





# Magnetotelluric measurements to explore deeper structures of the Tendaho geothermal prospect (Afar, NE Ethiopia)

by

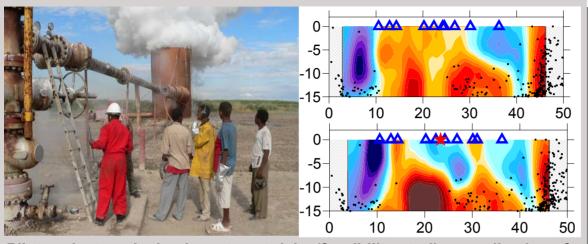
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**November 2008** 



### **Outline**

- A Introduction
- B The MT method
- C MT results from Tendaho



Pilot projects reducing investment risks (feasibility studies, application of geoscientific methods for site evaluation)



Policy advice and awareness raising

Goal:
Partner countries
use their
geothermal
potential









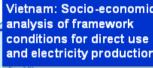
Uganda: Investigation of Buranga geothermal site



Eritrea: Training of geo -thermal experts in cooperation with the United Nations University



Yemen: Feasibility Study at Al Lisi as part of a GEF Project





Rwanda: geothermal studies training



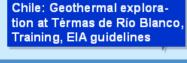
Tanzania: Geothermal exploration, training, and awareness raising of decision makers



Kenya: Training in GIS, airborne thermal camera survey



Ethiopia: geophysical exploration at Tendaho, Support to the African Geothermal Conference 2006



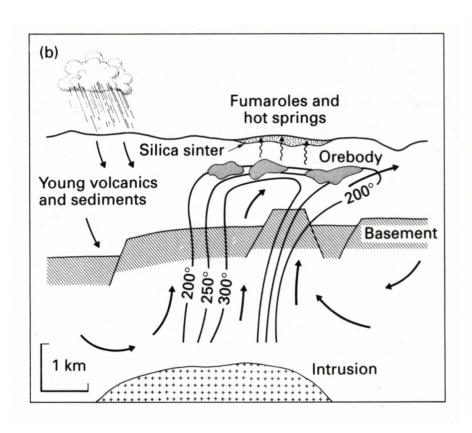


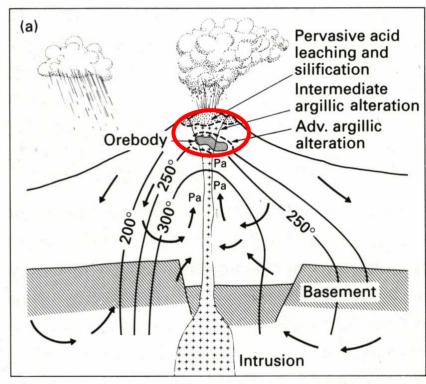
#### Main phases of surface exploration

- Literature survey (desk top study)
- Regional review (with prioritizing sites)
- Site selection (apply for concession of most promising site)
- Remote sensing (satellite images, aerial photographs, IR, InSAR)
- Geologic survey (rock units, tectonic setting, active faulting, age of youngest volcanic activity, surface manifestations, alteration zones)
- Hydrologic survey (meteorological data, discharge rates of springs, water table, hydraulic gradient, mean residence time)
- Geochemical survey (chemical and isotopic composition of fluids and gases, geo-thermometry, soil gas survey)
- Interim conceptual model (for geophysical survey planning)
- Geophysical survey (temperature gradient, resistivity methods, gravity, magnetics, micro-seismicity)
- Synthesis (final conceptual model with suggestion of sites for exploratory drilling)



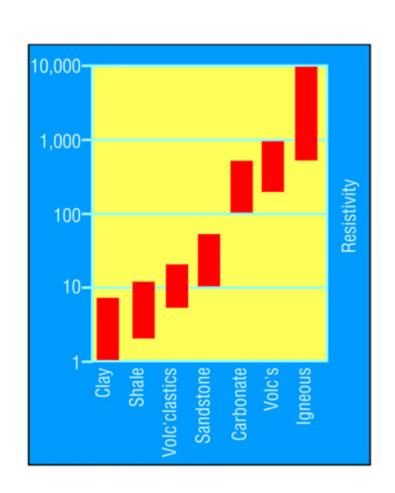
### Schematical geothermal reservoir types

