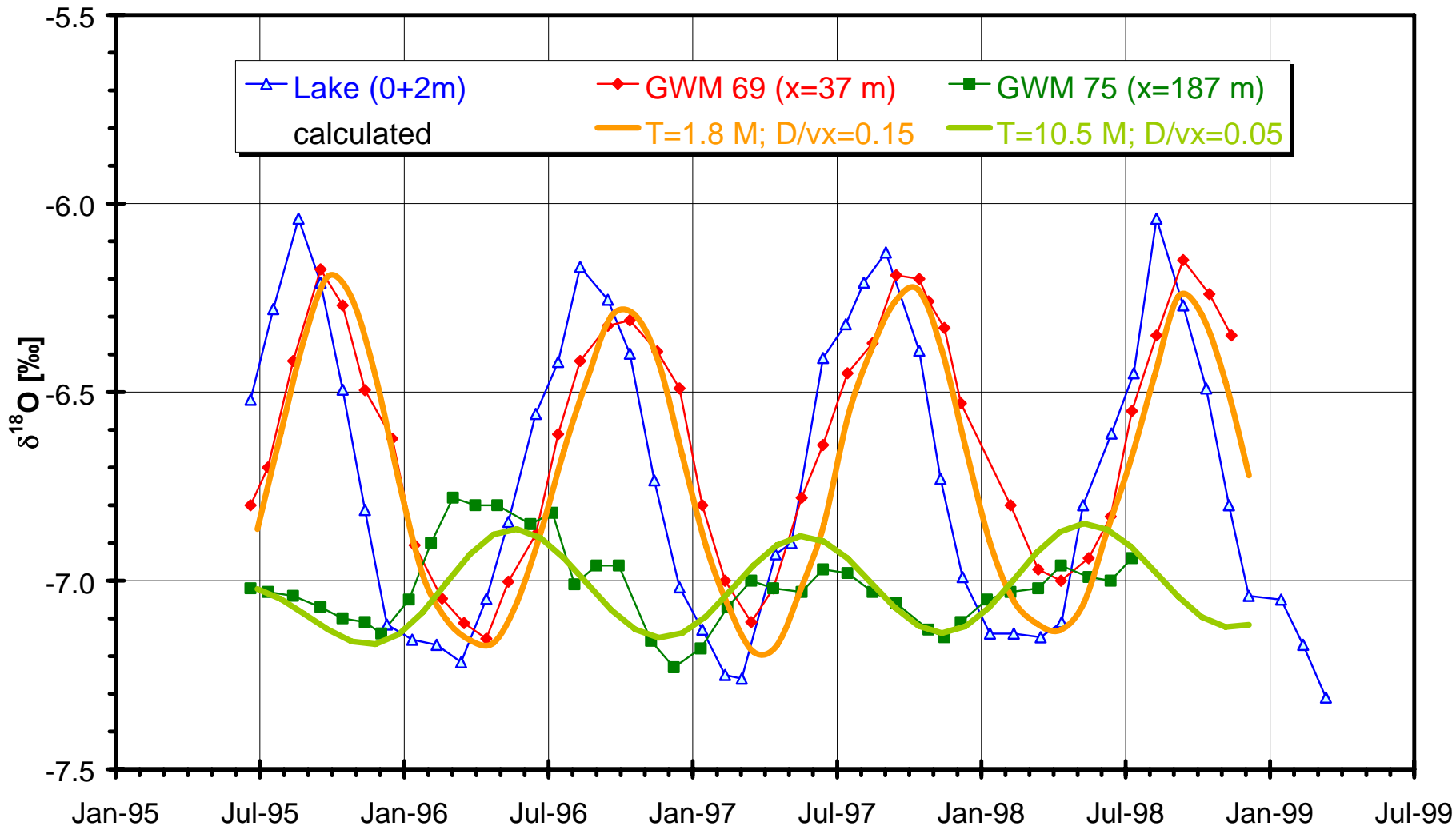


Location Map







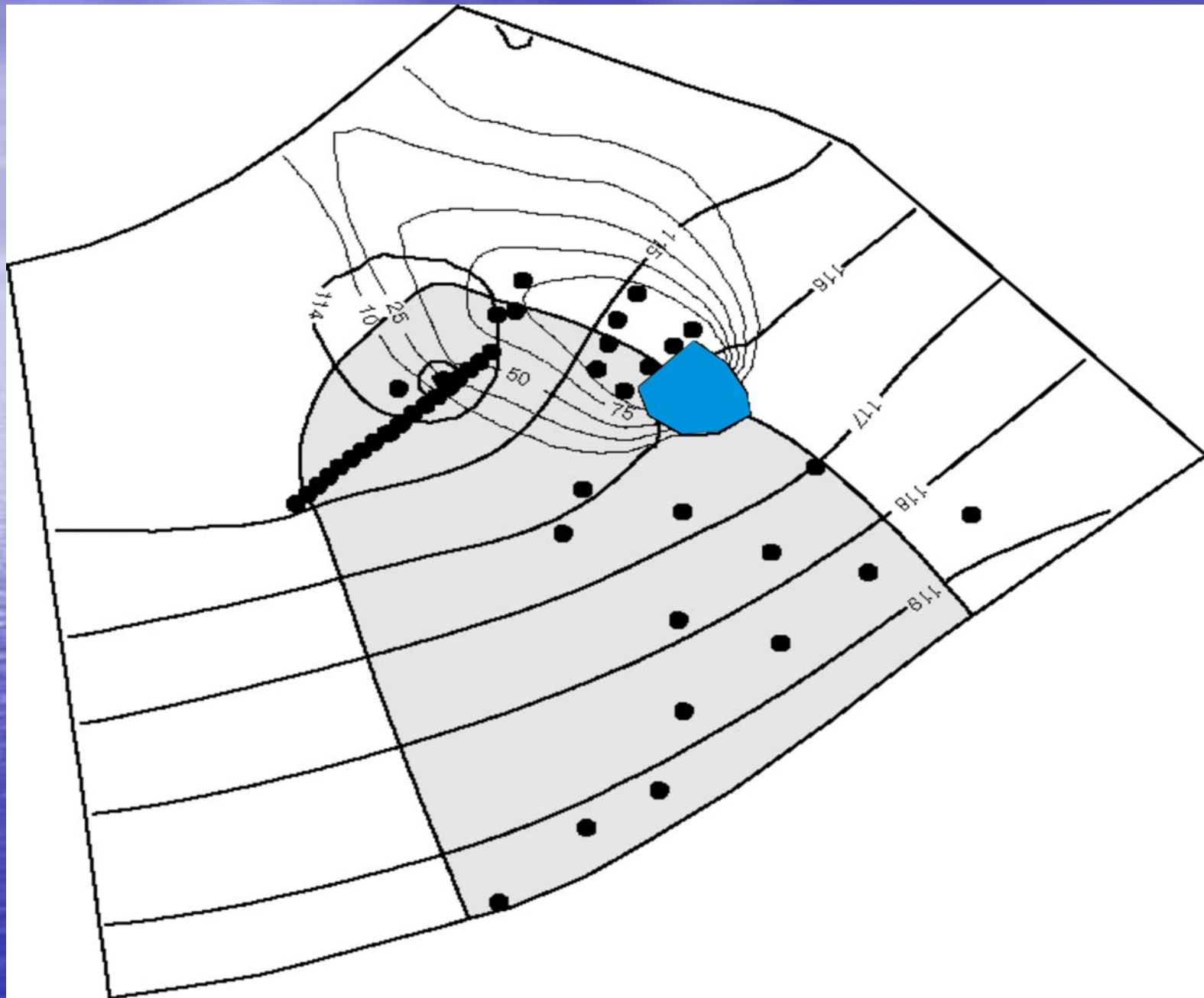


MEAN $\delta^{18}\text{O}$ -VALUES OF WATER SAMPLES AND VALUES OF MODEL PARAMETERS OBTAINED BY CALIBRATION OF A TWO-DIMENSIONAL ANALYTICAL DISPERSION MODEL USING THE ISOTOPE DATA

Sampling point	Distance from the lake	Mean isotope content	Portion of lake water	Mean transit time of water	Dispersion parameter
-	x (m)	$\delta^{18}\text{O}$ (‰)	p (%)	T (d)	P_D (-)
Ground water	-	- 8.22	0	-	-
Lake water	0	- 6.59	100	0	0
P 69	37	- 6.63	98	37 ± 3	0.15 ± 0.02
P 70	46	- 7.22	60	n.d.	n.d.
P 71	34	- 6.63	98	42 ± 4	0.18 ± 0.02
P 72	47	- 6.57	100	70 ± 4	0.20 ± 0.02
P 73	258	- 6.84	90	330 ± 30	0.05 ± 0.01
P 74	237	- 6.73	99	360 ± 30	0.10 ± 0.03
P 75	187	- 7.00	85	300 ± 30	0.05 ± 0.01
P 76	141	- 7.08	80	300 ± 30	0.10 ± 0.03

MEAN VALUES OF HYDRAULIC PARAMETERS CALCULATED FOR THE FLOW DISTANCE BETWEEN LAKE BANK AND OBSERVATION WELLS

Sampling point	Distance from the lake	Mean ground water velocity	Longitudinal dispersivity	Mean hydraulic conductivity