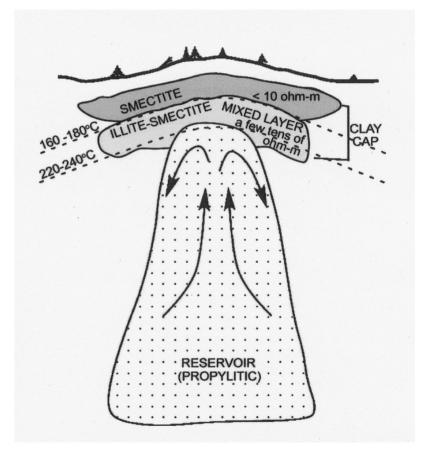




from: Evans 1997

Working model of a geothermal reservoir, produced by hydrothermal alteration





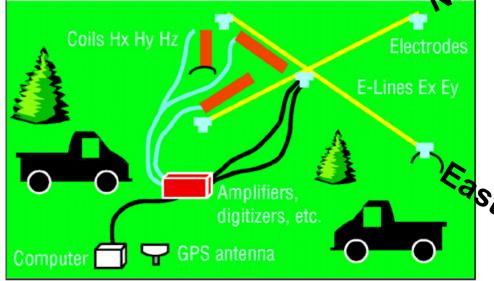
Modified after: Johnston et al. 1992

#### B The MT method



Two 5-channel stations (Ex, Ey, Bx, By, Bz) Time synchronised recording 10 kHz bis 0,01 Hz (100 s)

Sensors: Induction coil magnetometers
Pb-PbCl electrodes, dipole length 100 m



#### B The MT method

No transmitter: passive method. What are the sources?



- Variations of the Earth's magnetic field cover a broad frequency range
- Lower frequencies penetrate deeper into the subsurface and thus allow conclusions on electrical conductivity structures at depth



#### B The MT method

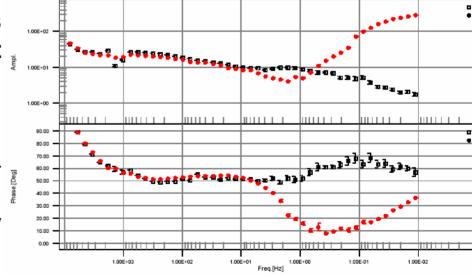
Exploration depth depends on frequency and conductivity of subsurface:

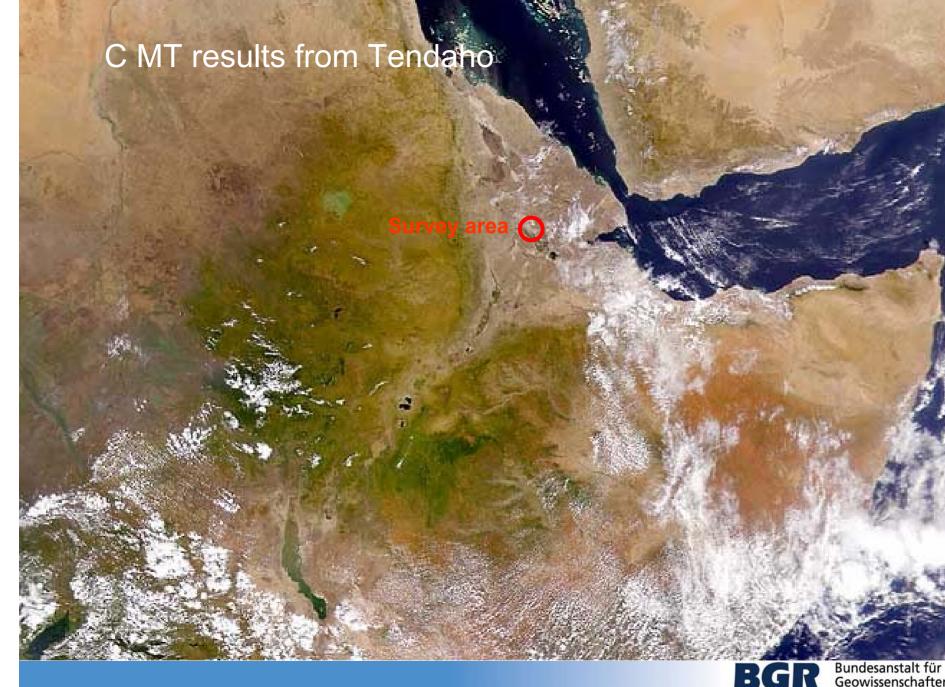
e.g.: 10 Ohm\*m / 0.1 Hz 5 km depth of exploration

$$\delta \cong 0.5 \sqrt{\frac{\rho}{f}}$$

The ratio of E- to H- field variations is used to calculate **apparent resistivities** for selected frequencies.

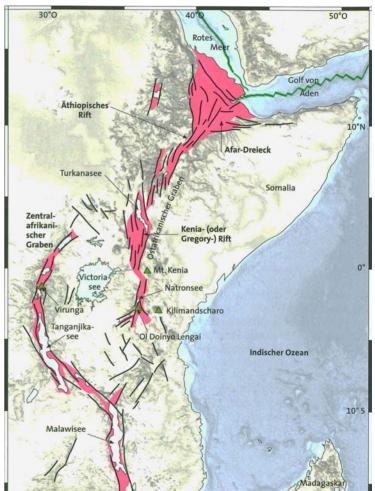
The **phase** shift between E- and H- fields is also indicative for changes in subsurface resistivity





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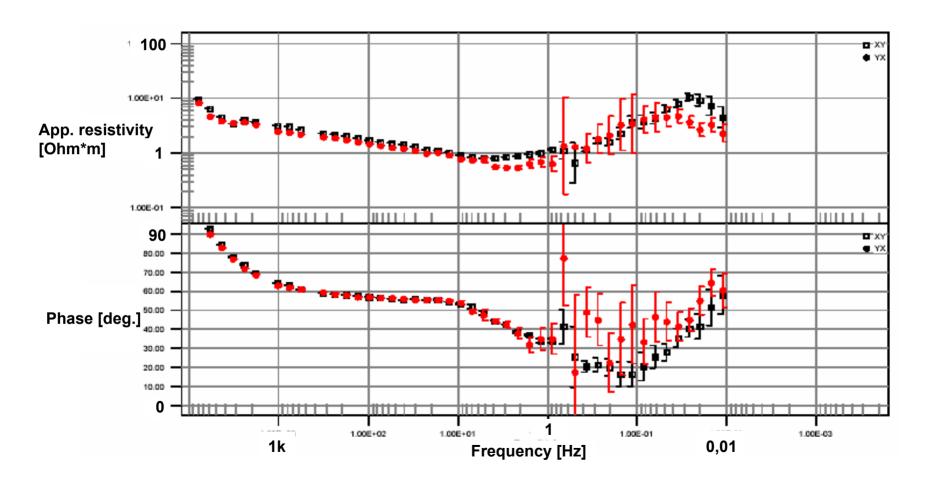
source: Frisch & Meschede 2007