-	x (m)	v (m/d)	α_L (m)	k (m/s)
P 69	37	1.00 ± 0.10	5.5 ± 0.5	$7.2 \pm 0.7 \cdot 10^{-4}$
P 71	34	0.81 ± 0.08	6.1 ± 0.6	$5.9 \pm 0.6 \cdot 10^{-4}$
P 72	47	0.67 ± 0.05	9.4 ± 0.9	$4.8 \pm 0.5 \cdot 10^{-4}$
P 73	258	0.78 ± 0.08	12.9 ± 2.5	$5.7 \pm 0.5 \cdot 10^{-4}$
P 74	237	0.66 ± 0.06	23.7 ± 7.7	$4.8 \pm 0.4 \cdot 10^{-4}$
P 75	187	0.62 ± 0.10	9.4 ± 1.8	$4.5 \pm 0.7 \cdot 10^{-4}$
P 76	141	0.67 ± 0.07	14.1 ± 4.2	$4.8 \pm 0.5 \cdot 10^{-4}$



Conclusions:

- The results of isotope studies facilitate the description of lake - ground water interaction in a qualitative and quantitative way.
- The quantification of lake water portions in the surrounding aquifer is also basis for the evaluation of possible changes in ground water quality.
- The use of isotope data reduces parameter uncertainty in numerical ground water modelling on a local scale.
- The combination of isotope and modelling techniques provides the unique possibility for the delineation of a ground water protection area.

Thank you for your attention