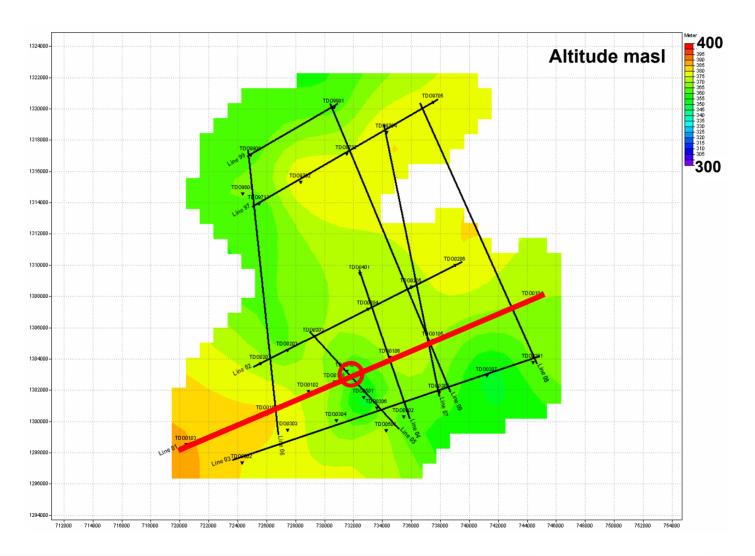


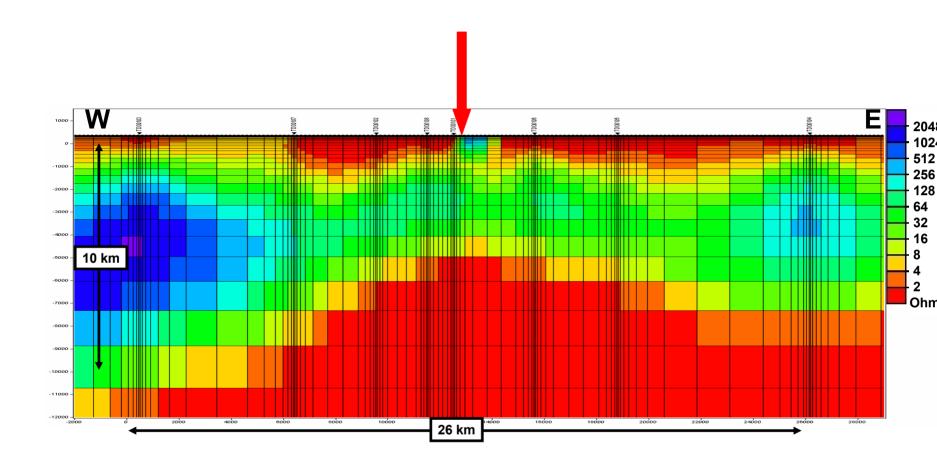




Geothermal manifestations in the survey area

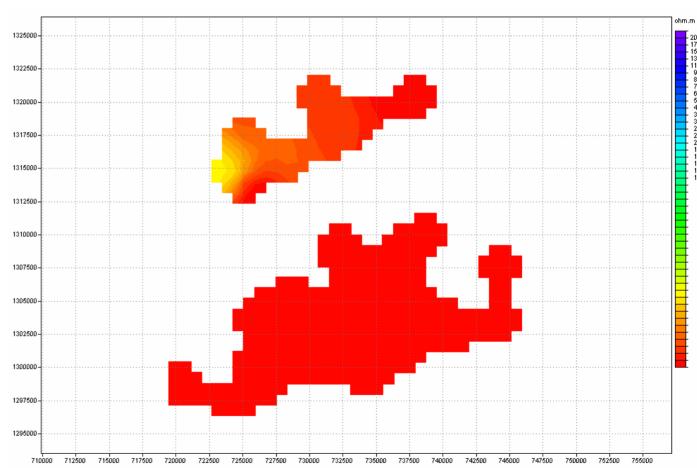






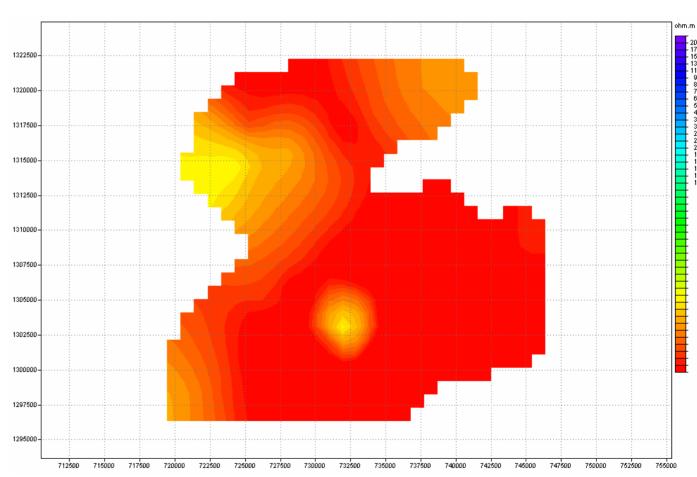
Line 1: resistivity section

Resistivity map at 200 masl



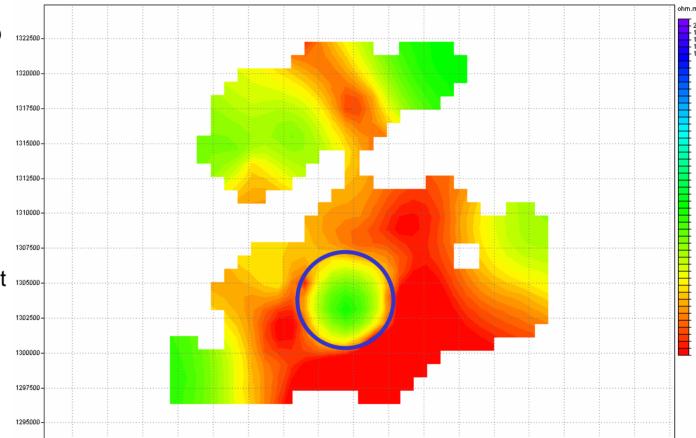


Resistivity map at 300 mbsl





Resistivity map at 1000 mbsl

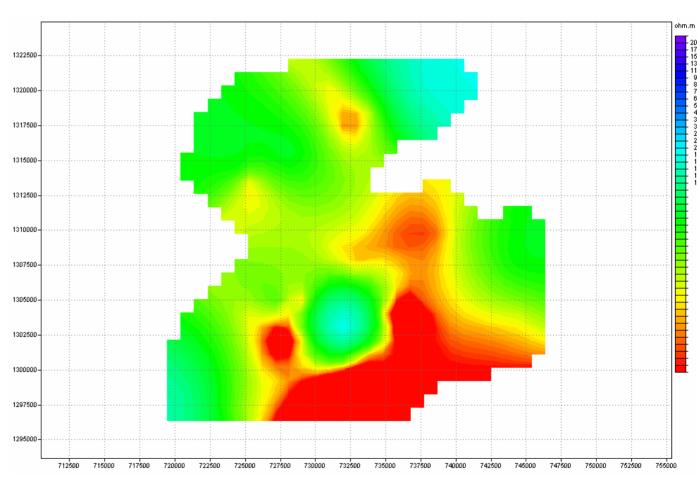


722500

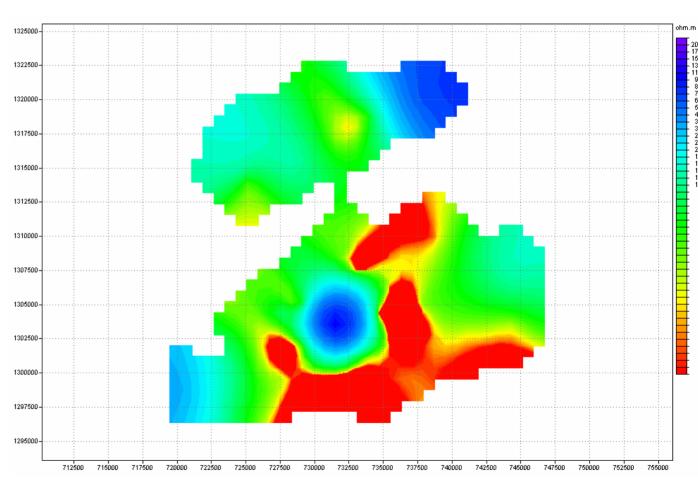
727500

deep reservoir at approx. 1400 m depth?

Resistivity map at 1500 mbsl

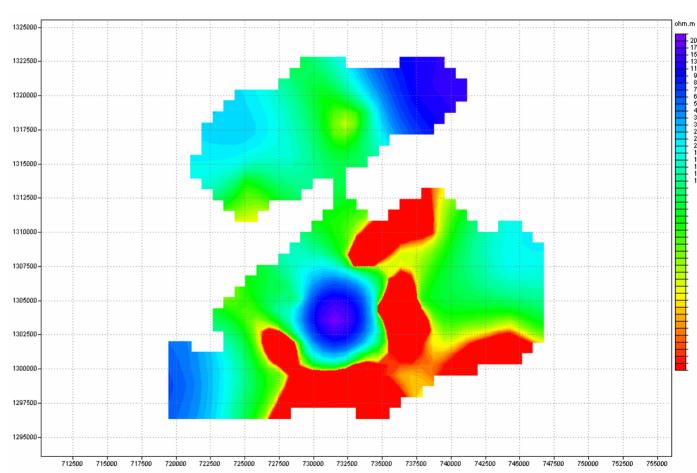


### Resistivity map at 2500 mbsl



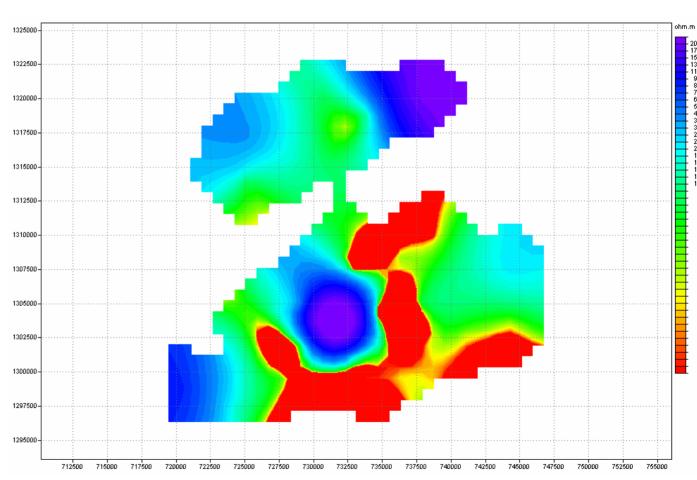


# Resistivity map at 3000 mbsl



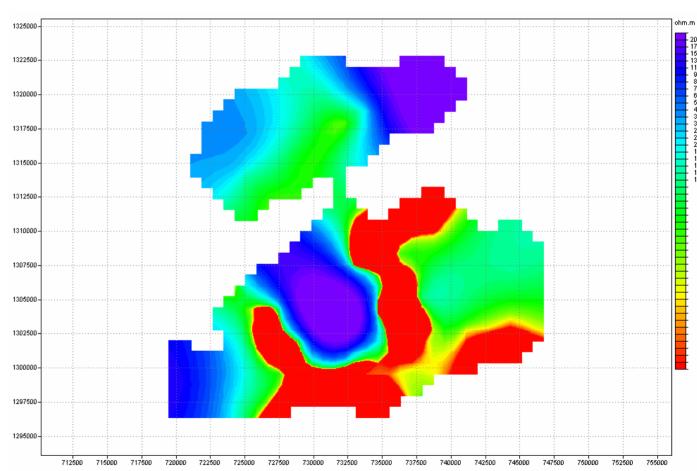


Resistivity map at 4000 mbsl



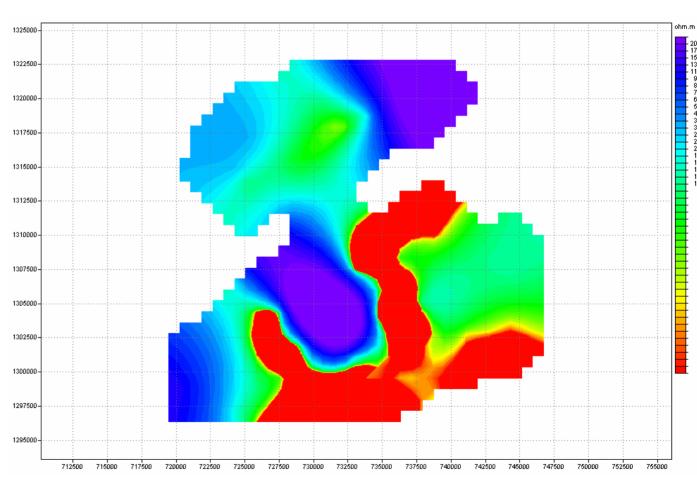


### Resistivity map at 5000 mbsl



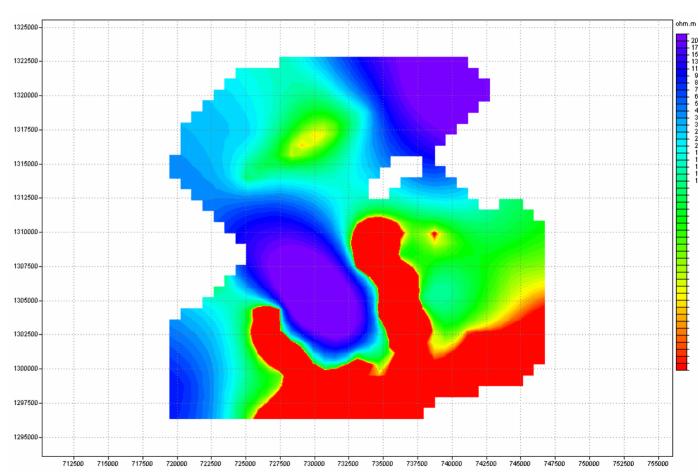


# Resistivity map at 6000 mbsl



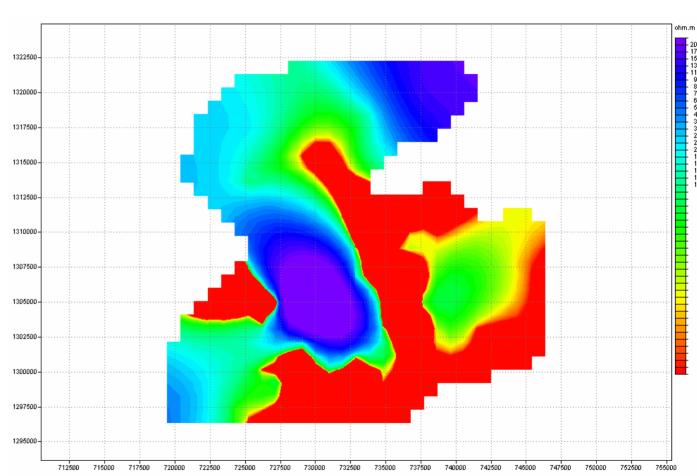


### Resistivity map at 7000 mbsl





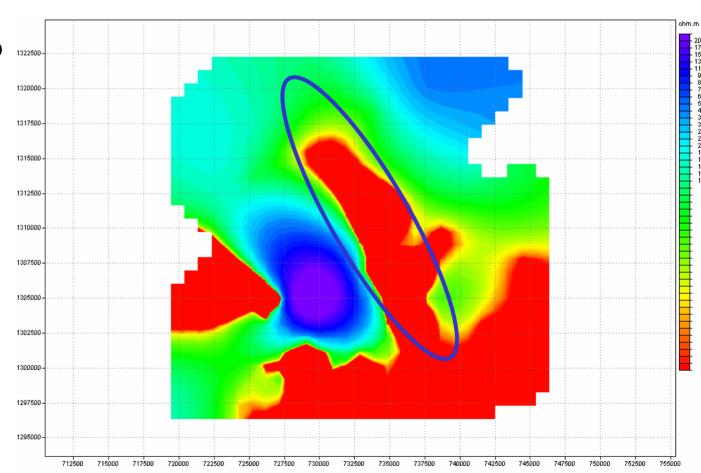
Resistivity map at 8000 mbsl



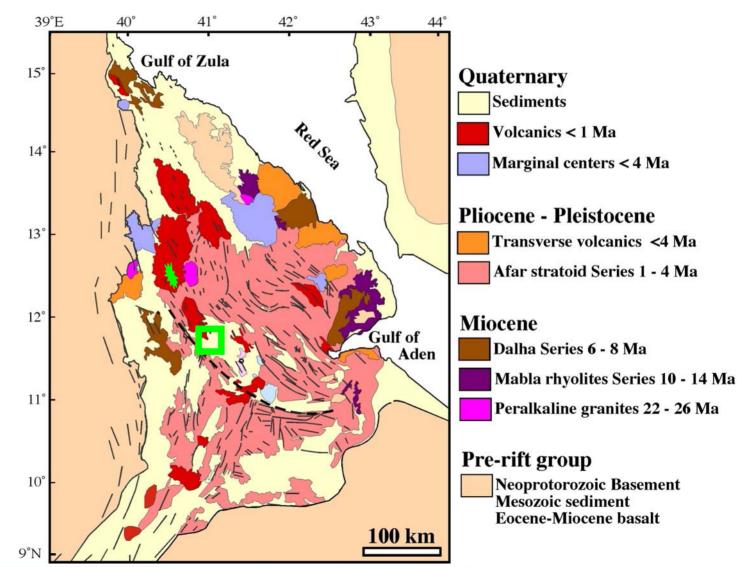


Resistivity map at 9000 mbsl

magma?









Boina vent (Dabbahu rift structure), developed approx.100 km NW of survey area, autumn 2005





